

Millimetre-Wave Site Characteristics at the Australia Telescope Compact Array

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Abstract

We present a millimetre-wave site characterisation for the Australia Telescope Compact Array (ATCA) based on nearly 9 yr of data from a seeing monitor operating at this facility. The seeing monitor, which measures the phase fluctuations in the signal from a geosynchronous satellite over a 230-m baseline caused by water vapour fluctuations along their sight lines, provides an almost gapless record since 2005, with high time resolution. We determine the root mean square (rms) of the path length variations as a function of time of day and season. Under the assumption of the ‘frozen screen’ hypothesis, we also determine the Kolmogorov exponent, α , for the turbulence and the phase screen speed. From these, we determine the millimetre-wave seeing at $\lambda = 3.3$ mm. Based on the magnitude of the rms path length variations, we estimate the expected fraction of the available observing time when interferometry could be successfully conducted using the ATCA, as a function of observing frequency and antenna baseline, for the time of day and the season. We also estimate the corresponding observing time fractions when using the water vapour radiometers installed on the ATCA in order to correct for the phase fluctuations occurring during the measurement of an astronomical source.

Keywords: atmospheric effects – instrumentation: interferometers – site testing – techniques: interferometric

1 INTRODUCTION

The Australia Telescope Compact Array (ATCA) is a radio interferometer located near the town of Narrabri in northwestern New South Wales, Australia at S30°18′46″ E149°33′00″. For over 10 yr it was the only millimetre interferometer located in the southern hemisphere (Stevens & Wark 2011). It was originally planned as a centimetre wavelength synthesis telescope with upgrades to the 12-, 7- and 3-mm observing bands to be implemented at a later stage. This, along with some accounts of how the site selection came to agree on Narrabri, is described in detail in Frater, Brooks, & Whiteoak (1992). Following a proposal to fit out the array with millimetre wave receivers, a site testing programme was initiated to establish its feasibility. Longer term observations however were limited to opacity measurements of the atmosphere at 30 GHz while single snapshot observations were proposed at a wavelength of 3 cm with the array antennae. This is described in an ATNF internal memo (Hall et al. 1992). From these measurements the fraction of nights when the transmission is suitable for observation at a given wavelength can be estimated, but they do not allow the phase stability for operating an interferometer as a function

of baseline to be determined. In this paper we present a comprehensive analysis of the millimetre site characteristics at the ATCA based on 8.5 yr of data and extending the analysis of Middelberg et al. (2006), which used the initial 1 yr of data gathered with a seeing monitor that was installed in 2004.

2 DATA DESCRIPTION

The seeing monitor is a fixed baseline interferometer with a 230-m baseline and two receivers. They observe a transmitter beacon on a geosynchronous satellite. From 2004 until the satellite’s decommissioning in September 2008, the signal used was the 30.48-GHz beacon on the Optus B3 satellite. The seeing monitor receiver hardware and software then was adapted to the new beacon frequency of 21.198 GHz on the Optus C1 satellite. The beacon signal is down converted and mixed to 50 kHz and sampled at 10 Hz with a phase locked loop integration time of 100 ms. Individual phase measurements, however, are stored every 5 s only. The phases were corrected to their zenith values by multiplying by the cosine of the zenith angles (51° and 57° respectively). From this data, a zenith path difference Δd is then formed for two

successive 5-s time intervals from $\lambda \Delta \Phi / 2\pi$, where λ is the wavelength of the beacon signal. A second-order polynomial is fitted and used to eliminate satellite drift and thermal effects, as per the description in Middelberg et al. (2006). The standard deviation as well as the root mean square (rms) of this Δd is computed for a 30-min window. Because 30 min is much larger than the crossing time of a parcel of the atmosphere over the 230-m baseline, using the standard deviation provides a robust measure of the magnitude of the phase fluctuations occurring over a small portion of the day. For all the parameters derived (see below) a 5- σ iterated sigma-clipping about the median value was also performed to remove deviant data. This resulted in the removal of no more than 2% of the relevant values. The design of the seeing monitor, along with a site analysis spanning approximately 1 yr of data has been described in further detail in Middelberg et al. (2006) so we do not include it here.

We have examined 8.5 yr worth of seeing monitor data spanning from April 2005 to October 2013. The site quality statistics we have derived include:

- the rms path length differences as a function of time of day and season,
- the lag phase structure functions (LPSFs) and Kolmogorov exponents, together with the phase screen speed,
- the visibility efficiency as a function of baseline and frequency,
- the millimetre-wave seeing at 3.3 mm based on the path length fluctuations and Kolmogorov exponents.

The Kolmogorov exponent determines whether the atmosphere above the ATCA can be treated as two-dimensional (2D) ‘frozen screen’ (see below). The other parameters provide useful metrics for the fraction of time that conditions are sufficiently good to observe in a particular millimetre waveband. We conclude with an analysis outlining the improvement in observing time that could be achieved when using the ATCA water vapour radiometers (WVRs) described in (Indermuehle, Burton, & Crofts 2013), as a function of baseline, frequency, time of day and season.

3 THEORY OF TURBULENCE

Andrey Kolmogorov in 1941 developed the now widely accepted theory that small-scale turbulent motions are statistically isotropic and thus do not have any preferential spatial direction (Kolmogorov 1941). The large-scale turbulent cells on the other hand are not isotropic because their behaviour is determined by the shape of the boundary. Kolmogorov’s hypothesis entertains the idea that a turbulent flow of velocity V contains large vortices which over time decompose into smaller vortices until they have reached a scale where the remaining kinetic energy is turned into internal energy through viscous dissipation. Kolmogorov introduced the structure function in order to describe the fluctuations in

the velocity field as a function of separation r :

$$D(r) = \int_{-\infty}^{\infty} dx (f(x) - f(x+r))^2, \quad (1)$$

which is equivalent to the mean square difference between two points of separation r :

$$D(r) = \langle (f(x) - f(x+r))^2 \rangle. \quad (2)$$

We model the layer containing the eddies of water vapour as a phase screen moving across the telescope aperture. Taylor in his seminal paper on ‘The Spectrum of Turbulence’ (Taylor 1938) found that for a given moving air mass with embedded eddies of varying properties such as temperature, humidity, or water vapour, as long as the velocity of the air mass is much greater than the turbulent velocity, the sequence of changes measured by a probe at a given point simply measures an unchanging pattern of eddies going by. This means the eddies, as related to the probe, are not changing and seen as statically embedded (or ‘frozen’). This is today referred to as the ‘frozen screen’ hypothesis (or Taylor hypothesis) and in our case applies to eddies of poorly mixed water vapour in the atmosphere. Assuming that the phase screen velocity, v_s , does not change significantly during the time period, τ , over which a measurement is made, we can state that τ becomes equivalent to the baseline b through the relationship

$$v_s = \frac{b}{\tau}, \quad (3)$$

where the lag time τ is the time it takes for the phase screen to pass over the baseline b . We can then formulate the temporal phase structure function, or LPSF:

$$D_\phi(\tau) = \left\langle \left(\phi(t) - \phi\left(t + \frac{b}{v_s}\right) \right)^2 \right\rangle, \quad (4)$$

where ϕ is the phase, as modified by the water vapour. By evaluating the LPSF using the seeing monitor data we can examine the scale of the turbulence, as well as the velocity of the phase screen v_s through the following relationship:

$$D_\phi(\tau) = D_\phi(b) \mid b = v_s \tau, \quad (5)$$

where D is the phase structure function of the turbulence. The square root of this corresponds to the rms phase variation in the following way:

$$\Phi_{\text{rms}}(\tau) \equiv \sqrt{D_\phi(b)}. \quad (6)$$

This can be formulated in a general way to determine a baseline and wavelength-dependent term K as shown by Coulman (1990) and Carilli & Holdaway (1999):

$$\Phi_{\text{rms}}(b) = K \left(\frac{\lambda_0}{\lambda} \right) \left(\frac{b}{b_0} \right)^\alpha, \quad (7)$$

where λ is the observing wavelength, λ_0 that of the seeing monitor, b is the antennae baseline length, and b_0 the

separation of the seeing monitors. α is the Kolmogorov exponent and K is measured in radians. With $\lambda = \lambda_0$ and $b = b_0$, $\Phi_{0,\text{rms}} = K$ is then the rms phase fluctuation measured with the seeing monitor itself. According to Kolmogorov's theory of turbulence, the phase noise *versus* baseline relationship should follow Equation 7 with $\alpha = \frac{1}{3}$ for turbulence with baselines longer than the width of the turbulent layer (i.e. $b > h$). This is referred to as 2D turbulence, because the phase screen assumes the statistical properties of a 2D thin screen.

The other case is where the baseline is shorter than the width of the turbulent layer ($b < h$) and has a Kolmogorov exponent of $\alpha = \frac{5}{6}$. The phase screen then exhibits three-dimensional (3D) behaviour and is called the thick screen.

There is also an outer scale, L_0 , beyond which the rms phase variation should no longer increase with baseline length. This corresponds to $\alpha = 0$. We cannot determine the outer scale however because we are not measuring multiple baselines at a single point in time.

The phase variations averaged over a given integration time cause coherence loss in the measured visibility. For a given visibility $V = V_0 e^{i\phi}$ the effect on the measured amplitude due to phase noise in a given averaging time is

$$\langle V \rangle = V_0 \times \langle e^{i\phi} \rangle = V_0 \times e^{-\phi_{\text{rms}}^2/2} \quad (8)$$

This assumes Gaussian random phase fluctuations of ϕ_{rms} over the averaging time. This is discussed in more detail in Carilli & Holdaway (1999).

4 THE RMS PATH LENGTH DIFFERENCE

The seeing monitor provides only two output parameters: (zenith) phase and time. Taking these we can build the phase differences between adjoining measurements. The phase difference rms and its standard deviation can then be analysed. Because the (zenith) path differences fluctuate around a mean near zero, there is minimal difference between these measures. We use rms as the preferred metric. The phase differences can easily be converted to path differences using the frequency of the observed signal, $d = \lambda \Delta \Phi / 2\pi$.

