

Surface waves and currents in aquatic vegetation

James C. McWilliams[†]

Department of Atmospheric and Oceanic Sciences, University of California, Los Angeles, CA 90095-1565, USA

(Received 11 July 2022; revised 5 December 2022; accepted 3 January 2023)

A multiscale asymptotic theory is formulated for surface gravity waves and currents in finite-depth water with a vegetation canopy that provides a drag force on both flows with known drag coefficients. It assumes that the density is uniform and the depth is uniform *pro tem* and that the wave frequency is fast compared to the current advective rate. It is a quasi-linear theory in which the wave dynamics is independent of current and drag to leading order but provides perturbative corrections, and in which wave nonlinear interactions are neglected while quadratic wave-averaged wave fluxes and quadratic wave-drag effects are retained. The primary surface wave is modified by drag and current interactions, and the wave-averaged current momentum balance includes a wave-augmented drag force and several vortex forces due to Earth's rotation, current vorticity, Stokes drift and drag-induced wave vorticity. The wave-averaged current equations derived here are a suitable basis for future large-eddy simulation and submesoscale circulation computational models.

Key words: ocean processes, coastal engineering, ocean circulation

1. Introduction

This paper presents a formal theory for surface gravity waves and currents in shallow water with a vegetation canopy that provides a drag force on both. The goal is a wave-averaged oceanic model that can be computationally integrated as a large-eddy simulation (LES) model with a rigid lid or a submesoscale circulation model – the essential distinction is whether boundary-layer turbulence is resolved or parametrized – with a time step much larger than would be required to resolve the primary wave propagation. The purpose is for intermediate-sized problems with heterogeneous currents and vegetation, not for the micro-mechanics of flow around plant fronds (e.g. Nepf 2012).

The paper is intended as a proof-of-concept demonstration of relevant flow–vegetation interaction processes. In particular, the model posits a parametrization of the horizontal

[†] Email address for correspondence: jcm@atmos.ucla.edu

drag force per unit water density associated with a horizontal flow \mathbf{u}^h through a canopy with mostly upright plants,

$$\mathbf{D}^h = -\mathcal{D}(\mathbf{x})|\mathbf{u}^h|\mathbf{u}^h \tag{1.1}$$

(Shaw & Schumann 1992); herein, vectors are in bold italic face, and the superscript h denotes a horizontal vector. The scalar factor,

$$\mathcal{D} = \alpha C_D \geq 0, \tag{1.2}$$

is the product of α , the foliage area density per unit volume (i.e. with units of m^{-3}) and non-dimensional C_D , the drag coefficient appropriate to the vegetation type. In the above, $\alpha(\mathbf{x})$ is proportional to the cross-sectional profile the plants present to the flow. It is usually spatially variable because plant density varies, both in the vertical profile of the canopy and in heterogeneous horizontal distributions (‘patches’), but it is assumed independent of time on the relevant flow scales.

Much of the potential value of the theory developed here is in learning more about how non-uniform vegetation affects coastal currents and *vice versa*. The formula (1.1) assumes there is a turbulent boundary layer adjacent to the vegetation, expressed by its quadratic dependence on the velocity. The presumption is that α and C_D will be determined experimentally for particular plant types and distributions. The formula also neglects any flexible plant motion as the flow varies (cf. Henderson 2019), which is most appropriate for plants with rigid or highly buoyant fronds such as macroalgae (giant kelp). Based on field experiments, Utter & Denny (1996) estimate a value of $C_D \approx 0.015$ for macroalgae, and Monismith *et al.* (2022a,b) analyse other aspects for kelp-forest drag. Luhar & Nepf (2011) discuss how plant compliance under flow alters its configuration, hence α . For a justification for neglecting plant motion for macroalgae, see Yan, McWilliams & Chameki (2021) – but note that there (1.1) is assumed to apply only to the current’s \mathbf{u} , not the total flow including wave orbital motions as here. There is also an extensive literature on atmospheric flows over and through terrestrial canopies (e.g. Finnigan 2000; Brunet 2020). The parametrization formula (1.1) is arguably too simple for many real situations; nevertheless, it has often been invoked in previous studies of canopy drag and it is worthwhile to further explore its consequences *faute de mieux*.

The effects of surface gravity waves on currents are widely investigated and modelled. Outside of the surf zone (as assumed here), a wave-averaged theory for the evolution of the current velocity \mathbf{u} highlights the importance of the Stokes vortex forces,

$$\mathbf{SVF} = \mathbf{u}^{St} \times (f\hat{\mathbf{z}} + \nabla \times \mathbf{u}), \tag{1.3}$$

where \mathbf{u}^{St} is the wave Lagrangian mean flow, the Stokes drift (Craik & Leibovich 1976) – here generalized to include the influence of the Coriolis frequency f due to Earth’s rotation projected on the local vertical direction $\hat{\mathbf{z}}$. Its influence is considered the primary explanation for the occurrence of Langmuir circulations in the surface boundary layer, as confirmed in many LES solutions (reviewed in Sullivan & McWilliams 2010). More recently, the importance of Stokes drift has been demonstrated in submesoscale circulation models (Hypolite *et al.* 2022). This canopy-interaction theory is a suitable counterpart to what has become the prevailing wave–current interaction theory (also known as Stokes vortex force, originating with Craik & Leibovich (1976)) used in wave-averaged LES and circulation models.

The wave-averaged theory is of the quasi-linear type. In its previous forms, it is based on two small parameters,

$$\epsilon = ak \ll 1 \quad \text{and} \quad \mu = \frac{V}{C} \ll 1. \quad (1.4a,b)$$

The first assumption says that the wave's free surface elevation has a small slope, where a is the elevation amplitude and k is the horizontal wavenumber. The second assumption says that the current speed V is small compared to the wave propagation speed, $C = \sigma/k$, where σ is the wave frequency; this implies a time scale separation between faster waves and slower currents and thus enables wave averaging for the current dynamics. Taken together, they motivate the ingredients of a quasi-linear theory: linearized wave dynamics with no current influence at leading order in μ and no nonlinear wave interactions, and wave-averaged quadratic wave fluxes (e.g. SVF) in the current dynamics that incorporate the current effect on the waves at $O(\mu)$.

Various extensions of quasi-linear wave-averaged theory have been made (Holm 1996; McWilliams, Restrepo & Lane 2004; Suzuki & Fox-Kemper 2016), including a recent generalization for higher-order effects in μ in u^{St} , SVF , etc. (McWilliams 2022, hereafter WCI22); however, only the leading-order current interactions are included. The latter paper (WCI22) contains an extended discussion of many aspects of a quasi-linear wave-current interaction theory, and specific references to it are made below to abbreviate the present paper. WCI22 does not include anything about canopy drag, and the wave and current relations here are for a finite-depth ocean (where vegetation mostly is found) rather than the deep-water (infinite depth) in WCI22.

This paper follows this quasi-linear derivation path by additionally including the canopy drag effects in (1.1) while again taking a perturbation approach with an assumption of weak drag, i.e.

$$v = aD_* \ll 1, \quad (1.5)$$

where D_* is a representative value for $\mathcal{D}(x)$. Canopy drag yields several important effects: dissipation of the waves, a wave-enhanced drag on the currents, and an augmentation of VF by the wave vorticity generated through drag. The ideas of wave dissipation and wave-enhanced current drag by vegetation have been addressed previously (e.g. Dalrymple, Kirby & Hwang 1984; Rosman *et al.* 2013; Hu *et al.* 2014; Losada, Maza & Lara 2016; Lei & Nepf 2019; Zhao *et al.* 2021; Yu, Rosman & Hench 2022), often using laboratory experiments or direct numerical simulations and with an emphasis on the wave damping; the reliance on an experimentally determined C_D is common. The particular formulae derived here for the partitioned wave and current drag forces in D^h are somewhat different than previously, and the augmented VF result is new. The theory assumes that there is a primary wave (or multiple primary waves; see § 12 and Appendix B) that is perturbatively modified by canopy drag and weaker currents.

The outline of the paper is as follows. The basic fluid model, primary approximations and wave-current decomposition are in § 2. The leading-order wave solution is in § 4. The perturbation effects of canopy drag and currents on the waves are in §§ 5–8 and Appendix A, and the combined drag and current effects of the perturbed waves on the currents are in §§ 9 and 10. Alternative situations of weaker or multiple primary waves are analysed in §§ 11 and 12. A summary and discussion of future prospects is in § 13. Finally, because of the complexity of a multi-parameter, multiscale theoretical formalism, a table of symbols is included for reference (table 1).

Symbol	Definition
g	gravitational acceleration
f	Coriolis frequency
H	resting water depth
$(a, \mathbf{k}, \sigma, C)$	primary wave amplitude, wavenumber vector, eigenfrequency and phase speed
$F(z)$	primary wave vertical eigenfunction
C, V	scale estimates for wave propagation speed and current velocity
ϵ	wave steepness parameter (1.4a,b)
μ	current amplitude parameter (1.4a,b)
$\mathbf{x} = (x, y, z, t)$	space and time coordinates
$\mathbf{x}^h = (x, y)$	horizontal coordinates
\bullet^h	horizontal vector
$\mathbf{u} = (u, v, w), \mathbf{u}^h = (u, v)$	3-D and horizontal velocity
∇, ∇^h	3-D and horizontal spatial gradient operators
ζ	vorticity vector
ϕ	geopotential function (normalized pressure)
η	surface elevation anomaly
\mathbf{u}^{St}	Stokes drift velocity (4.5)
La	Langmuir number (§ 3)
$VF = SVF + CVF$	wave-averaged total = Stokes + canopy vortex forces, (1.3) + (10.5)
D^h	canopy drag force (1.1)
$\mathcal{D}(\mathbf{x}), \alpha(\mathbf{x}), C_D$	drag coefficients (1.2)
v	drag amplitude parameter (1.5)
$\bar{\bullet}$ or $\langle \bullet \rangle, \bullet'$	wave averages and wave fluctuation components
$\mathcal{E} = \exp[i\Theta]$	primary wave phase functions (4.1)
c.c., \bullet^*	complex conjugate
D_w^h, D'_w, \bar{D}_w^h	drag forces associated with wave orbital motion (§ 5)
\bar{D}^h	wave-averaged drag force on currents (9.3)
$\tau = (\mu, \nu)t$	'slow' time coordinates of drag- and current-induced evolution
$\bullet'_w, \bullet'_d, \bullet'_c, \bullet'_m$	subscripts for the primary wave, drag, current and combined wave components
$\bullet'_\zeta, \bullet'_\phi$	vortical and potential wave field decomposition components
Φ', Q	wave correction potential functions (5.11), (6.5) and (7.2)
R_i, R_0	forcing terms for Φ' problem
\mathcal{P}	projection operator for solvability condition (8.3)
J, r	slow-time drag-induced wave decay rates (8.8a,b)
\bar{B}	wave-averaged Bernoulli head
\bar{C}	potential function associated with wave-averaged vortex force (10.4)
H, I, K	phase functions for drag force (Appendix A)
Δ	non-dimensional drag force function (Appendix B)

Table 1. Symbol definitions.

2. Basic model and wave-current decomposition

The theory in this paper analyses wave-current-drag interactions while making the unrealistic simplifications of constant water density, a flat bottom and an absence of non-conservative forces apart from the canopy drag; such generalizations can be pursued later.

Thus, the incompressible fluid equations are

$$\left. \begin{aligned} \partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla) \mathbf{u} + f \hat{\mathbf{z}} \times \mathbf{u} &= -\nabla \phi + \mathbf{D}^h, \\ \nabla \cdot \mathbf{u} &= 0, \end{aligned} \right\} \quad (2.1)$$

where $\mathbf{u} = (\mathbf{u}^h, w) = (u, v, w)$ is the velocity, ϕ is the density-normalized pressure as a deviation from the background hydrostatic profile $-gz$ that does not accelerate the flow (g is the gravitational constant), f is the Coriolis frequency (assumed constant) and \mathbf{D}^h is the drag force (1.1).

For boundary conditions, all fields are periodic in the horizontal coordinates (x, y) , and in the vertical,

$$\left. \begin{aligned} w &= \partial_t \eta + \mathbf{u}^h \cdot \nabla^h \eta, \\ \phi &= g\eta \end{aligned} \right\} \quad (2.2)$$

at the sea surface, $z = \eta(x, y, t)$, and

$$w = 0 \quad (2.3)$$

at the bottom, $z = -H$ (with H a constant). The second condition in (2.2) is a statement that the total pressure is uniform at the surface interface. (Boundary stresses are omitted for now.)

