AB INITIO MOLECULAR DYNAMICS STUDY OF Fe-CONTAINING SMECTITES

Xiandong Liu 1,* , Evert Jan Meijer 2,† , Xiancai Lu 1 , and Rucheng Wang 1

¹ State Key Laboratory for Mineral Deposit Research, School of Earth Sciences and Engineering, Nanjing University, Nanjing 210093, P.R. China

² Van't Hoff Institute for Molecular Sciences and Amsterdam Center for Multiscale Modeling, University of Amsterdam, Nieuwe Achtergracht 166, 1018 WV Amsterdam, The Netherlands

Abstract—In order to identify the influences imposed by Fe substitution, density functional theory-based Car-Parrinello molecular dynamics simulations were employed to study both oxidized and reduced Febearing smectites. The following basic properties were investigated: local structures in the clay layer, hydroxyl orientations, and the vibration dynamics of H and Si. Structural analyses indicated that the average Fe–O bond lengths are ~2.08 Å and 2.02 Å in the reduced and oxidized models, respectively, and the Fe substitutions did not affect the coordination structures of the Al–O and Si–O polyhedra. For hydroxyl orientations, Fe(III) substitution had no obvious influence but Fe(II) forces the coordinated hydroxyls to present a wide-angle distribution. Furthermore, the present work has shown that both substitution has no effect on the Si–O stretching, while Fe reduction causes a blue-shift of the out-of-plane stretching mode. The results provide quantitative constraints and clues for future research.

Key Words—CPMD, Density Functional Theory, Hydroxyl Orientation, Iron, Smectite, Vibration Dynamics.

INTRODUCTION

Iron-bearing smectites are of great interest in terms of both basic research and industrial applications (Stucki, 2006). These ubiquitous phyllosilicates form an important pool of Fe in natural ecosystems and have many connections with other inorganic (e.g. oxides and carbonates) and organic phases such as bacteria (Alimova et al., 2009; Murad and Fischer, 1988). Due to susceptibility to reduction and oxidation reactions, Fe-smectites participate in many reactions and processes, and therefore play active and important roles in the geochemical cycle of Fe (Murad and Fischer, 1988). In industry, smectites are used widely in civil engineering, in the synthesis of composite materials, and in oil exploration. Because of the reduction capability, Fe-bearing smectites are also used in degrading organic pollutants and reducing heavy-metal cations (Cervini-Silva et al., 2000; Favre et al., 2006; Jaisi et al., 2008a, 2008b; Stucki, 2006; Peretyazhko et al., 2008).

Many experimental works have been conducted to study the properties of Fe-smectites and, in particular, the effects of Fe oxidation states. Reduction by Fe is widely acknowledged to alter many properties, such as cation exchange and fixation capacity, specific surface area, layer stacking order, swelling capacity, and surface

 * E-mail address of corresponding author: xiandongliu@gmail.com
 * e.j.meijer@uva.nl
 DOI: 10.1346/CCMN.2010.0580109 acidity (Stucki et al., 2002; Stucki, 2006). Reliable physical mechanisms for these findings have been difficult to establish, however. In fact, to better understand these issues, some fundamental questions must be answered: (1) How do Fe substitutions (Fe(III) and Fe(II)) affect the crystal structures of smectites? The EXAFS (Extended X-ray Absorption Fine Structure) technique can provide atomic-level structural information, but unequivocal interpretation of EXAFS data is still very difficult (Denecke, 2006). (2) How do the substitutions affect hydroxyl orientation? This is very important for many properties of smectites such as adsorptive properties, and at present is nearly impossible to measure directly. (3) How do the substitutions influence the vibration dynamics of smectites? Techniques such as infrared (IR) and Raman spectroscopy are commonly used in clay sciences, but sometimes the band assignments are questionable and distinguishing the contribution of a single ion species is very difficult (Balan et al., 2001; Blanchard et al., 2008; Gates, 2008). In addition, smectites are usually fine-grained, poorly crystallized, and always coexist with other minerals, which also increase the problems with data collection in experimental measurements.