Figures 1 and 2 give examples of a week of path variations in January and July. The diurnal range of path fluctuation varies by about an order of magnitude in both the summer and winter data, but the amplitudes are higher in summer by a factor of about two. In the summer, they vary from about 250 μm at night to 2 500 μm during the day while in the winter, 100–950 μm of diurnal variation is more common.

As can be seen from Figure 3, which shows the histogram and cumulative distribution of the rms path length differences through the year, the variations between the summer and winter months are clear. In May and July, for 80% of the time the rms path difference is less than 400 μm , whereas in January, for the same fraction of time, the rms path length

is twice as high, i.e. about 800 μm . The histogram features a marked peak for both the May and July curves, indicating a prolonged period of time (lasting from May to August) where rms path difference conditions of just 150 μm or less are experienced for one-third of the time. Examination of Table 1, where these values are tabulated, confirms that all the winter months feature rms values close to their minimum value.

By further breaking this data down into time of day (i.e. 3-h time bands) statistics, as shown in Figure 4, it can be seen that the worst conditions in winter, occurring between 12 h and 15 h local time (LT) (with about 750- μm rms path difference), are similar to the average conditions in summer. However, while these are considerably better than conditions in the worst periods experienced during summer (12 h–21 h LT), they also are worse than the best periods in summer, which last from 21 h–09 h LT, i.e. from late night to mid-morning. The slope of the steeply rising portion of the cumulative distribution in Figure 4 is steeper for the winter data and the difference between daytime and nighttime conditions is half of that in summer, indicating that the good conditions in winter last for a longer period of time; for roughly 15 h per day (from 21 h–09 h LT), there is an 80% probability of the rms path noise being less than 250 μm . In comparison, during the best times in summer, lasting for 12 h per day, there is an 80% probability of the path noise being less than 500 μm (twice the typical winter-time rms path difference).

The distribution of the rms path differences, split into quartiles for each month and 3-h time band, are listed in Table 1, together with their maximum and minimum values within each month/time band. The maximum and minimum values encountered in these quartiles over the entire year are listed in Table 2. As an illustrative example of the use of these tables, in April at 12 noon, for 50% of the time the path differences will be less than 728 μm . The lowest value for the third quartile path difference occurs in May between midnight and 3 am, and is 224 μm .

A direct comparison of our results with the analysis of Middelberg et al. (2006) is not readily possible as these authors convert the rms path differences they determine to a 1-km baseline, making use of the Kolmogorov exponent they find for each path difference (i.e. scaling using Equation 7). We instead provide the statistics in Tables 1 and 2 for the 230-m seeing monitor baselines as these can be used for direct comparison with the values provided to the observer by the ATCA control system, to assist in deciding upon the appropriate project to pursue at a particular time. We discuss how this information can be used in Section 5.5. We have, however, compared our results to Middelberg et al. (2006) by converting the rms path differences we measure to a 1-km baseline using the local value for the Kolmogorov exponent. The distributions we see in June and November, as per those shown by Middelberg et al. (2006) in their Figure 5, are broadly similar. For instance, in June for roughly 80% of the time the 1-km path differences are better than 600 μm

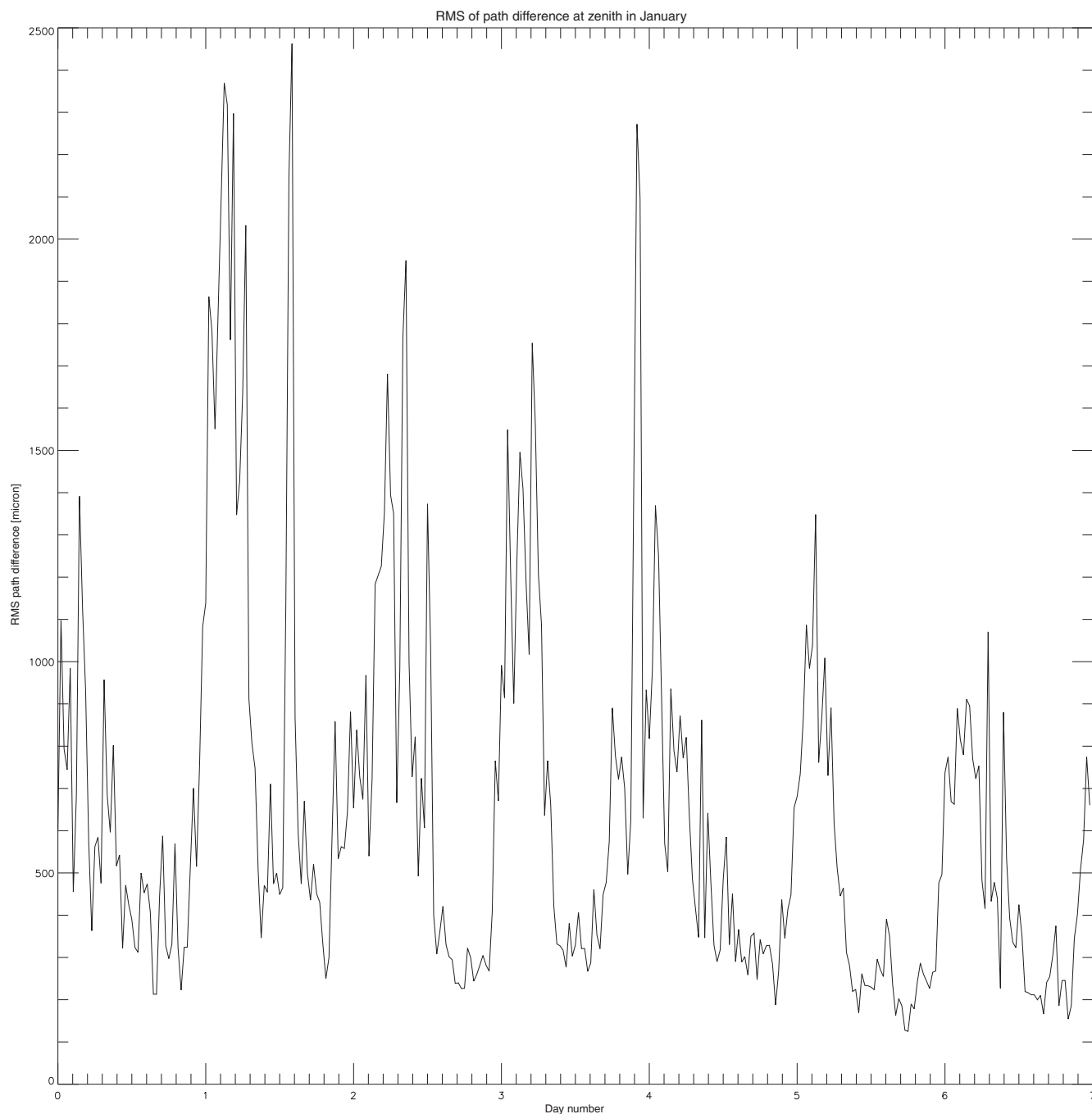


Figure 1. The rms path difference for a week's worth of data in January. The range of path fluctuation varies by about an order of magnitude, from $\sim 250 \mu\text{m}$ at night to $\sim 2500 \mu\text{m}$ during the day time.

for 15 h per day in both our studies. The corresponding difference is $1\,000 \mu\text{m}$ in November. However, for the lowest quality periods, we find there is significantly more time with large rms differences over the 9- versus 1-yr range of the two datasets. For instance, in November between 12 h and 15 h, we find that for roughly half the time the path differences are greater than $1\,500 \mu\text{m}$, in comparison to $<10\%$ of the time by Middelberg et al. (2006). Year to

year variations can be significant and might account for this difference.

5 DERIVED PARAMETERS

We can derive further information from the rms path length differences by analysing the behaviour in the time domain, applying the theory outlined in Section 3. This yields the