Direct numerical solutions of (2.1)–(2.3) are difficult for general waves and currents especially because of the disparate time scales of wave propagation and current evolution and the tendency for η to develop wave-related singularities (i.e. to break). So, to obtain a wave-averaged LES-type or circulation model for the currents, several approximations are made. These are listed in (1.4a,b) and (1.5). In particular, they imply that the wave solutions at leading order are linear and uninfluenced by currents and drag (§ 4).

The relations in this paper are all dimensional, but scaling estimates are made (§ 3) based on non-dimensionalizations using the primary wave scales: $a, \mathbf{x} \sim 1/k$ (with the current length scale ℓ assumed to be the same or larger), $t \sim 1/\sigma$, $\mathbf{u} \sim C = \sigma/k$, etc.

The velocity and other fields are decomposed into currents and waves, e.g.

$$\mathbf{u} = \bar{\mathbf{u}} + \mathbf{u}', \quad (2.4)$$

with the distinction that \mathbf{u}' has zero average over a wave period, $2\pi/\sigma$. A further decomposition is made for the wave velocity,

$$\mathbf{u}' = \mathbf{u}'_w + \mathbf{u}'_c + \mathbf{u}'_d. \quad (2.5)$$

Here \mathbf{u}'_w is the primary wave field with linear dynamics, its magnitude being $O(\epsilon C)$; \mathbf{u}'_c is the small modification of the waves due to the current effects on the waves (CEW; § 6), its magnitude being $O(\epsilon \mu C)$; and \mathbf{u}'_d is the small modification of the waves due to the transient drag force $\mathbf{D}^{h'}$ (§ 5), its magnitude being $O(\epsilon \nu C)$. The respective wave corrections due to drag and currents in \mathbf{u}'_d and \mathbf{u}'_c can be superimposed because their determining equations are linear (§§ 5 and 6). No decomposition is imposed on $\bar{\mathbf{u}}$ ($\sim \mu C$); the goal is to derive a wave-averaged dynamical balance equation for $\bar{\mathbf{u}}$ including all of these wave effects on currents (WEC; § 10).

3. Scaling assumptions and rationale

The primary scaling assumptions of $\epsilon, \mu, \nu \ll 1$ are stated in § 1, and the wave and current velocities have scaling estimates of

$$\mathbf{u}' \sim \epsilon C, \quad \bar{\mathbf{u}} \sim \mu C, \quad (3.1a,b)$$

respectively. For strict validity of a quasi-linear perturbation theory, ϵ must be small compared to μ or ν , which often is not true in nature. Wind–wave equilibrium in the surface layer with Ekman currents scales as $\mu \sim \epsilon^2$; i.e. $\mu \ll \epsilon$ with wave orbital motions

($\sim \epsilon C$) stronger than currents ($\sim \mu C$). Hanley, Belcher & Sullivan (2010) show the empirical distribution of the so-called turbulent Langmuir number, $La = \sqrt{u_* / |\mathbf{u}^{St}|(0)} \sim \mu^{1/2} / \epsilon$, which is peaked near the equilibrium value of 0.3. Stronger currents would increase the μ / ϵ ratio and big swells decrease it. With the Langmuir turbulence context in mind, the primary target regime is where the wave orbital motion is faster than the current speed, which itself is often comparable to the Stokes drift:

$$\mathbf{u}'_w \gg \bar{\mathbf{u}} \sim \mathbf{u}^{St}. \tag{3.2}$$

This ordering allows simpler formulae for the canopy drag on the waves and for the wave-induced canopy drag on the currents (§§ 5 and 9). Furthermore, for generality of the results, we assume that drag effects are comparable to current effects, both on the wave perturbations (§§ 5–7) and on the currents (§ 10); i.e. $\nu \sim \mu \ll 1$.

The theoretical focus in wave–current interaction is usually on the interactions associated with the spectrum-peak waves (i.e. at most a few wavenumber components that have the largest amplitudes a in η). For these waves, the nonlinear evolution effects are often weak, even if the full wave spectrum has stronger nonlinear interactions. The success of quasi-linear theory and even the simplification to monochromatic primary waves as a modelling framework for Langmuir turbulence (McWilliams, Sullivan & Moeng 1997) provides motivation to continue this approach with the inclusion of canopy drag.

Once a commitment is made to a quasi-linear theory for wave–current interaction, in part because it is a doable theory for an otherwise very difficult problem, then the magnitude of ϵ need not be declared. However, with the nonlinear canopy drag law (1.1), the ratio μ / ϵ is important for how the drag force is approximated and partitioned between waves and currents. In this paper the choice is made to assume that wave orbital velocities are stronger than currents – the most common natural situation – as the main line of argument, with tangential remarks about the alternatives (§ 9).

4. Primary wave

For simplicity, the primary wave field is assumed to be monochromatic, i.e. a single eigenmode for linear wave dynamics (but also see § 11 and Appendix B). We adopt a complex notation; e.g. for the surface elevation,

$$\eta'_w = a \frac{\mathcal{E}}{2} + \text{c.c.} \tag{4.1}$$

Here a is the amplitude, $\mathcal{E} = \exp[i\Theta]$ is the phase function, $\Theta = \mathbf{k} \cdot \mathbf{x}^h - \sigma t$ is its horizontal-propagation argument, \mathbf{k} is the horizontal wavenumber vector (with magnitude k), σ is the frequency eigenvalue and c.c. denotes the complex conjugate of the preceding expression. In finite-depth water ($kH < \infty$), the dispersion relation is

$$\sigma = \sqrt{gk \tanh[kH]} > 0. \tag{4.2}$$

The propagation is in the direction of \mathbf{k} . The other eigenfunction fields are

$$\left. \begin{aligned} \phi'_w &= gaF(z) \frac{\mathcal{E}}{2} + \text{c.c.}, \\ \mathbf{u}'_w &= \frac{ga}{\sigma} F(z) \mathbf{k} \frac{\mathcal{E}}{2} + \text{c.c.}, \\ w'_w &= -\frac{iagk}{\sigma} G(z) \frac{\mathcal{E}}{2} + \text{c.c.} \end{aligned} \right\} \tag{4.3}$$

The non-dimensional vertical structure functions are

$$\left. \begin{aligned} F(z) &= \cosh[kz] + \tanh[kH] \sinh[kz], \\ G(z) &= \sinh[kz] + \tanh[kH] \cosh[kz]. \end{aligned} \right\} \quad (4.4)$$

Functions F and G decay monotonically with depth. At $z = -H$, $G = 0$ while $F \neq 0$. If a is real, then η'_w , ϕ'_w and \mathbf{u}'_w are all $\propto \cos[\Theta]$, and $w'_w \propto \sin[\Theta]$. Furthermore, there is no vorticity in this eigenmode, $\boldsymbol{\zeta}'_w = \nabla \times \mathbf{u}'_w = 0$, and this is quite helpful in solving for the drag and current wave corrections in §§ 5 and 6.

In a multiscale theory, as here, a can be a function of ‘slow’ time τ due to drag and current effects (§§ 5–8), in which case $a(\tau)$ can become complex due to a slow evolution in the primary wave phase. In this paper, for simplicity, a is assumed to be real and positive in discussing specific canopy-related results (§§ 5 and 9–11 and Appendices A and B), while its more general complex form is retained for the current interactions (§§ 6 and 7). The latter allows for ‘wave–current resonance’ that might also incorporate canopy effects, but that is not a central focus in this paper; also see the extended discussion of wave–current resonance and several examples in WCI22, chapters 7 and 8.

An important property of surface waves is their Lagrangian mean flow, the Stokes drift velocity. For small ϵ , this is defined by

$$\mathbf{u}^{St} = \left\langle \left(\int^t \mathbf{u}' dt \cdot \nabla \right) \mathbf{u}' \right\rangle, \quad (4.5)$$

where the angle brackets denote an average over the wave period (i.e. this has the same meaning as the overbar symbol). In this paper only the leading-order approximation to \mathbf{u}^{St} is needed. With $\mathbf{u}' \approx \mathbf{u}'_w$,

$$\left. \begin{aligned} \mathbf{u}^{hSt}(z) &\approx \frac{|a|^2 \sigma \mathbf{k}}{2 \sinh^2[kH]} \cosh[2k(z+H)], \\ w^{St}(z) &\approx 0. \end{aligned} \right\} \quad (4.6)$$

The drift is only in the horizontal direction. Also, $\nabla \cdot \mathbf{u}^{St} = 0$, as is usual in conservative quasi-linear wave theories (WCI22, appendix C).

For completeness, these primary wave formulae are also listed in the deep-wave limit, $kH \rightarrow \infty$ (even though that is inconsistent with bottom-rooted marine vegetation affected by surface wave motions):

$$\sigma \rightarrow \sqrt{gk}, \quad F(z), G(z) \rightarrow e^{kz}, \quad \frac{\cosh[2k(z+H)]}{2 \sinh^2[kH]} \rightarrow e^{2kz}. \quad (4.7a-c)$$

Everything in this section is well known.

5. Effect of drag on the waves

With the decomposition (2.5) and the assumption that wave orbital motions are stronger than currents (3.2), the largest canopy drag term is

$$\mathbf{D}_w^h = -\mathcal{D} |\mathbf{u}_w^{h'}| \mathbf{u}_w^{h'}. \quad (5.1)$$

Substituting from (4.3) with real $a > 0$,

$$\mathbf{D}_w^h = -\mathcal{D} \left(\frac{(ga)^2 F^2 k}{\sigma^2} \right) |\cos[\Theta]| \cos[\Theta] \mathbf{k}. \quad (5.2)$$

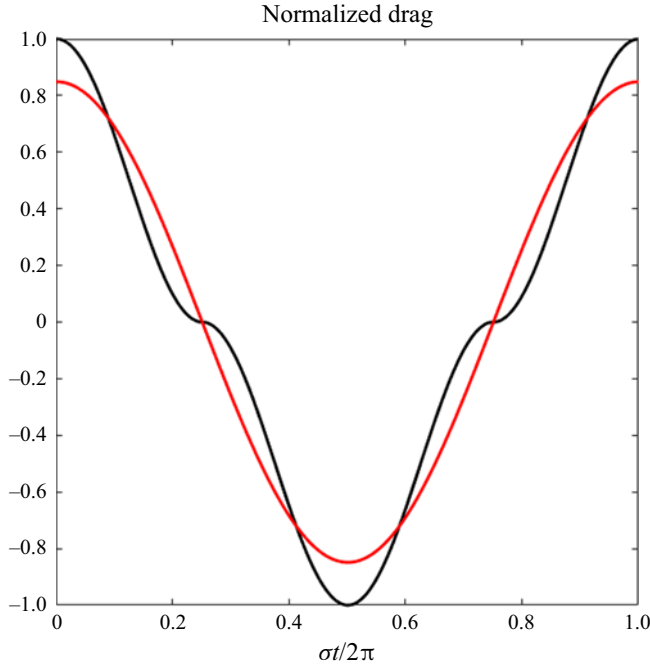


Figure 1. Plots of $|\cos[\Theta']| \cos[\Theta']$ (black) and $c_1 \cos[\Theta']$ (red) over one wave period in $\Theta' = \sigma t/2\pi$.

The functional dependence on Θ can be approximated as

$$\langle |\cos[\Theta]| \cos[\Theta] \rangle = 0, \quad |\cos[\Theta]| \cos[\Theta] \approx c_1 \cos[\Theta], \quad (5.3a,b)$$

as demonstrated in figure 1. The coefficient $c_1 = 8/3\pi \approx 0.849$ is the projection of the left-hand side of (5.2) on $\cos[\Theta]$, i.e. it is the value that gives the least-square error in this approximation averaged over a primary wave period; the root-mean-square (r.m.s.) relative error in the approximation is $<20\%$. Thus, the wave forcing by canopy drag is purely oscillatory, with $\bar{D}_w^h = 0$, and it is approximately monochromatic. The latter allows more extensive analytic results compared to the exact form of D_w^h ; see Appendices A and B for generalizations.

