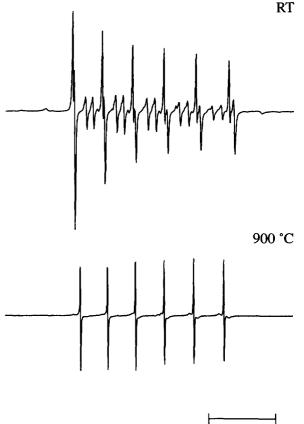
RESIDUAL MANGANESE (II) SPECIATION IN MONTMORILLONITE

McBride (1995) has called attention to the possibility that the residual Mn(II) species in untreated SWy-1 montmorillonite, investigated spectroscopically by Gehring *et al* (1993), could be Mn²⁺ in the structure of calcite, a known impurity associated with the clay mineral. In support of his proposal, McBride (1995) notes that the EPR spectrum for Mn(II) originates primarily from the >2 μ m fraction of the SWy-1 sample and that the EPR lineshape is similar to that observed for Mn²⁺ substituted in the calcite structure, both at room temperature and after heating at 700°C.

Gehring *et al* (1993) concluded on the basis of variable-temperature EPR spectroscopy that residual Mn(II) in untreated SWy-1 montmorillonite could be either an isomorphically-substituted species or an adsorbed species. They rejected the former possibility because of the known ready exchangeability of Mn(II) in SWy-1 montmorillonite and instead proposed that residual Mn(II) is a partially-solvated, adsorbed species in the montmorillonite interlayers. This hypothesis was consistent with the observed persistence of the Mn(II) EPR spectrum after heating at 600°C and with the drastic changes in the spectrum occurring along with the dehydroxylation of montmorillonite at 600°-700°C and with its decomposition at 800°C.

The EPR spectrum of Mn(II) after heating at 700°C is indeed similar in SWy-1 montmorillonite (Gehring et al 1993, Figure 6) and calcite (McBride 1995, Figure 1), giving credence to the proposal of McBride (1995). However, the temperature effect on the signal intensity for the SWy-1 sample was to reduce it by about 90% (Gehring et al 1993), whereas little or no reduction of signal intensity occurred for the calcite sample (Mc-Bride 1995). The proposal that the drastic change in Mn(II) spectral lineshape observed by Gehring et al (1993) after heating at 800°C is simply the result of calcite conversion to CaO (McBride 1995) is also not consistent with the evident persistence of the EPR signal in calcite heated at 900°C (Figure 1). Gehring et al. (1993, Figure 5), by contrast, observed disappearance of the signal already after heating at 800°C, coincident with the decomposition of the clay mineral at this temperature.

In our view, the assignment of the Mn(II) EPR signal in SWy-1 montmorillonite remains problematic, requiring further experimental investigation. We thank Professor McBride for contributing to the discussion of this point with his most helpful insights.



20 mT

Figure 1. Room temperature EPR spectra of $CaCO_3$ powder, untreated and after heating at 900°C for 5 hours.

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