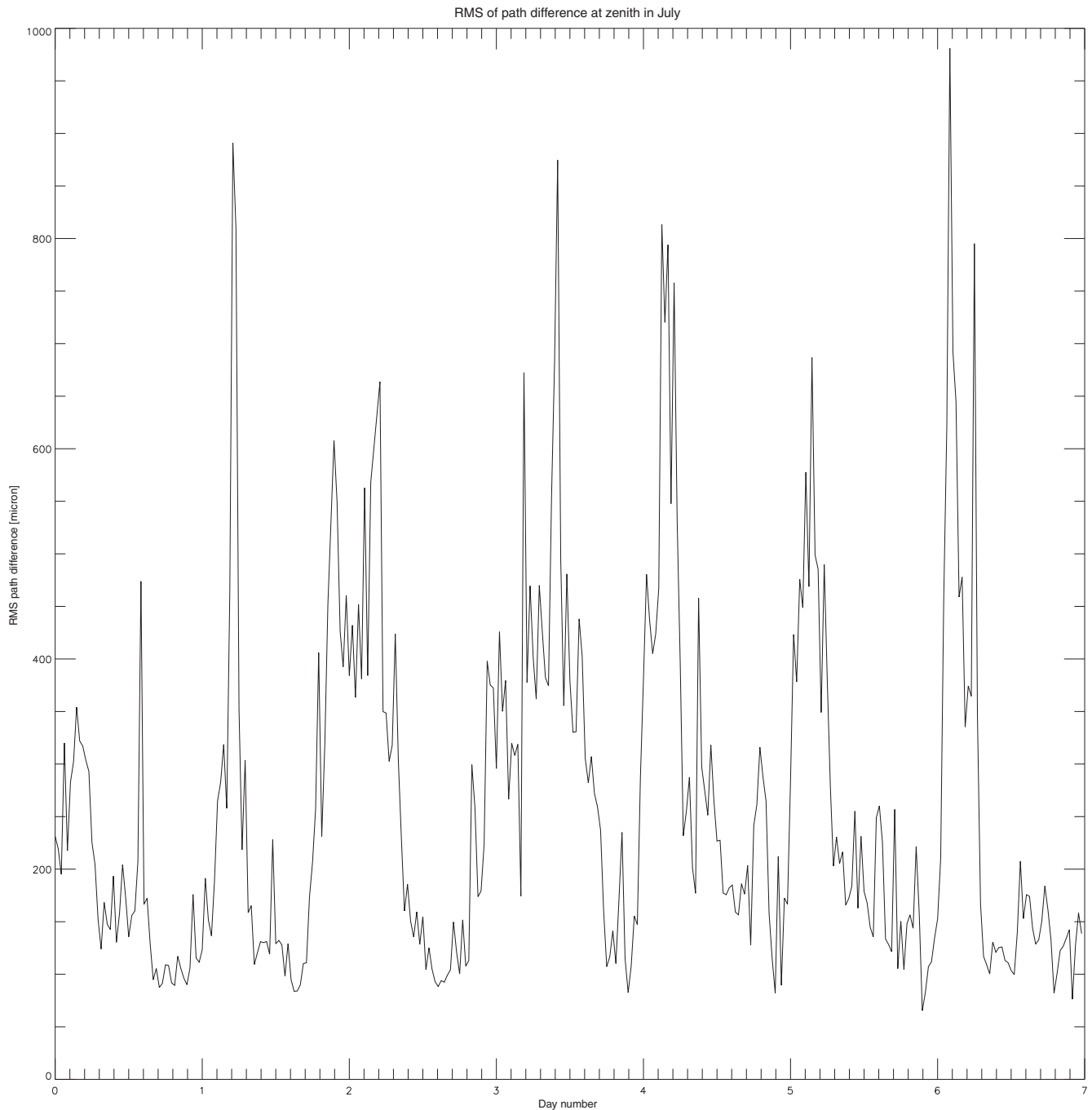


Figure 2. The rms path difference for a week's worth of data in July. The fluctuations are similar in magnitude to the summer data in Figure 1, but their overall level is much lower, varying between 100 and 950 μm .

Kolmogorov exponent (α), corner time (t_c), saturation path length (p_{sat}), and phase screen speed (v_s), as we discuss below.

5.1 The Kolmogorov exponent

Figure 5 shows an example of a lag phase plot used to extract the LPSF parameters. This illustrates the behaviour of the fluctuating phase screen as it moves across the baseline. As

long as the structure of the air mass is not self similar, the rms of the path length variation in the measurement span increases to a maximum after which subsequent data result in a reduction in the rms because a self-similar scale size has been reached. This behaviour continues until the scale size of the turbulent layers in the atmosphere have become too extended and the data have become statistically independent, thus marking the point where the outer scale length for the turbulence has been reached.

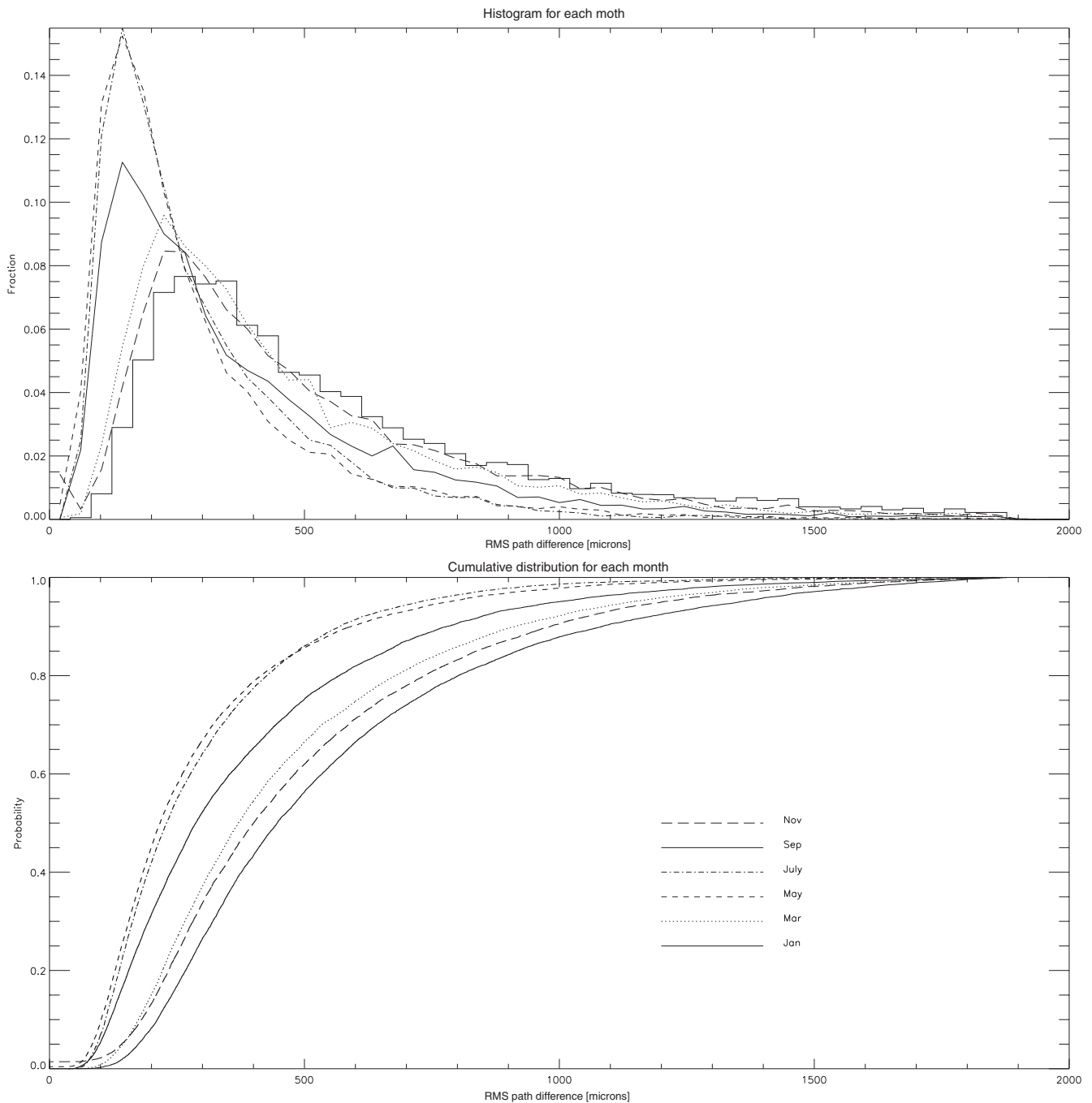


Figure 3. Histogram and cumulative distribution of the rms path differences in μm . The differences between summer and winter months are clear, as discussed in Section 4.

The parameters that can be derived from the LPSF include the corner time t_c , which is the time that elapses until the phase noise has peaked for the first time as well as the saturation path length noise $p_{\text{rms,sat}}$. We determine the corner time as the time when two or more successive path length noise points are smaller than the preceding points. The saturation path noise p_{rms} at t_c can then be compared to the standard deviation of the path difference σ_{pd} and because the data can be approximated by a series of sinusoids fluctuating about a

zero point, it should satisfy $p_{\text{rms}} \approx \sigma_{\text{pd}} \times \sqrt{2}$.¹ We then use the corner time t_c and the baseline length b to determine the velocity of the phase screen v_s according to the relationship shown in Equation 3. The gradient fit in log space to the phase noise before the corner time has been reached determines the Kolmogorov exponent α .

¹ The rms of any sinusoid is equal to its amplitude/ $\sqrt{2}$.

Table 1. Quartile values for the rms zenith path differences over the 230-m seeing monitor baseline determined for the 8.5-yr dataset from April 2005 to October 2013. The data has been split by time of day into 3-h intervals (the starting time for each is listed; i.e. 00 ≡ 0–3 h, etc.) and month. Also shown are the maximum and minimum values and the hours when they occur, for each quartile and month combination. Times are in Australian Eastern Standard Time (AEST).

Time/Month	25%, 50%, and 75% quartiles in μm								Min	Hour	Max	Hour
	00	03	06	09	12	15	18	21				
Jan	229	212	265	465	621	543	319	251	212	3	621	12
	303	287	356	661	883	722	430	347	287	3	883	12
	423	412	495	945	1243	1026	604	478	412	3	1243	12
Feb	221	205	250	467	653	525	303	254	205	3	653	12
	296	281	357	660	931	725	414	338	281	3	931	12
	417	404	487	935	1259	985	552	477	404	3	1259	12
Mar	195	177	200	407	582	427	253	215	177	3	582	12
	262	248	275	559	800	589	346	291	248	3	800	12
	359	350	382	782	1053	822	467	422	350	3	1053	12
Apr	150	144	153	308	523	311	193	161	144	3	523	12
	206	196	212	434	728	431	253	221	196	3	728	12
	296	293	303	639	958	614	356	311	293	3	958	12
May	116	111	112	197	370	219	145	125	111	3	370	12
	159	154	160	305	528	306	194	171	154	3	528	12
	224	234	233	476	752	438	269	236	224	0	752	12
Jun	125	118	122	194	328	211	145	139	118	3	328	12
	180	177	173	286	461	292	200	195	173	6	461	12
	269	270	258	405	628	405	282	277	258	6	628	12
Jul	121	114	118	198	370	226	149	133	114	3	370	12
	171	162	162	294	497	317	208	178	162	6	497	12
	250	237	251	433	671	441	288	251	237	3	671	12
Aug	121	111	117	256	419	254	150	128	111	3	419	12
	167	161	166	371	574	356	210	176	161	3	574	12
	259	242	252	550	789	512	294	256	242	3	789	12
Sep	130	126	148	371	482	288	167	143	126	3	482	12
	191	184	228	532	679	415	234	202	184	3	679	12
	289	282	348	757	945	571	335	284	282	3	945	12
Oct	144	139	190	416	516	333	183	159	139	3	516	12
	212	199	284	592	725	471	251	226	199	3	725	12
	307	305	423	875	993	662	374	341	305	3	993	12
Nov	204	191	268	478	583	398	239	214	191	3	583	12
	273	263	363	678	808	557	336	298	263	3	808	12
	410	362	507	947	1103	796	490	442	362	3	1103	12
Dec	196	192	261	504	589	456	253	209	192	3	589	12
	280	270	367	704	873	643	350	295	270	3	873	12
	418	396	536	1008	1187	890	505	440	396	3	1187	12

In Figure 6 we show the histogram and cumulative distribution of the Kolmogorov exponents for the complete date range from 2005–2013. The distributions for the time of day (in 3-h bands) for July (i.e. winter) and January (i.e. summer) are shown in Figure 7. Neither the diurnal nor the seasonal variations are large. As can be seen from Table 3, the median value is 0.41, and the median ranges between 0.40–0.43 over the months of the year. The time of day dependence has a preference towards a larger α during daytime hours from 09:00 to 18:00 local time. The slope of the January cumulative distribution in Figure 7 is steeper than in July with their peaks being approximately the same, $\alpha = 0.65$. At the lower end however, the winter time Kolmogorov exponent extends to smaller values by about $\Delta\alpha \approx 0.1$. Variations in the Kolmogorov exponent

are clearly dominated by the time of day rather than the season.