The wave drag has the same phase function as the eigenmode, *viz.*

$$D_w^{h'} \approx -D c_1 \left(\frac{(ga)^2 F^2 k}{\sigma^2} \right) \cos[\Theta] \mathbf{k}; \quad (5.4)$$

and \mathbf{u}'_d is needed for the WEC model (§ 10). With this approximation for $D_w^{h'}$ as the forcing term for the $O(\nu)$ correction to \mathbf{u}'_w , this makes the solution for \mathbf{u}'_d itself monochromatic. Obviously, a more complete solution with $D^{h'}$ could be obtained with higher harmonics of Θ , and Appendix A shows that the errors incurred by this simplification are minor.

The solution procedure for \mathbf{u}'_d (as for \mathbf{u}'_c in § 6 and in WCI22, chapter 5) is decomposed into vortical and irrotational parts:

$$\mathbf{u}'_d = \mathbf{u}'_d{}^\zeta + \mathbf{u}'_d{}^\phi, \quad (5.5)$$

where $\nabla \times \mathbf{u}'_d{}^\phi = 0$. As a consequence of $\zeta'_w = 0$, the vortical component is relatively easy to obtain, while the irrotational component may be more difficult.

The wave vorticity equation with respect to drag at $O(\nu)$ is

$$\partial_t \zeta'_d = \text{curl}[\partial_t \mathbf{u}'_d] = \text{curl}[\mathbf{D}'_w]. \tag{5.6}$$

There is no vertical component in the drag force (1.1), and

$$\text{curl}[\mathbf{D}'_w] = c_1 \frac{g^2 a^2 k}{\sigma^2} \begin{pmatrix} k^y \partial_z [\mathcal{D}F^2] \\ -k^x \partial_z [\mathcal{D}F^2] \\ F^2 (\hat{\mathbf{z}} \cdot \mathbf{k} \times \nabla^h \mathcal{D}) \end{pmatrix} \cos[\Theta]. \tag{5.7}$$

To invert (5.6), the time integral of the phase function is

$$\int^t \cos[\Theta] dt = -\frac{1}{\sigma} \sin[\Theta]. \tag{5.8}$$

Therefore, using (5.4)–(5.7),

$$\zeta'_d = c_1 \frac{g^2 a^2 k}{\sigma^3} \begin{pmatrix} -k^y \partial_z [\mathcal{D}F^2] \\ +k^x \partial_z [\mathcal{D}F^2] \\ -F^2 (\hat{\mathbf{z}} \cdot \mathbf{k} \times \nabla^h \mathcal{D}) \end{pmatrix} \sin[\Theta]. \tag{5.9}$$

Furthermore, the vorticity relation for $\zeta'_d = \nabla \times \mathbf{u}'_d$ in (5.9) can be inverted for the vortical component of the wave velocity to obtain

$$\mathbf{u}'_d{}^{h\zeta'} = \mathcal{D}c_1 \left(\frac{(ga)^2 F^2 k}{\sigma^3} \right) \sin[\Theta] \mathbf{k}, \quad w'_d{}^{\zeta'} = 0. \tag{5.10}$$

Notice that the vertical component of ζ'_d is zero if $\nabla^h \mathcal{D} = 0$, i.e. if the vegetation is horizontally uniform.

The irrotational component $\mathbf{u}'_d{}^{\Phi'}$ in (5.5) requires the solution of an elliptic boundary-value problem. It is derived by taking the divergence of the momentum equation at $O(\nu)$ and utilizing a combination of the vertical momentum equation and the conditions (2.2) and (2.3) at the boundaries to eliminate the velocity in favour of a velocity potential function Φ'_d defined by

$$\mathbf{u}'_d{}^{\Phi'} = -\nabla \int^t \Phi'_d dt \tag{5.11}$$

(WCI22, chapter 5). The derivation includes recognizing that $\nabla \cdot \partial_t \mathbf{u}'_d = 0$ and $\zeta'_w = 0$. It also includes the linearized wave forms of (2.2) and (2.3) in a Taylor series expansion about $z = 0$:

$$\left. \begin{aligned} w' &= \partial_t \eta' + \nabla^h \cdot (\eta' \bar{\mathbf{u}}^h) \quad \text{and} \quad \phi' = g\eta' \quad \text{at } z = 0, \\ w' &= 0 \quad \text{at } z = -H. \end{aligned} \right\} \tag{5.12}$$

For this drag-effect derivation, the $\bar{\mathbf{u}}^h$ boundary-condition term is not considered, but it will be included for the CEW component $\mathbf{u}'_c{}^{\Phi'}$ in §6. Also, the wave vertical momentum equation has no drag component (i.e. $\hat{\mathbf{z}} \cdot \mathbf{D}'_w = 0$). The combination of the relations in this

paragraph is the following partial differential equation (PDE) system:

$$\left. \begin{aligned} \nabla^2 \Phi'_d &= \text{div}[\mathbf{D}_w^{h'}], \\ \left[\partial_z - \frac{\sigma^2}{g} \right] \Phi'_d &= -\frac{2}{g} \partial_t \partial_\tau \phi'_w \quad \text{at } z = 0, \\ \partial_z \Phi'_d &= 0 \quad \text{at } z = -H. \end{aligned} \right\} \quad (5.13)$$

The ‘forcing’ term in (5.13) includes the possibility of a ‘slow-time’ evolution of the primary wave field (4.1)–(4.3) through its amplitude, $a(\tau)$, i.e.

$$-\frac{2}{g} \partial_t \partial_\tau \phi'_w = -2\sigma \sin[\Theta] \partial_\tau a \quad (5.14)$$

for a slow-time coordinate, $\tau = \nu t$. This slow-time dependence is determined through a solvability condition for (5.13) that assures there is no resonant forcing of Φ'_d , i.e. that there is no projection of the forcing on the primary wave eigenmode (WC122, appendix E).

It will be shown that canopy drag induces a slow amplitude decay for the primary wave, and this condition must be combined with an analogous constraint on Φ'_c , hence providing an amplitude dependence also on $\bar{\mathbf{u}}$; both of these statements will be further explained in §§ 6–8. The system (5.13) must be solved in conjunction with the evolution of the current $\bar{\mathbf{u}}$ also on the slow time scale τ (§ 10).

With (5.4) the drag forcing in the PDE in (5.13) is

$$\begin{aligned} \text{div}[\mathbf{D}_w^{h'}] &= -c_1 \left(\frac{(gaF)^2 k}{\sigma^2} \right) (\mathbf{k} \cdot \nabla^h) [\mathcal{D} \cos[\Theta]] \\ &= c_1 \left(\frac{(gaF)^2 k}{\sigma^2} \right) \{k^2 \mathcal{D} \sin[\Theta] - (\mathbf{k} \cdot \nabla^h \mathcal{D}) \cos[\Theta]\}. \end{aligned} \quad (5.15)$$

In the particular case of horizontally uniform vegetation, $\nabla^h \mathcal{D} = 0$, and we can factor Φ'_d for its (\mathbf{x}^h, z, t) dependences,

$$\Phi'_d = \varphi_d(z) \sin[\Theta], \quad (5.16)$$

where φ_d satisfies a one-dimensional (1-D) boundary-value problem from (5.13):

$$\left. \begin{aligned} [\partial_z^2 - k^2] \varphi_d &= \mathcal{D}(z) c_1 \frac{g^2 a^2 k^3}{\sigma^2} F^2(z), \\ \left[\partial_z - \frac{\sigma^2}{g} \right] \varphi_d &= -2\sigma \partial_\tau a \quad \text{at } z = 0, \\ \partial_z \varphi_d &= 0 \quad \text{at } z = -H. \end{aligned} \right\} \quad (5.17)$$

Furthermore, this 1-D problem only needs to be solved once at each slow time τ as $a(\tau)$ evolves. Or, even more efficiently, it can be partitioned into two problems that only need to be solved once – one with only ordinary differential equation forcing by the interior factor of a^2 , the other with only surface boundary forcing by the factor of $\partial_\tau a$ – then the two answers are superposed with the current values of $a(\tau)$ and $\partial_\tau a(\tau)$ as their factors.

6. Current effects on the waves (CEW)

In WCI22 (chapter 5) the quasi-linear theory is presented for CEW in u'_c , and the relevant parts of it will be repeated briefly. The wave momentum equation is linearized with respect to wave–current advection, i.e. CEW occurs at $O(\mu)$ compared to the primary wave balance with respect to the current effects, and (2.1) is approximated as

$$\left. \begin{aligned} \partial_t u'_c + f \hat{z} \times u'_c &= -\nabla \phi'_c - \nabla(u'_w \cdot \bar{u}) + u'_w \times \bar{\zeta}, \\ \nabla \cdot u'_c &= 0, \end{aligned} \right\} \quad (6.1)$$

where $\bar{\zeta}$ is the current vorticity. Besides the pressure gradient, the right-hand-side wave forces are the gradient of the wave Bernoulli head and the wave vortex force. The associated vertical boundary conditions on u'_c , ϕ'_c and η'_c are (5.12) now with a non-zero \bar{u}^h . Because the wave components for both drag and current satisfy linear equations, they can be superposed in u' .

In this and the next sections a is complex (see § 4).

As before in (5.5), the wave component u'_c is decomposed into vortical and divergent components,

$$u'_c = u_c^{\zeta'} + u_c^{\Phi'}. \quad (6.2)$$

Taking the results from WCI22 (section 5.2), with modifications for finite kH , the wave vorticity associated with (6.1) is

$$\begin{aligned} \zeta'_c &= \nabla \times \left(\int^t u'_w dt \right) \times (f \hat{z} + \bar{\zeta}) \\ &= \nabla \times \frac{ag}{\sigma^2} \{ iF(k \times (f + \bar{\zeta}^z) \hat{z}) + Gk(\hat{z} \times \bar{\zeta}^h) \} \frac{\mathcal{E}}{2} + \text{c.c.} \end{aligned} \quad (6.3)$$

The gradient cross-product in this relation can be inverted to yield the vortical velocity component $u_c^{\zeta'} = (\nabla \times)^{-1} \zeta'_c$, disregarding a possible contribution to the divergent component. The result is

$$u_c^{\zeta'} = \frac{ga}{\sigma^2} \begin{pmatrix} ik^y(f + \bar{\zeta}^z)F(z) - k\bar{\zeta}^y G(z) \\ k\bar{\zeta}^x G(z) - ik^x(f + \bar{\zeta}^z)F(z) \\ i(k^x \bar{\zeta}^y - k^y \bar{\zeta}^x)F(z) \end{pmatrix} \frac{\mathcal{E}}{2} + \text{c.c.} \quad (6.4)$$

As with $u_d^{\Phi'}$ in § 5, the irrotational (divergent) velocity component is the gradient of a potential function,

$$u_c^{\Phi'} = -\nabla \int^t \Phi'_c dt, \quad (6.5)$$

where $\Phi'_c = \phi'_c + u'_w \cdot \bar{u}$ is a combined pressure and Bernoulli-head potential. Following the same derivation path as outlined in § 5, the associated boundary-value problem for Φ'_c

is the following:

$$\left. \begin{aligned}
 \nabla^2 \Phi'_c &= -\nabla \left(\int^t \phi'_w dt \right) \cdot \nabla^2 \bar{\mathbf{u}} \\
 &= \frac{ag}{\sigma} (F(z) \mathbf{k} \cdot \nabla^2 \bar{\mathbf{u}}^h - iG(z)k\nabla^2 \bar{w}) \frac{\mathcal{E}}{2} + \text{c.c.}, \\
 \left[\partial_z - \frac{\sigma^2}{g} \right] \Phi'_c &= -\frac{2}{g} \partial_t \partial_\tau \phi'_w - \frac{2}{g} \nabla^h (\partial_t \phi'_w) \cdot \bar{\mathbf{u}}^h - \nabla^h \left(\int^t \phi'_w dt \right) \cdot \partial_z \bar{\mathbf{u}}^h \\
 &\quad - \frac{1}{g} (\partial_t \phi'_w) (\nabla^h \cdot \bar{\mathbf{u}}^h) \quad \text{at } z = 0 \\
 &= \left(2i\sigma \partial_\tau a + a\sigma \left(-2\mathbf{k} \cdot \bar{\mathbf{u}}^h + \frac{g}{\sigma^2} \mathbf{k} \cdot \partial_z \bar{\mathbf{u}}^h + i\nabla^h \cdot \bar{\mathbf{u}}^h \right) \right) \frac{\mathcal{E}}{2} + \text{c.c.} \\
 &\quad \text{at } z = 0, \\
 \partial_z \Phi'_c &= \hat{\mathbf{z}} \cdot \mathbf{u}'_w \times \bar{\boldsymbol{\xi}}^h = \frac{a\sigma \cosh[2kH]}{2k \sinh[kH]} \mathbf{k} \cdot \partial_z \bar{\mathbf{u}}^h \frac{\mathcal{E}}{2} + \text{c.c.} \quad \text{at } z = -H
 \end{aligned} \right\} \tag{6.6}$$

(cf. WCI22, section 5.2). This system is functionally consistent with (5.13) for Φ'_d except for the replacement of drag-effect forcing there with current-interaction forcing and the presumption that the slow-time coordinate is $\tau = \mu t$. As in § 5 for \mathbf{u}'_d , the elliptical system (6.6) is to be solved as $\bar{\mathbf{u}}$ evolves, along with the evaluation of (6.4), to obtain \mathbf{u}'_c using (6.2) and (6.5).