From these data it is evident that for the great majority of time the site experiences 2D turbulence since α is close to the value of 1/3 (see Section 3). This means that the vertical extent of the turbulent layer is thinner than the length of the baseline (i.e. < 230 m). α is never as large as 5/6 as would be required for 3D turbulence. For 50% of the entire time $\alpha = 0.41$ with the lowest 50% quartile located in August with $\alpha = 0.40$ and the highest in March with $\alpha = 0.43$, a relatively small annual variation. The largest values for α found in summer, occurring from 09 h–18 h, are consistent with there being an agitated troposphere caused by the Sun’s thermal heating at these times. Conversely, the smaller values in winter suggest a thinner phase screen than in summer, on

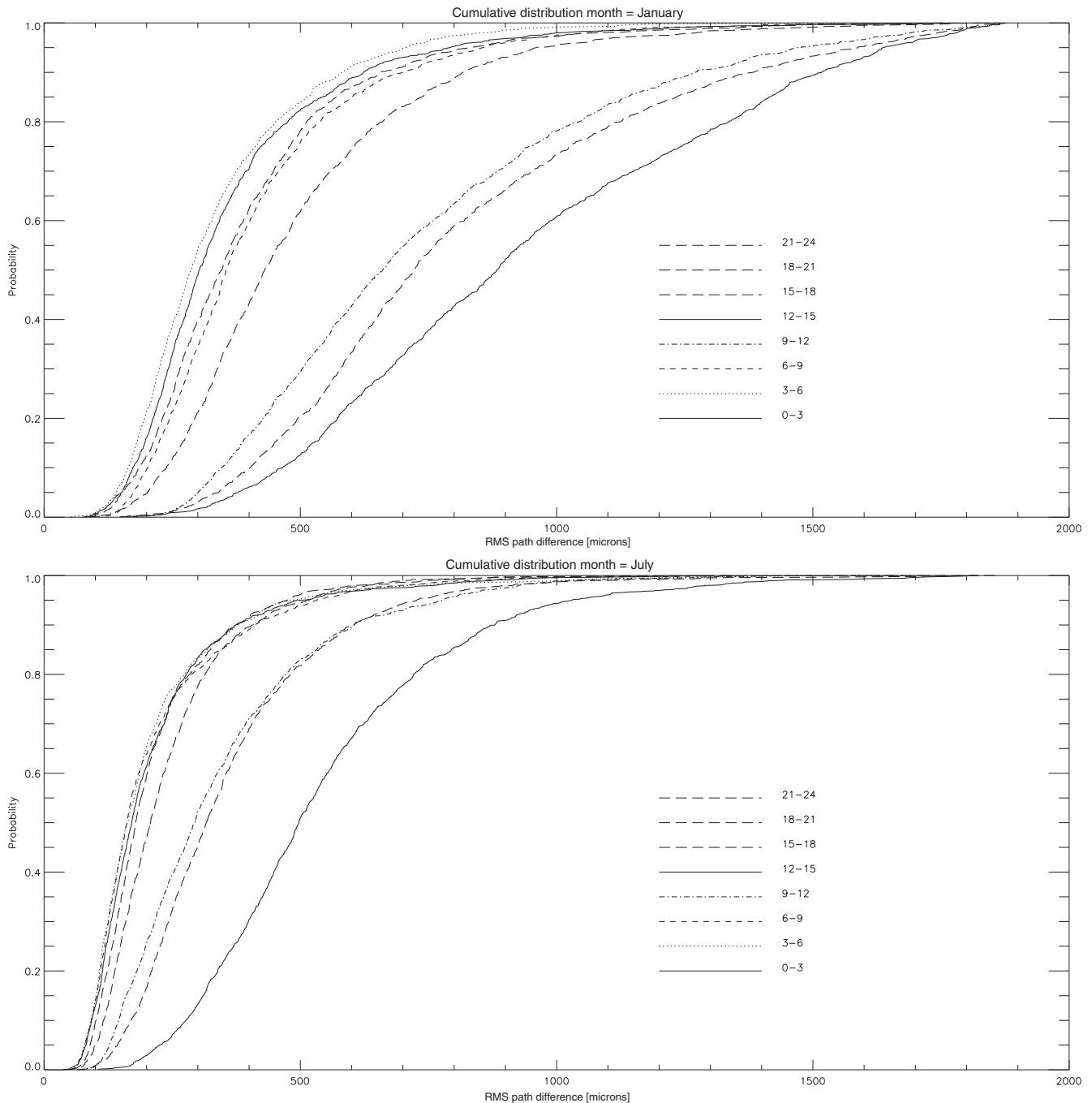


Figure 4. Cumulative distributions of the zenith rms path differences in μm for time of day (in 3-h bands) in summer (January, top) and winter (July, bottom). Summer nights have similar values to winter days.

average, presumably due to there being lesser thermal heating in winter.

5.2 Phase screen speed

From the LPSF the corner time is obtained, then, by using Equation 3, the phase screen speed. In Figure 8, we show the histogram and cumulative distribution for the phase screen speed on a monthly basis for the entire 9 yr of data. The

histogram conforms to a classical Weibull distribution:

$$F(x, \beta, \gamma) = \frac{\beta}{\eta} \left(\frac{x - \gamma}{\eta} \right)^{\beta-1} e^{-\left(\frac{x - \gamma}{\eta} \right)^\beta}. \quad (9)$$

In this case a location parameter $\gamma \sim 3.5$, scale parameter $\eta \sim 2.0$, and shape parameter of $\beta \sim 1.6$ provide a good fit. This later value is between that for an exponential ($\beta = 1$) and a Rayleigh ($\beta = 2$) distribution. Wind speed data throughout the Meteorological literature is often

Table 2. The maximum and minimum quartiles for each 3-h time band, together with the month in which they occur, as extracted from Table 1.

Time/ Quartiles	Minimum quartiles in μm							
	00	03	06	09	12	15	18	21
25%	116	111	112	194	328	211	145	125
	May	May	May	Jun	June	June	May	May
50%	159	154	160	286	461	292	194	171
	May	May	May	Jun	Jun	Jun	May	May
75%	224	234	233	405	628	405	269	236
	May	May	May	Jun	Jun	Jun	May	May
	Maximum quartiles in μm							
25%	229	212	268	504	653	543	319	254
	Jan	Jan	Nov	Dec	Feb	Jan	Jan	Feb
50%	303	286	367	704	931	725	430	347
	Jan	Jan	Dec	Dec	Feb	Feb	Jan	Jan
75%	423	412	536	1008	1259	1026	604	478
	Jan	Jan	Dec	Dec	Feb	Jan	Jan	Jan

seen to follow a Weibull distribution (Murthy, Xie, & Jiang 2004).

Of note is the quantisation effect due to the time resolution (5 s) as well as the minimal seasonal variation in the phase screen speed. While in winter there are slightly higher wind speeds, the difference with summer is small, of order 0.5 m/s. In Figure 9 the cumulative distributions for the phase screen speed are shown for 3-h time bands in January and in July. A complementary picture to Figure 4 is seen. The lowest wind speeds occur during the same time span, between midnight and 9 am, when the rms path variations are also lowest. This is consistent with the flow being more laminar and thus less turbulent when it is associated with lower wind speeds.

5.3 Seeing limitations

Another important effect caused by tropospheric phase fluctuations is the limitation it imposes on the ability of the telescope to spatially resolve structure. Analogous to the seeing limitations caused by microthermal fluctuations in the optical regime, in the radio regime the water-vapour-induced phase delays cause a positional variation in the source observed because interferometric phase corresponds to the measurement of a point source (Perley 1989).

Following Carilli & Holdaway (1999), the half-power point in the visibility (Equation 8) occurs when $\Phi_{\text{rms}} = 1.2$ radians. From the seeing monitor measurements $\Phi_{0,\text{rms}} = p_{\text{rms}}/\lambda_0 \times 2\pi$ where p_{rms} are the rms path differences. Then, using Equation 7 and substituting in Equation 8 we can define a half-power baseline length (HPBL), $b_{1/2}$, by

$$b_{1/2} = \left(\frac{1.2\lambda}{K\lambda_0} \right)^{1/\alpha}, \tag{10}$$

with $K = \Phi_{0,\text{rms}}$.

The seeing is then given by

$$\theta = \frac{\lambda}{b_{1/2}}. \tag{11}$$

For example, on 2006 June 16, the phase structure function parameters were determined, as shown in Figure 5. The rms path difference was 227 μm , the corner time 110 s, and therefore the phase screen speed was 2.1 m/s and Kolmogorov exponent $\alpha = 0.53$. From this we obtain $K = 0.095$ radians (5.4°). For observations to be conducted at 3.3 mm (90 GHz), with $\lambda_0 = 15$ mm and $b_0 = 230$ m, the half-power baseline is 1.6 km and the corresponding millimetre seeing is 0.43 arcsec.