If the currents themselves are three-dimensional (3-D), then (6.6) cannot be reduced to a 1-D vertical problem like (5.17), but they can be if $\bar{\mathbf{u}} = \bar{\mathbf{u}}^h(z)$ only, in which case the familiar Doppler-shifting effect on the frequency and wave vertical profile is recovered (WCI22, appendix J.1).

7. Synthesis of the wave perturbations

In §§ 5 and 6 the effects on the wave modifications due to wave drag and CEW are expressed in functionally similar ways. To make a more generally useful theory, it seems worthwhile to combine them, taking advantage of the fact that the associated \mathbf{u}' fields are determined by linear relationships, hence are superposable. Both formulations include the possibility of a slow-time evolution of the primary wave amplitude, $a(\tau)$, and they have a solvability condition (§ 8) that brings these two effects together. Therefore, a formal synthesis is now made under the scale-ordering assumption that $\nu \sim \mu$ (§ 3), and this can be considered the distinguished limit for this multiscale asymptotic theory. If these parameters had disparate magnitudes, then either the drag or current effects would be dominant, with corresponding consequences for the wave-averaged evolution (§ 10); however, the assumption that they are comparable is the more general case that incorporates both possibilities. The combined slow-time coordinate is thus the fastest of the preceding τ coordinates.

In this section and the following one, the wave amplitude $a(\tau)$ is a complex function only of slow time (§ 4); thus, the drag-related formulae are slightly modified from § 5.

Consider a combined wave-modification velocity,

$$\mathbf{u}'_m = \mathbf{u}'_d + \mathbf{u}'_c, \tag{7.1}$$

which itself is decomposed into vortical and irrotational components, $\mathbf{u}'_m{}^{\zeta'}$ and $\mathbf{u}'_m{}^{\Phi'}$. The accompanying fields η'_m , ϕ'_m and Φ'_m are analogous combinations.

The goal is a wave-averaged theory, and thus the ‘fast’ dependence should be separated from the wave-averaged dependence. This is done by factoring $\mathcal{E}/2$ in the various \mathbf{u}'_m relations and focusing on their generally complex coefficient fields; e.g. for

$$\Phi'_m = Q \frac{\mathcal{E}}{2} + \text{c.c.}, \tag{7.2}$$

the complex factor $Q(\mathbf{x}, \tau)$ is entirely a wave-averaged quantity. For \mathbf{u}'_m , choose an analogous complex factor, $U_m = U^\zeta + U^\Phi$, so that

$$\mathbf{u}'_m = U_m \frac{\mathcal{E}}{2} + \text{c.c.} \tag{7.3}$$

For the vortical velocity components in (5.10) and (6.4), the combined factor is

$$U_m^\zeta = -\frac{i\mathcal{D}c_1(gF(z))^2|a|ak}{\sigma^3} \mathbf{k} + \frac{ga}{\sigma^2} \begin{pmatrix} ik^y(f + \bar{\zeta}^z)F(z) - k\bar{\zeta}^yG(z) \\ k\bar{\zeta}^xG(z) - ik^x(f + \bar{\zeta}^z)F(z) \\ i\hat{\mathbf{z}} \cdot \mathbf{k} \times \bar{\zeta}^h F(z) \end{pmatrix}. \tag{7.4}$$

For the irrotational velocity component, the elliptical PDE system for the combined factor Q from (5.13) and (6.6) is the following:

$$\left. \begin{aligned} [\nabla^2 + 2i\mathbf{k} \cdot \nabla^h - k^2]Q &= -c_1 \frac{(gF(z))^2|a|ak}{\sigma^2} (\mathbf{k} \cdot \nabla^h \mathcal{D} + ik^2 \mathcal{D}) \\ &\quad + \frac{ag}{\sigma} \left(F(z) \mathbf{k} \cdot \nabla^2 \bar{\mathbf{u}}^h - iG(z)k\nabla^2 \bar{w} \right), \\ \left[\partial_z - \frac{\sigma^2}{g} \right] Q &= 2i\sigma \partial_\tau a + a\sigma \left(-2\mathbf{k} \cdot \bar{\mathbf{u}}^h + \frac{g}{\sigma^2} \mathbf{k} \cdot \partial_z \bar{\mathbf{u}}^h + i\nabla^h \cdot \bar{\mathbf{u}}^h \right) \quad \text{at } z = 0, \\ \partial_z Q &= \frac{a\sigma \cosh[2kH]}{2k \sinh[kH]} \mathbf{k} \cdot \partial_z \bar{\mathbf{u}}^h \quad \text{at } z = -H. \end{aligned} \right\} \tag{7.5}$$

Notice that this system has combined the forcing term $\propto \partial_\tau a$ that is present in both (5.13) and (6.6). Finally, the associated irrotational velocity factor is

$$U_m^\Phi = \frac{1}{\sigma} (Q\mathbf{k} - i\nabla Q). \tag{7.6}$$

In both (7.5) and (7.6) the 3-D gradient operator ∇ has been expanded to include its action both on \mathcal{E} and on the wave-averaged factor Q .

Once $a(\tau)$ is known (§ 8), then (7.3)–(7.6) determine the \mathbf{u}'_m field sufficiently to evaluate the WEC terms in the current evolution equations (§§ 9 and 10).

8. Solvability condition for wave amplitude $a(\tau)$

In perturbation expansions, a common technique is to enforce an orthogonality condition between a leading-order quantity – here the primary wave \mathbf{u}'_w – and its small perturbation \mathbf{u}'_m to avoid ‘resonant’ forcing of the perturbation that upsets the asymptotic ordering in space and/or time. Combining (5.13) and (6.6) and formally defining the inhomogeneous

terms as R_\bullet factors in the PDE and boundary conditions gives the following PDE system for the wave-correction velocity potential:

$$\left. \begin{aligned} \nabla^2 \Phi'_m &= R_i, \\ \left(\partial_z - \frac{\sigma^2}{g} \right) \Phi'_m &= R_0 \quad \text{at } z = 0, \\ \partial_z \Phi'_m &= R_H \quad \text{at } z = -H. \end{aligned} \right\} \quad (8.1)$$

The homogeneous solution to this system is the ϕ' eigenmode, $\mathcal{G} = F(z)\mathcal{E}$, with the eigenvalue σ and wavenumber vector \mathbf{k} , i.e. the primary wave (§ 4). There is an analogous $\mathcal{G}^\dagger = F(z)\mathcal{E}^\dagger$ for any other \mathbf{k}^\dagger with the same wavenumber magnitude k (i.e. for a circle in \mathbf{k} -space of radius k).

The solvability condition that suppresses excitation of a homogeneous solution in Φ' is obtained as follows (WCI22, appendix E). Multiply the PDE in (8.1) by the complex conjugate of \mathcal{G}^\dagger (denoted by a $*$ superscript), integrate over the domain volume and then integrate by parts to insert the boundary conditions from (8.1). The result is the solvability condition:

$$\iiint \mathrm{d}\mathbf{x} \mathcal{G}^{\dagger*} R_i(\mathbf{x}) = \iint \mathrm{d}\mathbf{x}^h (\mathcal{G}^{\dagger*}(0)R_0(\mathbf{x}^h) - \mathcal{G}^{\dagger*}(-H)R_H(\mathbf{x}^h)). \quad (8.2)$$

Define a horizontal wavenumber projection operator,

$$\mathcal{P}^\dagger[\cdot] \equiv \frac{1}{\mathcal{A}} \iint \mathrm{d}\mathbf{x}^h \mathcal{E}^{\dagger*}[\cdot], \quad (8.3)$$

with the property that

$$\mathcal{P}^\dagger[\mathcal{E}'] = \delta_{\mathbf{k}', \mathbf{k}^\dagger}. \quad (8.4)$$

In the above, \mathcal{A} is the horizontal area of the domain.

The solvability condition (8.2) can be more compactly expressed as an integral only over z :

$$\int_{-H}^0 F(z) \mathcal{P}^\dagger[R_i] \mathrm{d}z = F(0) \mathcal{P}^\dagger[R_0] - F(-H) \mathcal{P}^\dagger[R_H]. \quad (8.5)$$

This condition must be satisfied for all \mathbf{k}^\dagger . Because R_0 contains a term with $(\partial_\tau a)\mathcal{E}$ and all the other R_\bullet terms depend only on $a\mathcal{E}$, this condition gives a first-order differential equation for $a(\tau)$. This assures that $Q(\mathbf{x})$ will not have a component proportional to $F(z)$ without any associated \mathbf{x}^h structure. Once the $a(\tau)$ are determined from (8.5), then (7.5) is a well-behaved PDE system for Φ'_m .

Notice that the solvability condition (8.5) for $a(\tau)$ depends only on \mathcal{D} , the primary wave properties, and the current $\bar{\mathbf{u}}$, i.e. it does not require knowing the solution for \mathbf{u}'_m . So $a(\tau)$ can be determined independently.

In advance of more complete solutions with both drag and current effects on the waves, the simple situation of drag without any current or Coriolis effect (i.e. $\bar{\mathbf{u}} = \mathbf{f} = 0$) and with a horizontally uniform canopy (i.e. $\nabla^h \mathcal{D} = 0$) can be evaluated using (8.5). In this case, the projection operations are trivial except for $\mathbf{k}^\dagger = \mathbf{k}$, and the solvability condition

becomes

$$\int_{-H}^0 F(z) \left(-i \frac{(g^2 |a| a k^3)}{\sigma^2} F^2 \mathcal{D} c_1 \right) dz = F(0) 2i\sigma \partial_\tau a. \quad (8.6)$$

After rearrangement, one obtains

$$\partial_\tau a = -J|a|a, \quad J = c_1 \frac{g^2 k^3}{2\sigma^3} \int_{-H}^0 F^3(z) \mathcal{D}(z) dz. \quad (8.7a,b)$$

With a slow-time initial condition of $a(\tau_0) = a_0 > 0$, the solution is a real, positive amplitude function:

$$a(\tau) = \frac{a_0}{1 + r(\tau - \tau_0)}, \quad r = a_0 J = c_1 \frac{g^2 k^3 a_0}{2\sigma^3} \int_{-H}^0 F^3(z) \mathcal{D}(z) dz. \quad (8.8a,b)$$

This shows a slow-time, inverse-linear decay in the primary wave amplitude at a rate r that depends on the primary wave properties, including its initial amplitude, the canopy density and the drag coefficient. (A similar linear-decay dependence for wave amplitude is found for canopy-drag effects in Dalrymple *et al.* (1984), but with somewhat different functional dependences than in r associated with how \mathcal{D} is specified.)