In Table 4, the limiting seeing values for an observing wavelength $\lambda = 3.3\text{mm}$ are listed based on the calculated half-power baselines for the median Kolmogorov exponents and path difference rms for each hour of the day during the best (June) and worst (December) periods of the year. The best seeing conditions are of order 0.3 arcsec and occur in June, between midnight and 6 am. The worst conditions occur during December midday, and are about 2.5 arcsec.

5.4 Correlations between derived parameters

In Figure 10 we present scatter plots to search for correlations between the parameters derived from the seeing monitor data. As is evident from these, only the saturation path length and the rms path difference correlate strongly, with the slope to a linear fit being 1.44 ± 0.01 , i.e. close to $\sqrt{2}$, as anticipated (see Section 5.1). There is a weak positive correlation between the Kolmogorov exponent and the rms path difference and a weak anti-correlation between it and phase screen speed. While the Kolmogorov exponent allows us to infer that the turbulence generally approximates 2D

Table 3. The median Kolmogorov exponents α for each 3-h time band during January and July, and over the whole year.

Month/hour band	Kolmogorov exponent α								
	00–03	03–06	06–09	09–12	12–15	15–18	18–21	21–24	Median
Jan	0.388	0.385	0.397	0.449	0.466	0.449	0.422	0.400	0.419
Jul	0.385	0.384	0.383	0.412	0.446	0.424	0.395	0.387	0.402
Overall	0.387	0.384	0.389	0.438	0.458	0.441	0.410	0.396	0.413

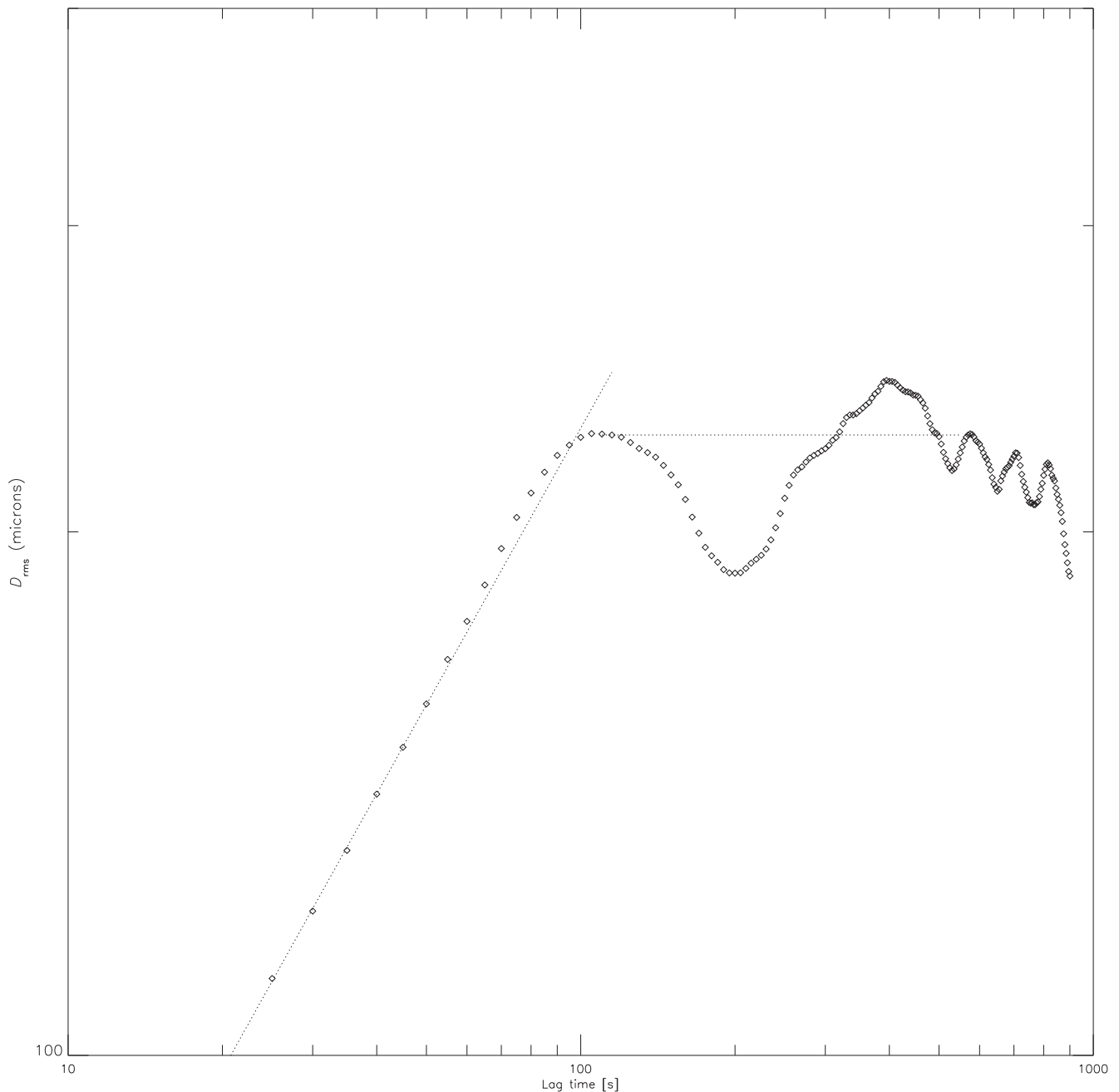


Figure 5. An example of lag phase structure function on 2006 June 16, plotting lag time (s) against the rms path difference (μm). The horizontal line shows where the saturation path length was determined as the structure function's first peak and the slope of the fit to the rising portion yields the Kolmogorov exponent. In this example, the rms path length was found to be $227 \mu\text{m}$ at the corner time $t_c = 110 \text{ s}$, yielding a phase screen speed of 2.1 m/s and Kolmogorov exponent $\alpha = 0.53$.

behaviour, the phase screen speed is not useful in inferring the millimetre-site conditions.

Only every 30th point is shown in Figure 10 for clarity. The alternative (and as shown by Middelberg et al. 2006) is to average the data in bins. This, however, results in a relatively smooth line which makes the correlation between the parameters appear to be better than it actually is.

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5.5 Available observing time in the millimetre wavebands

To successfully undertake interferometry on a given baseline and frequency, the visibility efficiency

$$\epsilon = \frac{\langle V_m \rangle}{V} = e^{-\sigma_v^2/2}, \quad (12)$$

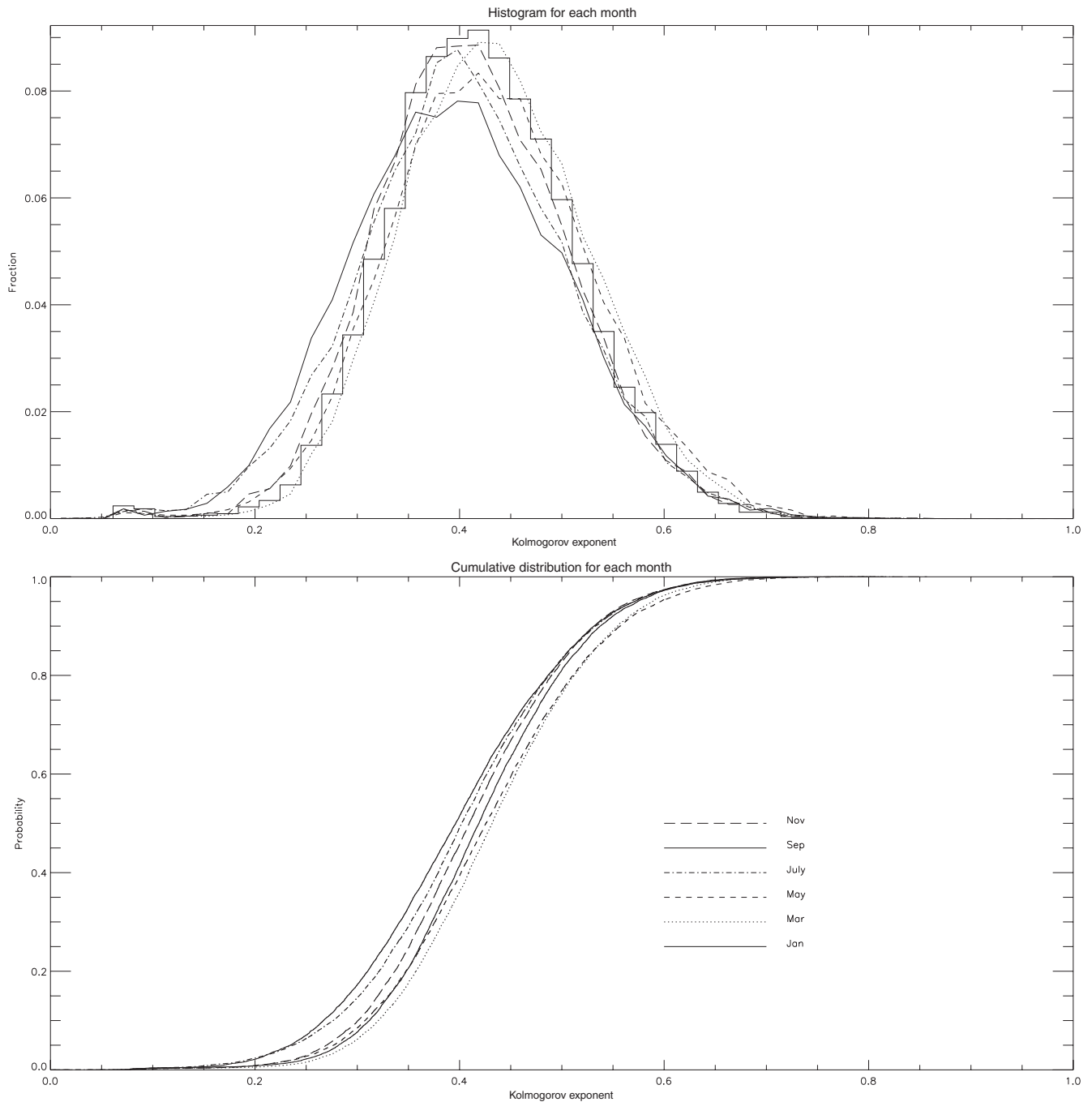


Figure 6. Histogram and cumulative distribution of the Kolmogorov exponent α as a function of the month. Seasonal variations are seen to be small. For clarity, only every other month is shown.

where σ_ϕ is the phase fluctuation rms and V the visibility, needs to be $\gtrsim 0.9$ (this is equivalent to a phase noise $\sigma_\phi = 30^\circ$). From this value the limiting rms path noise fluctuation can be determined for a given waveband ($d_{\text{rms}} = \frac{\phi_{\text{rms}}}{360} \lambda_{\text{obs}}$). In turn, by applying Kolmogorov turbulence theory (as discussed in Section 3) the maximum rms path noise on the 230-m seeing monitor baseline can be determined for interferometry to be possible for a given frequency and baseline, applying

Equation 13, as

$$d_{\text{rms}} = d_{\text{seemon-rms}} \times \left(\frac{b}{b_{\text{seemon}}} \right)^\alpha, \quad (13)$$

where $b_{\text{seemon}} = 230$ m. We have taken $\alpha = 0.4$, the median value found for the Kolmogorov index. These limiting values for the path fluctuations are shown in the top left plot of Figure 11 for observations at 22, 45, and 90 GHz.