With non-zero currents ($\bar{\mathbf{u}} \neq 0$), the implications of (7.4)–(7.6) and (8.5) must mostly wait for combined drag and current solutions in a computational code built on top of the equations derived here. However, an important issue is the possibility of wave–current resonance, which is discussed extensively in WCI22 (chapters 7 and 8 and appendices J and K). If the current field has a wavenumber component \mathbf{p} , and if $|\mathbf{k}^\dagger| = |\mathbf{k} \pm \mathbf{p}| = k$, then there can be a slow-time energy exchange between the amplitudes $a(\tau; \mathbf{k})$ and $a(\tau; \mathbf{k}^\dagger)$. This occurs through the quadratic product of primary wave and current fields in the quasi-linear advection terms in (6.1). That is, new primary waves can arise through the wave–current interaction, and wave energy can be exchanged among these resonant components. In principle, their amplitudes can become as large as that of the original primary wave. Several idealized examples of wave–current resonance for particular $\bar{\mathbf{u}}(\mathbf{x})$ fields are analysed in WCI22, and in all such cases the slow-time amplitude behaviour is oscillatory between primary wave modes under the condition of steady currents. For most simple currents $\bar{\mathbf{u}}(\mathbf{x})$ and primary wave \mathbf{u}'_w fields, however, there is no resonance and no current-induced change in a . If there are multiple primary waves present (Appendix B), then the solvability conditions are still valid, although this would compound the resonant possibilities.

As the currents evolve with WEC and other influences (§ 10), it is presently unknown how often wave–current resonance might be an important phenomenon. (All of the papers cited in § 1 ignore this possibility except for WCI22.) A computational simulation strategy for dealing with it is also discussed in WCI22 (chapter 9), but it has not yet been implemented. For now, it remains an open issue for solving a quasi-linear wave model like the one here.

9. Wave-averaged drag force

The wave-averaged drag force is

$$\bar{D}^h = -\mathcal{D} \langle |\mathbf{u}^h| \mathbf{u}^h \rangle, \quad (9.1)$$

and the total velocity is defined in (2.4) and (2.5). In support of the WEC model in § 10, the goal is to obtain a wave-averaged formula for \bar{D} without having to do wave-resolved evaluations, nor even having to solve the elliptic problems for $\mathbf{u}_d^{\phi'}$ or $\mathbf{u}_c^{\phi'}$.

For the results in this and the following section, a is real and positive (§ 4), and the approximation (5.4) is made for $\mathbf{D}_w^{h'}$.

As a first step, consider an approximate form for the quadratic term in the drag law for a sum of velocities, $\mathbf{u} + \mathbf{v}$, when $|\mathbf{v}| \ll |\mathbf{u}|$:

$$\begin{aligned} |\mathbf{u} + \mathbf{v}|(\mathbf{u} + \mathbf{v}) &\approx |\mathbf{u} + \mathbf{v}|\mathbf{u} + |\mathbf{u}|\mathbf{v} \\ &\approx |\mathbf{u}|\mathbf{u} + (\mathbf{u} \cdot \mathbf{v}) \frac{\mathbf{u}}{|\mathbf{u}|} + |\mathbf{u}|\mathbf{v}, \end{aligned} \tag{9.2}$$

neglecting second-order terms in \mathbf{v} .

Assume, as in (3.2) and §§ 3–5, that the largest velocity is $\mathbf{u}_w^{h'}$. Inserting (2.4) and (2.5) into (9.1) – with the preceding approximation and $\nu \sim \mu$ ordering for $\mathbf{u}_m^{h'}$ and $\bar{\mathbf{u}}^h$ – gives

$$\bar{\mathbf{D}}^h = -\mathcal{D} \left\langle \left(|\mathbf{u}_w^{h'}| \mathbf{u}_w^{h'} + \mathbf{u}_w^{h'} \cdot (\bar{\mathbf{u}}^h + \mathbf{u}_m^{h'}) \frac{\mathbf{u}_w^{h'}}{|\mathbf{u}_w^{h'}|} + |\mathbf{u}_w^{h'}| (\bar{\mathbf{u}}^h + \mathbf{u}_m^{h'}) \right) \right\rangle. \tag{9.3}$$

The first right-hand-side term has already been evaluated to have a zero average in (5.3a,b). Because both $\mathbf{u}_w^{h'}$ and $\mathbf{u}_m^{h'}$ are $\propto \mathcal{E}$ in their fast-time dependence, here two other wave quantities that are either linear or cubic in their signed dependence on \mathcal{E} also have zero average because of the odd sign symmetry in different phases of Θ . Thus,

$$\left. \begin{aligned} \langle |\mathbf{u}_w^{h'}| \mathbf{u}_w^{h'} \rangle &= 0, \\ \left\langle (\mathbf{u}_w^{h'} \cdot \mathbf{u}_m^{h'}) \frac{\mathbf{u}_w^{h'}}{|\mathbf{u}_w^{h'}|} \right\rangle &= 0, \\ \langle |\mathbf{u}_w^{h'}| \mathbf{u}_m^{h'} \rangle &= 0. \end{aligned} \right\} \tag{9.4}$$

This is because they are proportional to the trigonometric functions of Θ in a way that has equal magnitude and opposite sign over the interval of $[0, 2\pi]$.

In contrast, $(\mathbf{u}_w^{h'} \cdot \bar{\mathbf{u}}^h) \mathbf{u}_w^{h'} / |\mathbf{u}_w^{h'}|$ and $|\mathbf{u}_w^{h'}|$ are single-signed and have non-zero averages. (These cancellations also extend to the relatively weak higher harmonics in $\mathbf{u}_w^{h'}$ (figure 1); see Appendix A.) Furthermore, for terms linear in the magnitude of $|\cos[\Theta]|$, the averaging operation gives

$$\frac{1}{2\pi} \int_0^{2\pi} |\cos[\Theta]| \, d\Theta = \frac{2}{\pi}. \tag{9.5}$$

Thus, with (4.3) the wave-averaged drag force (9.1) is

$$\begin{aligned} \bar{\mathbf{D}}^h &= -\mathcal{D} \left\langle (\mathbf{u}_w^{h'} \cdot \bar{\mathbf{u}}^h) \frac{\mathbf{u}_w^{h'}}{|\mathbf{u}_w^{h'}|} + (|\mathbf{u}_w^{h'}| + \bar{\mathbf{u}}^h) \bar{\mathbf{u}}^h \right\rangle \\ &= -\mathcal{D} \left(\frac{2ag}{\pi\sigma} F(z) \left(k\bar{\mathbf{u}}^h + \frac{\mathbf{k} \cdot \bar{\mathbf{u}}^h}{k} \mathbf{k} \right) + |\bar{\mathbf{u}}^h| \bar{\mathbf{u}}^h \right). \end{aligned} \tag{9.6}$$

In an *ad hoc* way (i.e. beyond leading order in μ), the last term is added to have a well-behaved limit in the absence of a primary wave when the drag is entirely due to the current.

Thus, the canopy exerts a horizontal drag force on the currents that is enhanced by the primary wave. The spatial structure of this drag is the product of the canopy distribution $\mathcal{D}(\mathbf{x})$, the wave eigenmode profile $F(z)$ and the horizontal current itself $\mathbf{u}^h(\mathbf{x})$. It is retarding in the direction of the current in all three of the terms in (9.6); i.e. the drag work,

$\bar{D}^h \cdot \bar{u}^h$, is negatively proportional to each of $(\bar{u}^h)^2$, $(\mathbf{k} \cdot \bar{u}^h)^2$ and $|\bar{u}^h|(\bar{u}^h)^2$. The first two terms in (9.6) have a scaling magnitude of $(\nu\mu)C^2/\ell$, while the last term is $(\nu\mu^2/\epsilon)C^2/\ell$, where ℓ is a length scale for both the waves and the currents. As desired, (9.6) is expressed entirely in terms of wave-averaged quantities. Notice that there is no dependence on \mathbf{u}'_m , so the wave perturbations need not be solved for, apart from whatever $a(\tau)$ dependence is caused by the solvability condition (§ 8).

It would be straightforward to extend the representation of \bar{D}^h to include a bottom drag stress and again have a wave enhancement over the dependence on only \bar{u}^h . The form is likely to be similar to (9.6), which is functionally different from presently used parametrization formulae for wave-enhanced bottom stress in a turbulent bottom boundary layer (e.g. Feddersen *et al.* 2000). This issue is worth further consideration.

10. Wave effects on the currents (WEC)

The approximate wave-averaged current evolution equations for the general model (2.1) are

$$\left. \begin{aligned} [\partial_\tau + \bar{\mathbf{u}} \cdot \nabla] \bar{\mathbf{u}} + f \hat{\mathbf{z}} \times \mathbf{u}^h + \nabla \bar{\phi} &= -\nabla \bar{B} + \langle \mathbf{u}' \times \boldsymbol{\zeta}' \rangle + \bar{D}^h, \\ \nabla \cdot \bar{\mathbf{u}} &= 0, \end{aligned} \right\} \quad (10.1)$$

with vertical boundary conditions of $\bar{w} = 0$ at $z = 0$ and $-H$ (i.e. a rigid flat top and bottom; see WCI22, chapter 6, but with \bar{D}^h added). The time coordinate is again $\tau = \mu t$, based on the advective time scale of current evolution, although one can also view this as valid for $\tau = \nu t$ when $\nu \sim \mu$ (§ 7).

The added forces due to the presence of waves are on the right-hand side of (10.1). The wave-averaged drag \bar{D}^h is in (9.6). The wave-averaged Bernoulli head is

$$\bar{B} = \frac{1}{2} \langle \mathbf{u}'^2 \rangle / 2 \quad (10.2)$$

(WCI22, chapter 6), and it does not have to be evaluated explicitly because the ‘pressure Poisson’ equation from the incompressibility condition in (10.1) gives a combined expression for the potential force, $\bar{\phi} + \bar{B}$, as part of the time integration of (10.1) for $\bar{\mathbf{u}}$. This assumes that the lateral boundary conditions are consistent with their combination, as certainly is true for horizontal periodicity. The approximate rigid-lid condition on \bar{w} is associated with a higher-order kinematic surface boundary condition that is a diagnostic relation for wave-averaged sea level, $\bar{\eta}$, which is much smaller than η'_w (WCI22, appendix B).

The total wave-averaged vortical force is a combination of drag and current effects on the waves:

$$\mathbf{VF} = \langle \mathbf{u}' \times \boldsymbol{\zeta}' \rangle \approx \langle \mathbf{u}'_w \times (\boldsymbol{\zeta}'_d + \boldsymbol{\zeta}'_c) \rangle, \quad (10.3)$$

where the wave vorticities are in (5.9) and (6.3). In a way that is familiar in the papers referenced in the introduction, the wave-averaged wave vortex force associated with the currents is

$$\left. \begin{aligned} \mathbf{SVF} &= \langle \mathbf{u}'_w \times \boldsymbol{\zeta}'_c \rangle = \mathbf{u}^{St} \times (f \hat{\mathbf{z}} + \bar{\boldsymbol{\zeta}}) - \nabla \bar{C}, \\ \bar{C} &= \hat{\mathbf{z}} \cdot \int_{-H}^0 \mathbf{u}^{St}(z') \, dz \times (f \hat{\mathbf{z}} + \bar{\boldsymbol{\zeta}}) \end{aligned} \right\} \quad (10.4)$$

(WCI22, section 6.1 and appendix G). The right-hand-side terms are evaluated using (4.3) and (6.4). The first term is the familiar Stokes–Coriolis and Stokes–vorticity forces.

The second term $\nabla\bar{C}$ does not have to be evaluated explicitly because it, too, combines into the total potential force for the $\bar{\mathbf{u}}$ evolution.

The WEC vortical force associated with the wave-drag vortex force is evaluated using (4.3) and (5.9):

$$CVF = \langle \mathbf{u}'_w \times \boldsymbol{\zeta}'_d \rangle = -c_1 \frac{g^3 a^3 k^2}{2\sigma^4} G(z) \partial_z [DF^2] \mathbf{k}. \tag{10.5}$$

Because \mathcal{D} , F , G and $\partial_z F^2$ are all positive from (1.2) and (4.4), this force can be either retarding or accelerating in the direction of primary wave propagation, depending on the depth profile of the canopy structure, $\partial_z \mathcal{D} \propto \partial_z \alpha(\mathbf{x})$ in (1.2). A scaling estimate is $CVF \sim \nu \epsilon^2 C^2 / \ell$, compared to $SVF \sim \mu \epsilon^2 C^2 / \ell$; i.e. they are of similar magnitude if $\nu \sim \mu$. As remarked following (9.6), the wave-enhanced drag terms are $\bar{\mathbf{D}}^h \sim \nu \mu C^2 / \ell$; therefore, all three wave forces are comparable if $\epsilon^2 \sim \mu \sim \nu$.