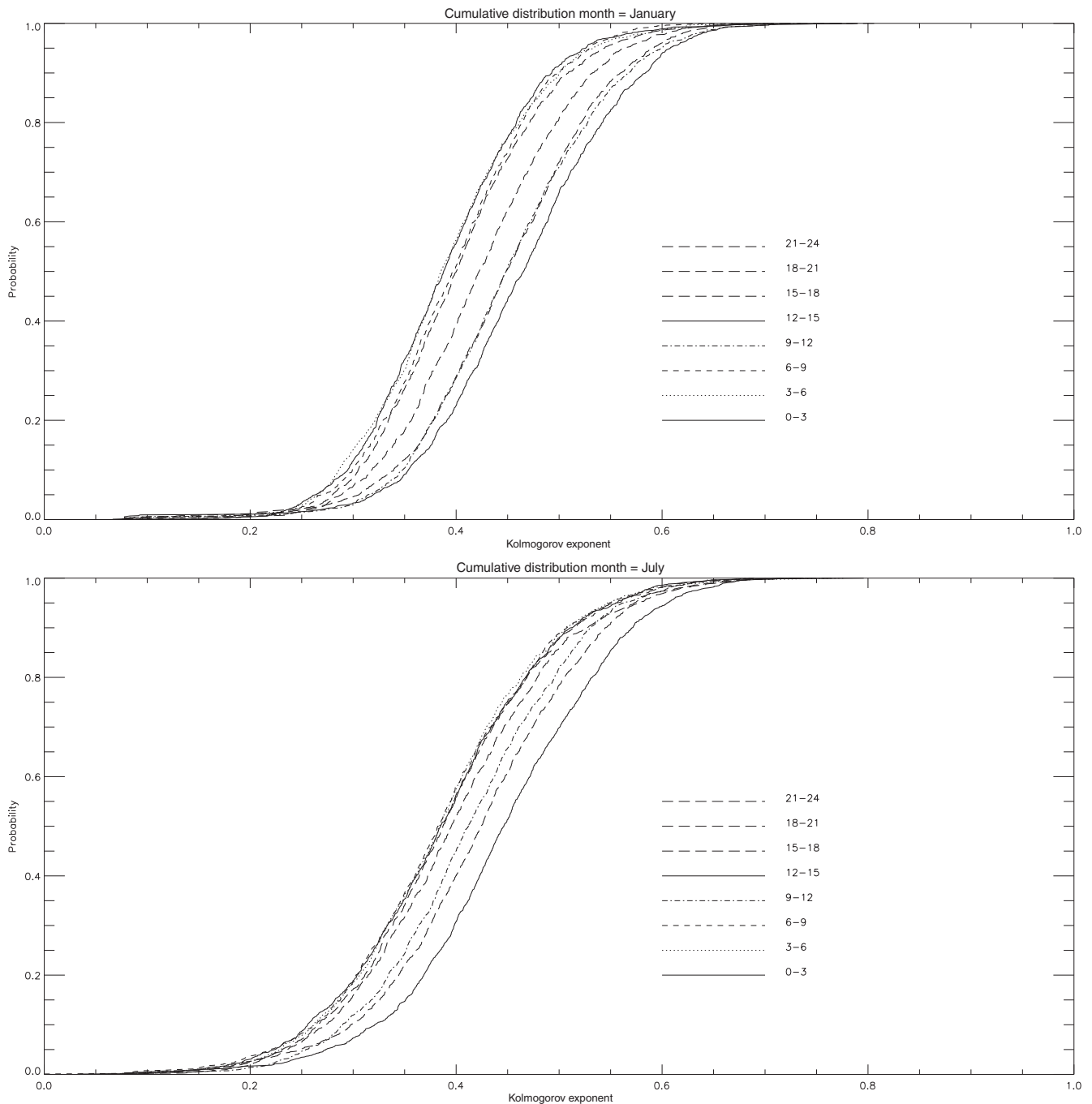


Figure 7. Cumulative distribution of the Kolmogorov exponent α for the months of January (top) and July (bottom) shown in 3-h time bands. Diurnal variations are relatively small, but are larger than the seasonal variations shown in Figure 6.

Furthermore, a set of WVRs has recently been developed for the ATCA (see Indermuehle et al. 2013). In that paper it is demonstrated that phase fluctuations which produce a visibility efficiency of 0.5 (i.e. phase fluctuations of 67°) can be corrected to provide a visibility efficiency of 0.9 using the WVRs. This, in turn, can be used to determine the limiting rms path fluctuations between the seeing monitors for successful observations when the WVRs are in use. These

values are also shown in the top left plot of Figure 11. As can be seen, the limiting rms path values with WVRs in use are significantly higher than the limits without.

The rms path lengths we have presented in this paper can also be used to determine for what fraction of the time it is possible to observe on a given frequency and baseline as a function of hour and/or season, with and without the use of the WVRs. The remaining three plots in Figure 11 show the

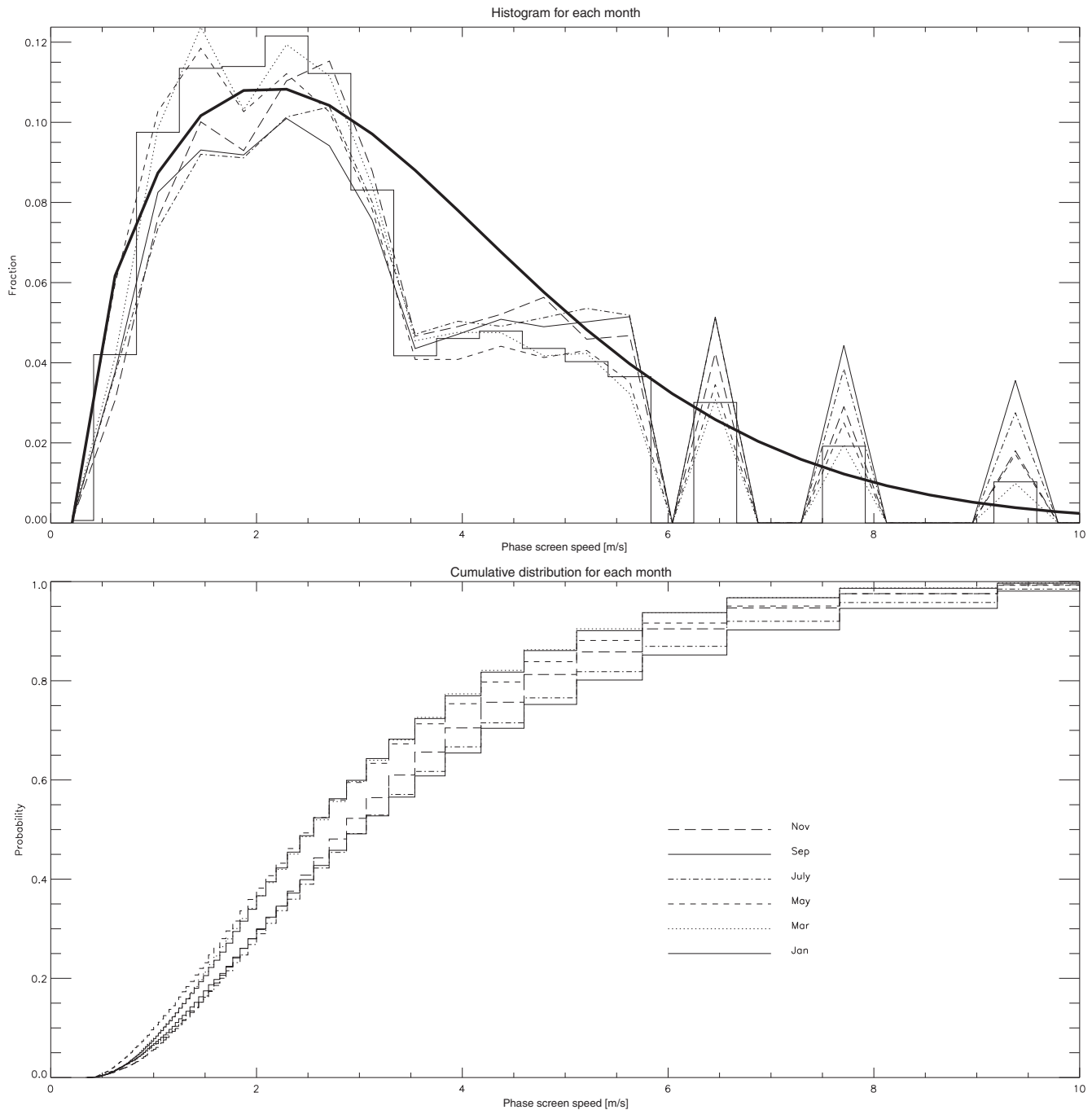


Figure 8. Histogram and cumulative distribution of the phase screen speed distribution as a function of month (for clarity only every other month is shown). Note that the quantisation effect arises out of the time resolution of 5 s. The thick line overlay in the histogram plot shows the Weibull distribution as shown in Equation 9, with location parameter $\gamma \sim 3.5$, shape parameter of $\beta \sim 1.6$ and scale parameter $\eta \sim 2.0$.

results for three cases: (i) averaged over the whole year, (ii) in January, and (iii) in July.

Some striking conclusions can be drawn from inspection of these plots. For instance, at 90 GHz observing efficiency using WVRs becomes similar to that at 45 GHz without their use. At 90 GHz observing efficiency on a 3-km baseline is only possible for about 7% of the time in winter, and so is only rarely attempted. This could rise to half the time if WVRs

were to be used. 45-GHz interferometry is possible out to 6-km baselines in winter for one-quarter of the time, and only possible in summer on short baselines of less than 1 km. With the use of WVRs, 45-GHz observations in summer could become a regular practice out to 3 km baselines (and at 6 km for one-third of the time). 22-GHz interferometry can be conducted through the winter on all baselines for two-thirds of the time, but only for one quarter of the time in summer.

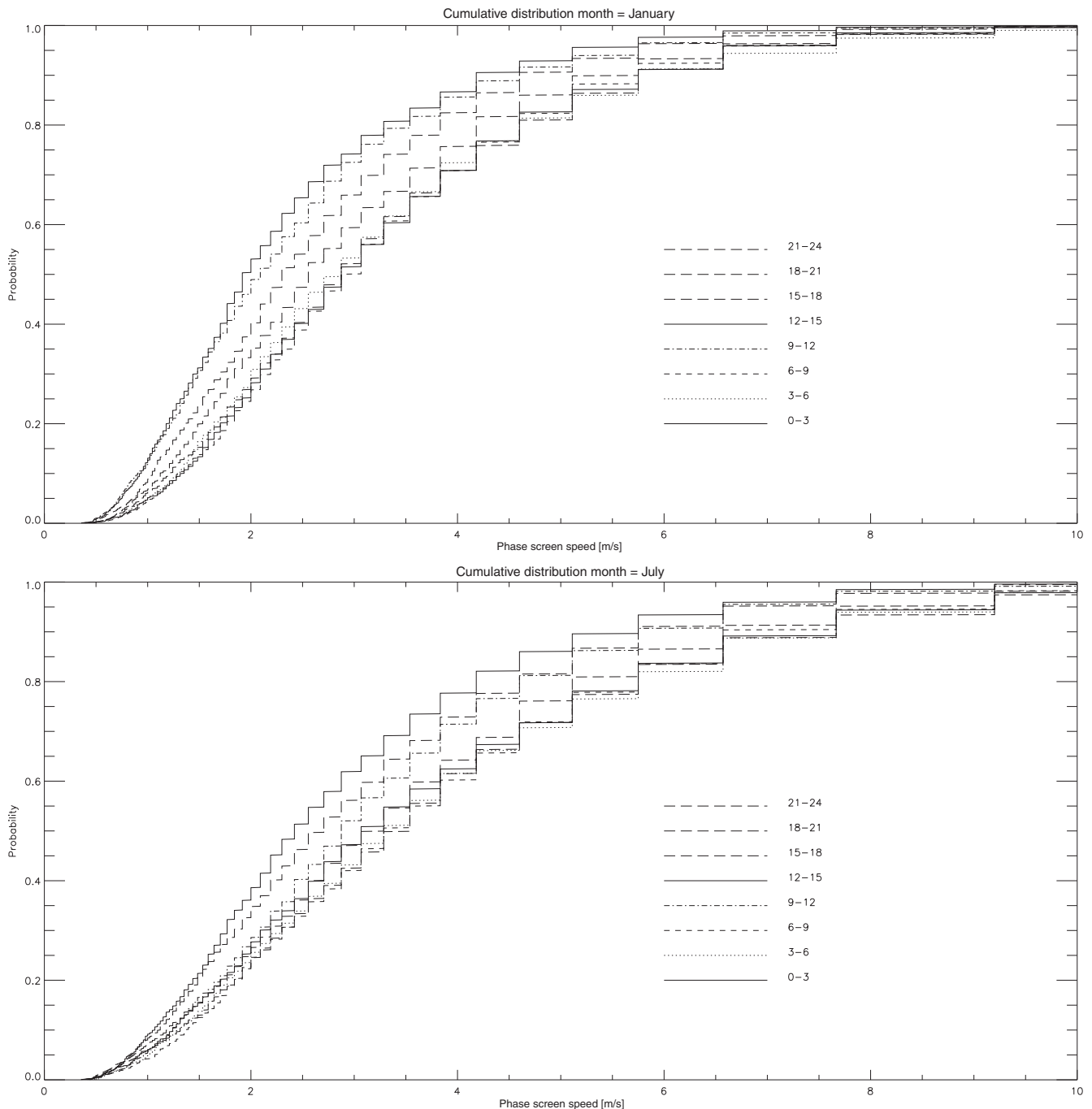


Figure 9. Cumulative distributions for the phase screen speed as a function of time of day (in 3-h bands) in summer (January, top) and winter (July, bottom).

With the use of WVRs it could be conducted on all baselines, for most of the time.

Table 1 and Figure 11 can also be used together to estimate the useable observable fractions in every 3-h time band for each month of the year. To do so, use the top left plot in Figure 11 to determine the maximum rms path length variations for the observing frequency (i.e. 22, 45, or 90 GHz) and antenna baseline of interest, both with and without the use of WVRs. For the time band and month under consideration for an observing programme compare these values to the 25%,

50%, and 75% quartiles listed in Table 1 to find the relevant observing fractions. For example, for a 3-km antenna baseline at 90 GHz, from Figure 11 we infer that path fluctuations of up to $100 \mu\text{m}$ and $220 \mu\text{m}$ can be used, without and with the use of WVRs, respectively. Without WVRs, from Table 1 it can be seen that the lowest quartile value for June is $118 \mu\text{m}$, so 90-GHz interferometry on a 3-km baseline will only rarely be possible. Table 1 also shows that the median night time path fluctuations are less than $200 \mu\text{m}$ through the winter months. Thus in winter, the use of WVR assisted

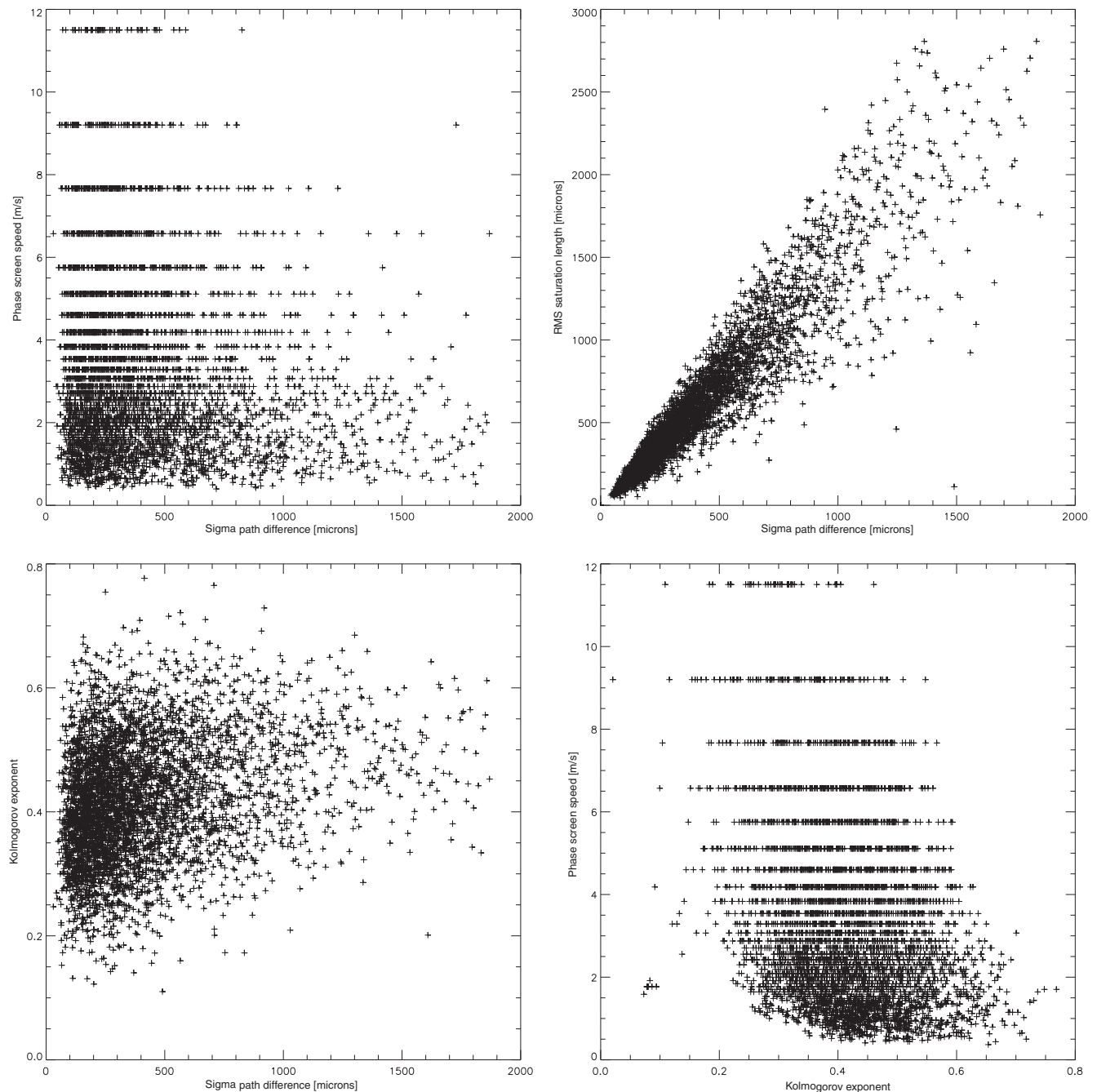


Figure 10. Correlations between the derived values for the rms path difference and phase screen speed (top left), rms path difference and saturation path length (top right), rms path difference and Kolmogorov exponent (bottom left) and the Kolmogorov exponent and phase screen speed (bottom-right). The corresponding correlation coefficients are -0.12 , 0.93 , 0.25 , and -0.28 respectively. Note that the quantisation in two of the plots arises from the limited values possible for the phase screen speed (see text). For clarity, only every 30th point is plotted.

interferometry at 90 GHz on this baseline will be possible during nights for more than 50% of the time, and indeed in daylight for more than 25% of the time (aside from 12 h–15 h). In January, on the other hand, such observations should only be attempted during night-time hours, between 21 h and 06 h, but even then will only be possible for less than 25% of the time.