Thus, the final form of the current momentum balance is

$$\begin{aligned} & [\partial_\tau + \bar{\mathbf{u}} \cdot \nabla] \bar{\mathbf{u}} + f \hat{\mathbf{z}} \times \mathbf{u}^h + \nabla \bar{\Pi} \\ &= VF + \bar{\mathbf{D}}^h \\ &= -c_1 \frac{g^3 a^3 k^2}{2\sigma^4} G(z) \partial_z [DF^2] \mathbf{k} + \mathbf{u}^{St} \times (f \hat{\mathbf{z}} + \bar{\boldsymbol{\zeta}}) + \bar{\mathbf{D}}^h, \end{aligned} \tag{10.6}$$

where $\bar{\Pi} = \bar{\phi} + \bar{B} + \bar{C}$ is the combined force potential function, $VF = CVF + SVF$ is the combined vortex force and $\bar{\mathbf{D}}^h$ is the combined drag force on the currents (9.6). The right-hand-side terms are, respectively, the vortex force due to canopy drag on the waves CVF , the Stokes vortex forces SVF and the wave-enhanced current drag. This equation is the principal result in this paper.

Notice that (10.6) is entirely specified in terms of \mathcal{D} , $\bar{\mathbf{u}}$ and the primary wave properties. Among the latter is the slow-time evolution of the primary wave amplitude $a(\tau)$ as determined from the solvability condition (8.5) that includes canopy wave damping and possible wave–current resonances. This system is readily implemented in an LES or circulation model, with the most challenging new element being the solvability condition (§ 8) that has heretofore been neglected in wave–current interaction studies.

The vortex force is a product of the quasi-linear CEW theory (§ 6) at leading order in μ (WCI22, section 5.1); the higher-order CEW and WEC terms presented in WCI22 (sections 5.2 and 6.6) are not included here, although they could be, as their forms are known. But this would require explicit solutions of the elliptic problems in §§ 5–7 that are not required for (10.6). Recall that another approximation is indicated in (5.3a,b), where the weak higher harmonics in \mathbf{D}'_w are neglected; this simplifies the theory, and Appendix A shows that this approximation is fairly accurate for representing the further consequences of \mathbf{D}'_w in § 5, hence in $\bar{\mathbf{D}}^h$ as well.

11. Weaker waves

For the preceding \mathbf{u}'_c , \mathbf{u}^{St} and SVF results, it does not matter how small $\epsilon \ll 1$ is. However, this is not true for the canopy effects. They were derived for the scaling relation (3.2), i.e. $\mu \ll \epsilon$, which yields the analytic relations in § 5 that carry through to the current effects in § 10. If the wave orbital velocities are weaker, with $\mu \sim \epsilon$, then the total drag force (1.1)

at leading order, viz.

$$\mathbf{D}^h = -\mathcal{D}|\mathbf{u}_w^{h'} + \bar{\mathbf{u}}^h|(\mathbf{u}_w^{h'} + \bar{\mathbf{u}}^h) \equiv \mathbf{D}_w^{h'} + \bar{\mathbf{D}}^h, \tag{11.1}$$

does not have explicit analytic expressions for its fluctuation and wave-averaged drag-force components for the primary wave eigenmode (4.3) and a general current field $\bar{\mathbf{u}}(\mathbf{x})$. Thus, at every location \mathbf{x} and slow time τ , these forces must be computed by taking a numerical average for $\bar{\mathbf{D}}^h$ over the primary wave period as an auxiliary calculation for the current evolution equation.

The resulting $\mathbf{D}^{h'} = \mathbf{D}^h - \bar{\mathbf{D}}^h$ can then be used to evaluate (5.6) as

$$\begin{aligned} \partial_t \boldsymbol{\zeta}'_d &= \text{curl}[\mathbf{D}^{h'}] \\ &= \begin{pmatrix} -\partial_z D^{y'} \\ +\partial_z D^{x'} \\ -\hat{\mathbf{z}} \cdot \nabla^h \times \mathbf{D}^{h'} \end{pmatrix}, \end{aligned} \tag{11.2}$$

with discretized spatial derivatives in the right-hand-side quantities. Similarly, a spatially discretized $\nabla^h \cdot \mathbf{D}^{h'}$ can be used to evaluate the solvability condition (8.5), also by numerical averaging with the projection operator \mathcal{P} . There is no need to solve for \mathbf{u}'_m with weaker waves.

For the current evolution equation analogous to (10.6), the current drag $\bar{\mathbf{D}}^h$ can be inserted directly, and the canopy vortex force \mathbf{CVF} is calculated as follows:

$$\begin{aligned} \mathbf{CVF} &= \langle \mathbf{u}'_w \times \boldsymbol{\zeta}'_d \rangle = \left\langle \mathbf{u}'_w \times \int^t \text{curl}[\mathbf{D}^{h'}] dt \right\rangle = - \left\langle \left(\int^t \mathbf{u}'_w dt \right) \times \text{curl}[\mathbf{D}^{h'}] \right\rangle \\ &= \left\langle \begin{pmatrix} \left(\int^t v'_w dt \right) (\hat{\mathbf{z}} \cdot \nabla^h \times \mathbf{D}^{h'}) + \left(\int^t w'_w dt \right) \partial_z D^{x'} \\ \left(\int^t w'_w dt \right) \partial_z D^{y'} - \left(\int^t u'_w dt \right) (\hat{\mathbf{z}} \cdot \nabla^h \times \mathbf{D}^{h'}) \\ - \left(\int^t u'_w dt \right) \cdot \partial_z \mathbf{D}^{h'} \end{pmatrix} \right\rangle. \end{aligned} \tag{11.3}$$

With the primary wave (4.3) and real a , this becomes

$$\begin{aligned} \mathbf{CVF} &= \frac{ga}{\sigma^2} (F \sin[\Theta] (\hat{\mathbf{z}} \cdot \nabla^h \times \mathbf{D}^{h'}) (\hat{\mathbf{z}} \times \mathbf{k}) \\ &\quad + kG \cos[\Theta] \partial_z \mathbf{D}^{h'} \\ &\quad + F \sin[\Theta] (\mathbf{k} \cdot \partial_z \mathbf{D}^{h'}) \hat{\mathbf{z}}). \end{aligned} \tag{11.4}$$

Again a fast-time numerical average over a primary wave period is needed, as indicated by the angle brackets, where Θ and $\mathbf{D}^{h'}$ are the fast-time variables. In (11.4) the first two lines are horizontal components of \mathbf{CVF} , and the final line is a vertical component. Thus, in this weaker wave case, the wave-averaged effects on the currents (WEC) are still well determined, but at the cost of added numerical averages of the wave-induced drag and canopy vortex force.

If the primary wave is even weaker, with $\mathbf{u}'_w \ll \bar{\mathbf{u}}$ and $\epsilon \ll \mu$, then $\bar{\mathbf{D}}^h \approx -\mathcal{D}|\bar{\mathbf{u}}^h|\bar{\mathbf{u}}^h$, and \mathbf{CVF} is much smaller than, for example, the advective tendencies in the current evolution equation. That is, wave effects on currents (WEC) are negligible, and the most important interaction is current effects on waves (CEW) (e.g. a Doppler shift in σ) and a fluctuating canopy drag force on the waves.

12. Multiple primary waves

Now reverting to the principal scaling assumption (3.2), a wave-averaged theory is still achievable for multiple components in the primary wave field, i.e. for N components,

$$\eta'_w = \sum_{n=1}^N a_n \frac{\mathcal{E}_n}{2} + \text{c.c.}, \tag{12.1}$$

with amplitudes a_n and exponential phase factors $\mathcal{E}_n = \exp[i\Theta_n]$ with $\Theta_n = \mathbf{k}_n \cdot \mathbf{x} - \sigma_n t$. Again, these waves are eigenmodes with σ_n as in (4.2) for \mathbf{k}_n , accompanying fields as in (4.3), vertical structure functions F_n and G_n as in (4.4), and a Stokes drift $\mathbf{u}^{St}(z)$ as in (4.6) that is a sum over the N separate contributions.

However, such a theory has added complexity in dealing with the drag effects on both the waves and currents. In particular, for use in the drag force on the waves (§ 5), the formula (5.1) is evaluated with the primary wave horizontal velocity,

$$\mathbf{u}'_w = \sum_n \frac{g a_n}{\sigma_n} F_n(z) \mathbf{k}_n \frac{\mathcal{E}_n}{2} + \text{c.c.}, \tag{12.2}$$

which then propagates through the equivalents of §§ 5–10. The source of the complexity is that analytic expressions are no longer available for D'^h_w , \bar{D}^h and CVF . Besides necessary numerical averages over the primary wave periodic interval for these quantities and in the solvability condition (8.5) (as in § 11), the simplifications in (9.4) are no longer strictly valid, so that $\mathbf{u}'_m = \mathbf{u}'_d + \mathbf{u}'_c$ contributions to \bar{D}^h would also need to be solved for computationally with the elliptic PDE system (7.5). This adds up to a considerable computational burden beyond what is involved with the analytic expressions in §§ 9 and 10.

As discussed in Appendix B, some quantitative exploration of D^h with multiple waves gives support for thinking that there may be useful approximate shortcuts to the full programme of numerical evaluations in the preceding paragraph. In particular, solving for \mathbf{u}'_m to evaluate \bar{D}^h in (9.3) might not be necessary if their associated quantities in (9.4) are not too large compared to the \mathbf{u}'_w contributions. Of course, this would need to be tested in particular problems of interest.

On the other hand, the CEW corrections in §§ 6–8 are easily generalized with multiple primary waves by simple superposition (as discussed in WCI22), as are their associated WEC terms in § 10. This superposition can even be extended to broadband wave spectra, e.g. for \mathbf{u}^{St} in SVF .

For weaker waves (§ 11), multiple primary waves add only a modest burden to the indicated numerical evaluations, with \mathbf{u}'_w in the formulae (11.1) and (11.3) understood to be the summation over the multiple primary wave components. The wave perturbations \mathbf{u}'_d and \mathbf{u}'_c continue to be negligible in this case.

13. Summary and prospects

This paper presents a theory for the wave-averaged effects of surface gravity waves and currents in the presence of a vegetative canopy, suitable for a modified LES or fine-scale circulation model. The principal new effects – compared to previous wave-averaged models with only a mean-current drag force and the Stokes vortex forces associated with the currents (e.g. Yan *et al.* 2021) – are a wave-amplitude damping (8.5), a wave-enhanced current drag (9.6) and a wave-drag vortex force (10.5). The model includes WEC for

current evolution (§ 10) plus wave damping and possible resonances for the primary wave amplitude(s), $a(\tau; \mathbf{k})$. For a single primary wave field, the set of resonance amplitudes is limited to a circle in \mathbf{k} space, which should not be burdensome to calculate (versus a broadband wave field). Interestingly, apart from its role in setting up the wave amplitude equations, the irrotational part of the wave perturbations, $\mathbf{u}_d^{\phi'} + \mathbf{u}_c^{\phi'}$, need not be solved for this model. This feature is in common with the original Stokes vortex-force models (e.g. Craik & Leibovich 1976) at leading order in μ , where knowledge of $\mathbf{u}_c^{\zeta'}$ suffices for determining the WEC expression in *SVF*.

For weaker waves (§ 11) and *a fortiori* for multiple primary waves (§ 12 and Appendix B), a wave-averaged theory for the current evolution is feasible in principle but probably laborious in practice because of the extra computational evaluations to obtain the final WEC equations. From this perspective, the monochromatic primary wave theory presented here, with its analytic wave-averaged expressions, is useful for clarifying wave-enhanced drag and vortex force effects and for supporting a rather straightforward computational implementation for current evolution. However, the generalization to multiple primary waves or a realistic broadband wave spectrum is much more daunting and might eventually require some form of wave-resolved modelling. If this were done, some ‘long-wave’ approximation to the free-surface dynamics could avoid the complexities of resolving spectrum broadening and wave breaking (versus its parametrization). Also, it would likely be advantageous to do a split-explicit time stepping (as commonly done for circulation models by separating the barotropic and baroclinic vertical modes (Shchepetkin & McWilliams 2005)) with fast time stepping of the surface gravity wave component and subsequent numerical averaging on the slow time step of the currents. This type of hybrid wave–LES model would be a pioneering task. It could bypass the need for a formal wave-averaged theory, although it is likely that such theories as here would be useful for interpreting the computational results. Nevertheless, the community experience with Langmuir turbulence and depth-induced nearshore wave breaking indicates that important lessons can be learned even with only a monochromatic wave field representing the spectrum-peak component.