In summary, significant gains in observing time and increases in useable baseline length are seen to be achievable in all seasons by using WVRs for all three of the millimetre observing bands. The useable observing period for a given baseline and frequency is increased by about 4 months a year. This would make interferometry possible throughout the year in all the millimetre wavebands.

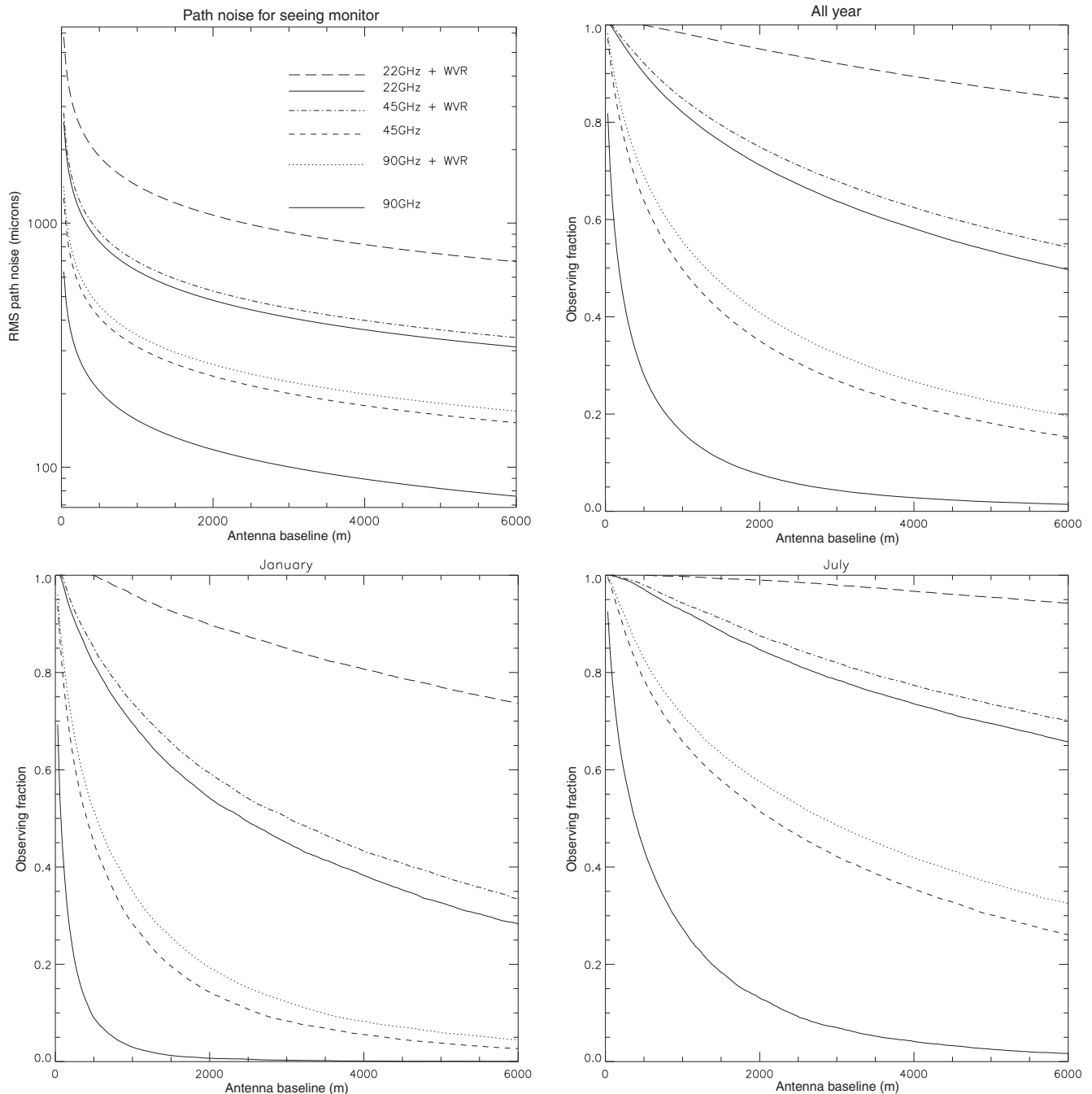


Figure 11. Observing fractions as a function of antenna baseline, frequency and season for the ATCA, with and without the use of the water vapour radiometers (WVRs) to provide millimetre-wave phase correction. The top left plot shows the rms path length noise as measured by the seeing monitor that is required to conduct successful observations on a given baseline and frequency, with and without the WVRs. The other three plots convert these into a fraction of the available observing time that could be used for the three cases: (i) the yearly average, (ii) January (i.e. summer), and (iii) July (i.e. winter). The calculations have been performed for three frequencies: 22, 45, and 90 GHz and assume the median value for the Kolmogorov exponent $\alpha = 0.4$. As described in Section 5.5, this Figure can be used together with Table 1 to estimate observing time fractions for any time band and month at one of these frequencies, on an antenna baseline of interest.

6 CONCLUSION

The magnitude of the path length fluctuations caused by variations in the water vapour columns between two antennae of an interferometer determines when observations may be suc-

cessfully attempted at a given frequency. We have analysed 8.5 yr of data from a seeing monitor at the ATCA site near Narrabri to provide statistics on the site conditions for this telescope, extending the initial 1-yr study conducted by Mid- delberg et al. (2006).

Table 4. The hourly seeing values in arcseconds for June and December for observations at $\lambda = 3.3$ mm. These are calculated using the median values determined for the rms path difference and Kolmogorov exponent for each of these time periods, as explained in Section 5.3.

Time AEST [h]	Seeing θ ["]	
	Jun $\bar{\theta}$	Dec $\bar{\theta}$
0–1	0.32	0.53
1–2	0.33	0.50
2–3	0.32	0.48
3–4	0.33	0.49
4–5	0.34	0.52
5–6	0.35	0.59
6–7	0.35	0.80
7–8	0.36	1.18
8–9	0.44	1.68
9–10	0.65	2.13
10–11	0.91	2.48
11–12	1.12	2.56
12–13	1.19	2.53
13–14	1.10	2.27
14–15	0.90	2.04
15–16	0.65	1.69
16–17	0.49	1.36
17–18	0.42	1.03
18–19	0.38	0.78
19–20	0.38	0.69
20–21	0.36	0.64
21–22	0.34	0.64
22–23	0.33	0.66
23–24	0.31	0.65

Two monitors located 230 m apart measure the signal from a geosynchronous satellite. We use the fluctuations in the path differences across this baseline to determine the rms value for the path difference, the Kolmogorov exponent, α , for the turbulence and the phase screen speed, under the ‘frozen-screen’ assumption (Taylor 1938) for the passage of the turbulent cells across the site. We also determine the seeing at 3.3 mm from the half-power points in the visibility.

Considerable variations in the rms path differences between the two monitors are found between summer and winter, as well as between day and night time, with 25% (first) quartile values ranging from a low of 111 μm in May between 03 h and 06 h, to a high of 653 μm in February between 12 h and 15 h. However, typical fluctuations during summer nights, ~ 300 μm , are similar to those occurring during winter days; in other words millimetre-wave interferometry can typically be undertaken during summer nights. Variations in the Kolmogorov exponent are much less, with a median value of ~ 0.4 and variations of $\sim \pm 0.04$. Similarly, variations in the phase screen speed are relatively small, with mode values found of ~ 2 m/s.

The Kolmogorov exponent is usually closer to the value expected for 2D turbulence, $1/3$, than that for 3D turbulence, $5/6$, consistent with the turbulent layer containing

the fluctuating water vapour being relatively thin, as per the frozen-screen hypothesis. The lowest values for α are found in winter, as would be expected for a less agitated troposphere than in summer.

The seeing at 3.3 mm shows significant seasonal and time of day variations, depending on the path length rms values. It ranges from a low of ~ 0.3 arcsec during winter nights to ~ 3 arcsec in the middle of summer days.

We have calculated the maximum path length fluctuations that can be tolerated for interferometric observations to be conducted with the ATCA on a given antenna baseline and frequency (in particular, the three receivers at 22, 45, and 90 GHz). Using the measured path length fluctuations, and assuming Kolmogorov turbulence with the median value measured for α of 0.4, we then estimate the useable observing fractions for time of day and month of the year. We also do this assuming that phase variations can be partially rectified using the ATCA WVRs. Gains are possible on all baselines, with useable observing periods typically increased by about 4 months per year for any given combination of frequency/baseline using the WVRs. In particular, observations at 90 GHz could be conducted out to the currently maximum 3-km baseline of the ATCA during winter months, or out to 6 km were the six kilometre antenna to be equipped with 3-mm capability.

With the use of WVRs, observations will often be undertaken with baselines longer than now being used. This would then facilitate flexible observing as when the conditions are not suitable for millimetre observations, the longer baselines will be conducive for many centimetre band programmes, as was proposed by Hall et al. (1992) in their original proposal for operation of the ATCA at millimetre wavelengths.

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