It seems worthwhile to proceed to a computation implementation of (10.6), e.g. for a kelp canopy, while also including the necessary non-conservative current effects (including some form of wave-enhanced bottom drag for the currents; see § 9), and, for now, without a formal extension of the theory to include buoyancy effects. The latter has been done previously for wave-averaged CEW and WEC theories (McWilliams *et al.* 2004; WCI22, chapter 10), and it probably would contain no surprises with a canopy. The waves are not influenced by buoyancy because the surface-interface gravitational force is so much stronger than the force associated with interior stratification, and the main additional WEC influence is Stokes-drift buoyancy advection. An important implementation ingredient is sufficient experimental information about $\mathcal{D}(x)$, which often is not easily obtained. Nevertheless, many nearshore and estuarine sites have marine vegetation that is likely to be influential on the local currents and material transports.

Another necessary extension for nearshore realism in the wave-averaged theory and computational model built upon it is inclusion of both bottom drag and variable bottom depth. The former is briefly alluded to at the end of § 9 and is not inherently difficult. The latter is naturally a part of the resolved current dynamics, and small-scale roughness elements are a part of a Monin–Obukhov specification of the bottom drag coefficient. However, further thought is needed for how to efficiently include finite-height, finite-width (versus horizontally slowly varying, where a ray theory would be appropriate) depth variations in the wave component of the flow. Perhaps this, too, would be an argument

for developing a long-wave free-surface LES model as mentioned in the second paragraph of this section. At the present time, oceanic LES examples for flow over such topography are rare, even in the absence of surface wave effects – but see Calhoun & Street (2001), Calhoun, Street & Kosoff (2001) and Skillingstad & Wijesekera (2004).

Beyond the derivation of wave-averaged equations, a measured programme of idealized computational experiments would be the right next step. Among the references in § 1 are various experimental results about wave damping in the presence of currents, and a computational implementation of this theory could take advantage of their various determinations of $\mathcal{D} = \alpha(x)C_D$. In an oceanically targeted LES model with finite depth and bottom-rooted vegetation such as kelp, simple geometries of different plant profiles and different horizontal patches could be explored with various surface stress and buoyancy flux and primary waves to assess how the combined wave–current–canopy interactions manifest. Our recent papers partly set this framework (Yan *et al.* 2021; Yan, McWilliams & Chamecki 2022), but as yet without the wave-enhanced drag effects. Beyond these idealized studies, some field conditions could be targeted for experimental testing. For various motivations, it would be timely to focus on kelp examples, both the heterogeneous natural patches found in shallow water outside the surf zone and the present and future suspended-platform aqua-farm configurations.

It is a delicate balancing act between do-able theories that enable informative complex-flow simulations and their underlying approximations that are only marginally defensible compared to nature’s true complexities. Formulating, solving and experimentally testing feasible, relevant problems involving waves, currents, topography and canopies still have a long way to go.

Acknowledgments. I appreciate discussions with M. Chameki and D. Dauhajre.

Funding. This work is supported by the ARPA-E MARINER Program no. DE-AR0000920 and the NSF grant no. OCE-2124174.

Declaration of interests. The author reports no conflict of interest.

Author ORCID.

 James C. McWilliams <https://orcid.org/0000-0002-1237-5008>.

Appendix A. Full drag force D_w^h for a monochromatic primary wave

In § 5 an approximation is made to (5.2) in (5.3a,b) based on their similarity of shape in figure 1. Now the modifications based on not making this approximation are shown.

The full drag formula is rewritten as

$$D_w^h = -\mathcal{D} \left(\frac{(ga)^2 F^2 k}{\sigma^2} \right) kH(\Theta) \tag{A1}$$

with

$$H(\Theta) = |\cos[\Theta]| \cos[\Theta]. \tag{A2}$$

As with its approximate form, $\langle H \rangle = 0$; i.e. the wave drag is entirely oscillatory. The phase function H is shown in figure 2 along with its approximate form.

Again, a vortical–irrotational decomposition is made for the wave perturbation u'_d . The vortical part again satisfies (5.6), but now the drag forcing term, $\text{curl}[D_w^{h'}]$, is the same as

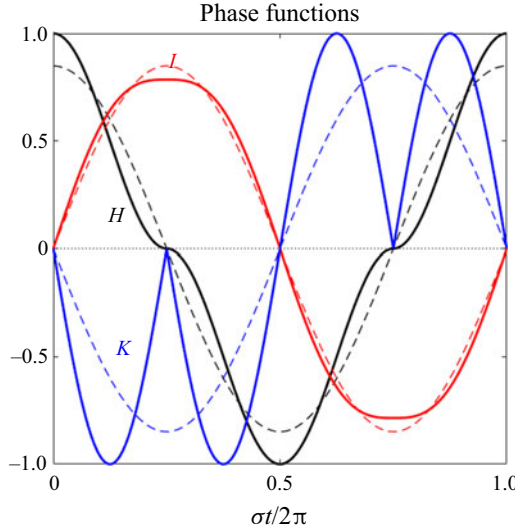


Figure 2. Plots of the full-drag phase functions, $H(\Theta')$ (black solid line), $I(\Theta')$ (red solid line) and $K(\Theta')$ (blue solid line), over one wave period in $\Theta' = \sigma t/2\pi$. Their counterparts to (A2), (A3) and (A6) with the approximation in (5.3a,b) are plotted with thin dashed lines with their corresponding colours. (The black lines replicate the lines in figure 1.)

(5.7) except that $c_1 \cos[\Theta]$ is replaced by $H[\Theta]$. To integrate the vorticity tendency, define

$$I(\Theta) \equiv \int^\Theta H(\Theta) d\Theta = \frac{1}{2} \text{sign}[\cos(\Theta)] \left(\Theta + \frac{1}{2} \sin[2\Theta] \right). \quad (\text{A3})$$

The second-line formula is taken from the Wolfram integrator (<http://integrals.wolfram.com/index.jsp>), and the function is plotted in figure 2. This function $I(\Theta)$ replaces $c_1 \sin[\Theta]$ in (5.9) and (5.10):

$$\left. \begin{aligned} \zeta'_d &= \frac{g^2 a^2 k}{\sigma^3} \begin{pmatrix} -k^y \partial_z [DF^2] \\ +k^x \partial_z [DF^2] \\ -F^2 (\hat{\mathbf{z}} \cdot \mathbf{k} \times \nabla^h \mathcal{D}) \end{pmatrix} I(\Theta), \\ \mathbf{u}_d^{h\zeta'} &= \mathcal{D} \left(\frac{(gaF)^2 F^2 k}{\sigma^3} \right) I[\Theta] \mathbf{k}, \quad w'_d = 0. \end{aligned} \right\} \quad (\text{A4})$$

The irrotational potential PDE forcing (5.15) is

$$\begin{aligned} \text{div}[\mathbf{D}_w^h] &= - \left(\frac{(gaF)^2 k}{\sigma^2} \right) (\mathbf{k} \cdot \nabla^h) [DH[\Theta]] \\ &= - \left(\frac{(gaF)^2 k}{\sigma^2} \right) \{k^2 \mathcal{D}K[\Theta] + (\mathbf{k} \cdot \nabla^h \mathcal{D})H[\Theta]\}, \end{aligned} \quad (\text{A5})$$

where

$$K(\Theta) \equiv \partial_\Theta H = -2|\cos[\Theta]| \sin[\Theta]; \quad (\text{A6})$$

see figure 2. This function $K(\Theta)$ replaces $-c_1 \sin[\Theta]$. Now the PDE system (5.13) for Φ'_d no longer has such a simple separable solution as in (5.16) even when $\nabla^h \mathcal{D} = 0$ because the differential equation and surface boundary condition have different dependences on Θ .

Section 6 is unaltered by the more general form of D_w^h here. The synthesis of drag and current effects on the waves in § 7 changes in the way indicated above for § 5 and will not be repeated. However, because the Θ dependences are no longer simply proportional to \mathcal{E} , the solution for Φ'_m and u'_m can no longer be simply factored as in (7.2) and (7.3).

The general discussion of the solvability condition (§ 8) still applies. For the particular example of drag in a uniform canopy (i.e. $\bar{u} = f = \nabla^h \mathcal{D} = 0$, as in (8.6)–(8.8a,b)), the relevant form of the projection operator is

$$\mathcal{P}[\bullet] = 2\langle \sin[\Theta] \bullet \rangle, \tag{A7}$$

which yields a value of 1 when applied to the coefficient $\sin[\Theta]$ of $\partial_z a$ in R_0 . When applied to the coefficient of $\text{div}[D'_h]$ in (A5) in R_i , the result using (A6) is

$$\mathcal{P}[-K(\Theta)] = 2 \frac{1}{2\pi} \int_0^{2\pi} 2|\cos[\Theta']| \sin^2[\Theta'] d\Theta' = \frac{8}{3\pi}, \tag{A8}$$

i.e. the same value as c_1 in § 5 after (5.3a,b); $R_{-H} = 0$ in this case. Thus, the solvability condition gives the same result for canopy-induced amplitude decay as in (8.8a,b) whether or not the approximation (5.3a,b) is made.

The mean current drag, \bar{D}^h in § 9, is unchanged with the more complete form of D_w^h in (9.6), because it does not depend on u'_d except insofar as the trivial relations in (9.4) are still valid. This is so because the more general functional shapes in figure 2 have the same symmetries in Θ as do their approximate counterparts.

In § 10 the drag-induced vortex force $\langle u'_w \times \zeta'_d \rangle$ involves the drag-induced wave vorticity,

$$\zeta'_d = \frac{g^2 a^2 k}{\sigma^3} \begin{pmatrix} -k^y \partial_z [\mathcal{D}F^2] \\ +k^x \partial_z [\mathcal{D}F^2] \\ -F^2 (\hat{z} \cdot \mathbf{k} \times \nabla^h \mathcal{D}) \end{pmatrix} I[\Theta], \tag{A9}$$

instead of (5.9). Thus, the wave-averaged vortex force with u'_w from (4.3) and ζ'_d from (A9) involves the following quantities:

$$\left. \begin{aligned} 2\langle \cos[\Theta] I[\Theta] \rangle &= 0, \\ 2\langle \sin[\Theta] I[\Theta] \rangle &= \frac{8}{3\pi} = c_1. \end{aligned} \right\} \tag{A10}$$

Thus, the outcome is the same as (10.5), and the final form of the current momentum balance is unaltered from (10.6).

In summary, the more complete quadratic functional form of the wave-drag force D_w^h has only slightly different outcomes for the drag-induced wave correction u'_d and no differences for the wave-averaged drag and vortex forces, justifying the simplifying assumption made in § 5.

Appendix B. Drag force D^h with multiple primary waves

With multiple primary waves (§ 12), the simple analytic expression for D_w^h in (5.4) is not available, nor is $\bar{D}_w^h = 0$. This is illustrated in this appendix and its implications considered.

The wave-averaged theory is based on a fast-time average over a period T to distinguish wave and current components. For a clean distinction, all wave components must have a commensurate frequency such that $2\pi/\sigma_n$ is an integer fraction of T . Once the

wave-averaged balances are derived, then T need not be explicit, but with the implicit assumption that it is short compared to the current evolution on the slow time τ . This is true asymptotically for μ and ν sufficiently small.

For illustration, consider the case of $N = 2$. Assume that $\sigma_1 \leq \sigma_2$. Choose T such that both σ_1 and σ_2 are commensurate frequencies; i.e. the period of the first component is T/m and the period of the second component is T/q (m and q being integers). The commensuration requirement is that $m \geq 1$ is the least integer such that $q = (\sigma_2/\sigma_1)m \geq m$ is also an integer. Possible examples are $(m, q) = (1, 2), (1, 8), (2, 3)$ and $(7, 15)$; notice that these integers can have no common divisors. Define a temporal phase function $\tilde{\Theta} = -\sigma_1 t$. Without loss of generality, choose the direction of \mathbf{k}_1 to be $\hat{\mathbf{x}}$. Then normalize $\mathbf{u}_w^{h'}$ by the amplitude and phase of its first wave component. Thus, the fast phase behaviour of the primary wave velocity is represented by the non-dimensional horizontal vector,

$$\mathbf{U}'(\tilde{\Theta}) = \cos[m\tilde{\Theta}]\hat{\mathbf{x}} + \alpha_2 \cos[q\tilde{\Theta} + \gamma_2]\hat{\mathbf{k}}_2. \tag{B1}$$

Here $\alpha_2 \geq 0$ is the relative amplitude of the second component; $\hat{\mathbf{k}}_2 = (\cos[\lambda_2], \sin[\lambda_2])$ is its relative wavenumber vector orientation with relative angle λ_2 ; and γ_2 is its relative phase. The function \mathbf{U}' is periodic in $\tilde{\Theta}$ over $[0, 2\pi]$, and it has zero average, $\langle \mathbf{U}' \rangle = 0$, for all parameter choices. With respect to the primary wave component drag force, the question is how the phase average of

$$\mathbf{\Delta}(\tilde{\Theta}) = |\mathbf{U}'| \mathbf{U}' \tag{B2}$$

depends on $(m, q, \alpha_2, \lambda_2, \gamma_2)$. To assess its average value in relation to its fluctuations, $\mathbf{\Delta}$ can be normalized by its r.m.s. value, $\text{r.m.s.}[\mathbf{\Delta}] = \langle \mathbf{\Delta}^2 \rangle^{1/2}$.

For $\alpha_2 = 0$ (a single component), $\langle \mathbf{\Delta} \rangle = 0$, as in § 4. Similarly, $R_\Delta \equiv \langle \mathbf{\Delta} \rangle / \text{r.m.s.}[\mathbf{\Delta}] \rightarrow 0$ as $\alpha_2 \rightarrow \infty$; thus, the mean primary wave drag is largest for $\alpha_2 \sim 1$. Extensive computational exploration of the parameters listed after (B2) shows that R_Δ is always small compared to unity, usually much smaller, and often zero. An apparent global maximum value is $R_\Delta = 0.225$ for the parameter set $(1, 2, 0.65, 0, 0)$. I know of no analytic formula for \bar{D}_w^h with multiple wave components, but it is readily evaluated in a numerical time average over the period T .

Thus, \bar{D}_w^h is not always zero with multiple waves, although its relative value seems usually to be small. This leads to a decomposition,

$$\mathbf{D}_w^h = \bar{\mathbf{D}}_w^h + \mathbf{D}_w^{h'}, \tag{B3}$$

with $\mathbf{D}_w^{h'}$ the larger right-hand-side quantity. The programme followed in § 5 is based on the generalization of $\mathbf{D}_w^{h'}$ instead of (5.4). Again, a reduced harmonic component approximation can be made:

$$\mathbf{D}_w^{h'} \approx \sum_n \mathbf{A}_n^h \cos[\Theta_n], \tag{B4}$$

where the fast-time dependence has been separated with $\Theta_n = \mathbf{k}_n \cdot \mathbf{x}^h - \sigma_n t + \gamma_n$ and spatially varying, real vector coefficients \mathbf{A}_n^h that are determined by a least-squares fit to the full $\mathbf{D}_w^{h'}(t)$ in a numerical time average over the commensurate frequency period T . These fits show that the \mathbf{A}_n^h components in general depend on all the primary component amplitudes $\{a_m\}$ and have orientations different from \mathbf{k}_n . As with the monochromatic approximation in (5.3a,b), experimentation with two components shows that the relative error in (B4) is small, i.e. at most a few tens of a per cent.

With the approximation (B4), the curl and divergence of $\mathbf{D}_w^{h'}$ are themselves sums over spatial differential functions of A_n^h times $\cos[\Theta_n]$ and $\sin[\Theta_n]$. Thus, ζ'_d and $\mathbf{u}_d^{h\zeta'}$ can be explicitly specified as in (5.9) and (5.10), and so can the right-hand-side forcing terms in the PDE system (5.13) for Φ'_d , hence $\mathbf{u}_d^{\Phi'}$ as in (5.11), both as in § 5. The former allows an explicit analytic wave averaging of the wave-induced vortex force *CVF* as a sum over the components of products of the primary wave $\mathbf{u}_w^{h'}$ harmonic coefficients and the functionals of A_n^h (cf. (10.5) in § 10). The latter allows a numerical evaluation of the solvability conditions for the primary component amplitudes $a_n(\tau)$ as averages over the commensurate period T and the vertical integral in (8.5) in § 8.

Just as $\bar{\mathbf{D}}_w^h \neq 0$ with multiple wave components, no longer are the several particular components of $\bar{\mathbf{D}}^h$ listed in (9.4) zero. As a first level of generalization, the previously evaluated $\bar{\mathbf{D}}_w^h$ could be retained in $\bar{\mathbf{D}}^h$ for the wave-averaged drag. Although it is formally larger than the other terms in (9.6) by a factor of $O(\nu^{-1})$ or $O(\mu^{-1})$, its relative value is small compared to its scaling estimate (as discussed in the fifth paragraph of this appendix). For the other two non-zero terms in (9.4), one could optimistically argue that they have small relative values compared to the non-zero components in the first line in (9.6), hence also small compared to $\bar{\mathbf{D}}_w^h$, and thereby neglect them. Otherwise, a numerical solution of (5.13) is required for $\mathbf{u}_d^{\Phi'}$, and numerical time averages over T are required for evaluating $\langle (\mathbf{u}_w^{h'} \cdot \mathbf{u}_d^{h'}) (\mathbf{u}_w^{h'} / |\mathbf{u}_w^{h'}|) \rangle$ and $\langle |\mathbf{u}_w^{h'}| |\mathbf{u}_d^{h'}| \rangle$; these are substantial auxiliary calculations. (Their counterparts with $\mathbf{u}_c^{h'}$ remain zero.) The other non-zero components listed in the first line in (9.6) could be evaluated with numerical time averages over T using (12.2).

REFERENCES

- BRUNET, Y. 2020 Turbulent flow in plant canopies: historical perspective and overview. *Boundary-Layer Meteorol.* **177**, 315–374.
- CALHOUN, R.J. & STREET, R.L. 2001 Turbulent flow over a wavy surface: neutral case. *J. Geophys. Res.* **106**, 9277–9293.
- CALHOUN, R.J., STREET, R.L. & KOSOFF, J.R. 2001 Turbulent flow over a wavy surface: stratified case. *J. Geophys. Res.* **106**, 9295–9310.
- CRAIK, A.D.D. & LEIBOVICH, S. 1976 A rational model for Langmuir circulations. *J. Fluid Mech.* **73**, 401–426.
- DALRYMPLE, R.A., KIRBY, J.T. & HWANG, P.A. 1984 Wave diffraction due to areas of energy dissipation. *ASCE J. Waterway Port Coastal Ocean Engng* **110**, 67–79.
- FEDDERSEN, F., GUZA, R., ELGAR, S. & HERBERS, T. 2000 Velocity moments in alongshore bottom stress parameterizations. *J. Geophys. Res.* **105**, 8673–8686.
- FINNIGAN, J. 2000 Turbulence in plant canopies. *Annu. Rev. Fluid Mech.* **32**, 519–571.
- HANLEY, K.E., BELCHER, S.E. & SULLIVAN, P.P. 2010 A global climatology of wind-wave interaction. *J. Phys. Oceanogr.* **40**, 1263–1282.
- HENDERSON, S.M. 2019 Motion of buoyant, flexible aquatic vegetation under waves: simple theoretical models and parameterization of wave dissipation. *Coast. Engng* **152**, 103497.
- HOLM, D.D. 1996 The ideal Craik–Leibovich equations. *Physica D* **98**, 415–441.
- HU, Z., SUZUKI, T., ZITMAN, T., UITTEWAAL, W. & STIVE, M. 2014 Laboratory study on wave dissipation by vegetation in combined current–wave flow. *Coast. Engng* **77**, 131–142.
- HYPOLITE, D., ROMERO, L., MCWILLIAMS, J.C. & DAUHAIRE, D. 2022 Langmuir circulations transfer kinetic energy from submesoscales to dissipative scales. *J. Phys. Oceanogr.* **53**, 253–268.
- LEI, J. & NEFF, H. 2019 Wave damping by flexible vegetation: connecting individual blade dynamics to the meadow scale. *Coast. Engng* **147**, 138–148.
- LOSADA, I.J., MAZA, M. & LARA, J.L. 2016 A new formulation for vegetation-induced damping under combined waves and currents. *Coast. Engng* **107**, 1–13.

Surface waves and currents in aquatic vegetation

- LUHAR, M. & NEPF, H.M. 2011 Flow-induced reconfiguration of buoyant and flexible aquatic vegetation. *Limnol. Oceanogr.* **56**, 2003–2017.
- MCWILLIAMS, J.C. 2022 Quasi-linear theory for surface wave–current interactions. In *Mathematics of Planet Earth*. Springer.
- MCWILLIAMS, J.C., RESTREPO, J.M. & LANE, E.M. 2004 An asymptotic theory for the interaction of waves and currents in coastal waters. *J. Fluid Mech.* **511**, 135–178.
- MCWILLIAMS, J.C., SULLIVAN, P.P. & MOENG, C.-H. 1997 Langmuir turbulence in the ocean. *J. Fluid Mech.* **334**, 1–30.
- MONISMITH, S., ALNAJJAR, M., DALY, M., VALLE-LEVINSON, A., JUAREZ, B., FAGUNDES, M., BELL, T. & WOODSON, C.B. 2022a Kelp forest drag coefficients derived from tidal flow data. *Estuar. Coast.* **45**, 2492–2503.
- MONISMITH, S.G., ALNAJJAR, M.W., WOODSON, C.B., BOCH, C.A., HERNANDEZ, A., VAZQUEZ-VERA, L., BELL, T.W. & MICHELL, F. 2022b Influence of kelp forest biomass on nearshore currents. *J. Geophys. Res.* **127**, e2021JC018333.
- NEPF, H. 2012 Flow and transport in regions with aquatic vegetation. *Annu. Rev. Fluid Mech.* **44**, 123–142.
- ROSMAN, J., DENNY, M.W., ZELLER, R.B., MONISMITH, S.G. & KOSEFF, J.R. 2013 Interaction of waves and currents with kelp forests (macrocystis pyrifera): insights from a dynamically scaled laboratory model. *Limnol. Oceanogr.* **53**, 798–802.
- SHAW, R.H. & SCHUMANN, U. 1992 Large-eddy simulation of turbulent flow above and within a forest. *Boundary-Layer Meteorol.* **61**, 47–64.
- SHCHEPETKIN, A.F. & MCWILLIAMS, J.C. 2005 The regional oceanic modeling system (ROMS): a split-explicit, free-surface, topography-following-coordinate oceanic model. *Ocean Model.* **9**, 347–404.
- SKYLLINGSTAD, E.D. & WIJESEKERA, H.W. 2004 Large-Eddy simulation of flow over two-dimensional obstacles: high drag states and mixing. *J. Phys. Oceanogr.* **34**, 94–112.
- SULLIVAN, P.P. & MCWILLIAMS, J.C. 2010 Dynamics of winds and currents coupled to surface waves. *Annu. Rev. Fluid Mech.* **42**, 19–42.
- SUZUKI, N. & FOX-KEMPER, B. 2016 Understanding Stokes forces in the wave-averaged equations. *J. Geophys. Res.* **121**, 3579–3596.
- UTTER, B. & DENNY, M. 1996 Wave-induced forces on the giant kelp *Macrocystis pyrifera* (Agardh): field test of a computational model. *J. Expl Biol.* **199**, 2645–2654.
- YAN, C., MCWILLIAMS, J.C. & CHAMECKI, M. 2022 Overlapping boundary layers in the inner shelf zone. *J. Phys. Oceanogr.* **52**, 627–646.
- YAN, C., MCWILLIAMS, J.C. & CHAMECKI, M. 2021 Generation of attached Langmuir circulations in the upper ocean mixed layer over a macroalgal farm. *J. Fluid Mech.* **915**, A76.
- YU, X., ROSMAN, J.H. & HENCH, J.L. 2022 Boundary layer dynamics and bottom friction in combined wave–current flows over large roughness elements. *J. Fluid Mech.* **931**, A11.
- ZHAO, C., TANG, J., SHEN, Y. & WANG, Y. 2021 Study on wave attenuation in following and opposing currents due to rigid vegetation. *Ocean Engng* **236**, 109574.