Ab initio simulations have been employed on phyllosilicate systems and have proven to be effective for complementing experiments and formulating hypotheses (*e.g.* Cygan and Kubicki, 2001; Kubicki and Bleam, 2003; Larentzos *et al.*, 2007; Boulet *et al.*, 2006). Only a few studies have focused on Fe-containing phyllosilicates, however. The fundamental questions presented above are rarely investigated and are, therefore, poorly documented. Hernandez-Laguna et al. (2006), for example, considered the influence of Fe(III) substitution on the total energies of phyllosilicates; Teppen et al. (2002) optimized the structure of trichloroethene-intercalated nontronite (an end-member of Fe-smectites in which almost all octahedral metals are Fe); and Rosso and Ilton (2003) focused on electron transfer in Fe-containing micas. The present work aimed to understand those questions with quantum molecular dynamics simulation (Car and Parrinello, 1985). Density functional theory-based molecular dynamics simulations were performed to study both the reduced and oxidized Fe-smectites at finite temperature. The structural and vibrational properties were derived from the trajectories to investigate the effects of Fe substitutions. By comparing the reduced and oxidized systems, the influences of Fe oxidation states were discussed in detail.

METHODOLOGY

The models

The smectite model was derived from the report of Viani *et al.* (2002). The simulated system (Figure 1) consisted of two unit cells $(2a \times b \times c)$ with the crystallographic parameters: a = 5.18 Å, b = 8.98 Å, c = 10 Å, $\alpha = \beta = \gamma = 90^{\circ}$. The unit-cell formula is Si₈[Al_{3.5}Fe_{0.5}]O₂₀(OH)₄ where one Al atom is replaced by an Fe in the octahedral sheet. With a single Fe atom, the magnetic ordering can be ignored. Some previous simulation works showed that this system size can properly reproduce the crystal properties of phyllosilicates (*e.g.* Churakov, 2006).



Figure 1. Snapshot of the reduced model.

In the oxidized system, the clay layer is neutral and the interlayer interaction is mainly van der Waals' interaction, so the *c*-axis length has no obvious influence on the coordination structure in the layer (Churakov, 2006). In the reduced form, the *c*-axis length is consistent with the experimental measurement of Li-smectite (Calvet, 1973). In this case, a Li⁺ cation was inserted in the interlayer space to compensate the negative charge of the clay layer. The cation can be above either the substituted (AlOHFe) or the non-substituted (AlOHAl) six-member ring. In the present study the former is considered and has a slightly lower energy. During the simulation, the interlayer cation always hovered above the six-member ring and did not escape. Smectites generally contain some water molecules in the interlayer spaces, which may have some effect on the vibration dynamics of the substrate (Larentzos et al., 2007). No interlayer water molecules were included in the present work in order to exclude the possible additional influences.

Simulation details

The electronic structure was calculated in the framework of density functional theory. The exchangecorrelation was described by the Becke-Lee-Yang-Parr (BLYP) function, which adopts the local density approximation augmented with the generalized gradient approximation for the exchange part proposed by Becke (1988) and for the correlation part by Lee *et al.* (1988). The BLYP function has been used to simulate the phyllosilicate systems and produces the bond lengths agreeing with the Perdew-Burke-Ernzerhof functional results and experiments (Churakov, 2006; Liu et al., 2008). The BLYP has been shown to describe H-bonding interactions accurately, e.g. it gives the best results in liquid water (Sprik et al., 1996). The Trouiller-Martins norm-conserving pseudopotentials (Troullier and Martins, 1991) were used to describe the interaction of the valence electrons and the core states, and the Kleinman-Bylander scheme (Kleinman and Bylander, 1982) was applied. The kinetic energy cutoff up to 90 Ry and one K-point in the center of the Brillouin zone ensure a convergence of the bond lengths within 0.02 Å. Iron is in the high-spin state in both oxidized and reduced models.

The molecular dynamics simulations were carried out using the Car and Parrinello approach as implemented in the *CPMD* package (*CPMD version* 3.11) (Car-Parrinello, 1985). The equation of motion was integrated with a time step of 6 a.u. (0.145 fs). In order to use this large time step, the deuterium mass was applied for all hydrogen atoms. The fictitious electronic mass was set at 800 a.u. which maintained the adiabatic conditions. The Nosé-Hoover thermostat was applied to control temperature at 298 K. After equilibration lasting 1 ps, the production number, volume, temperature-constant molecular dynamics runs were performed for 2.2 ps for both systems.



Figure 2. Al-O and Fe-O pair distribution functions in the octahedral sheet of the (a) oxidized and (b) reduced models.

RESULTS AND DISCUSSION

Structures of the clay layer

The pair distribution function of Fe(III)–O (Figure 2a) exhibits a narrow distribution with an average of ~2.02 Å; Fe(II)–O (Figure 2b) presents a slightly wider distribution and shows the peak value at ~2.08 Å. The results are consistent with a previous experimental study (Manceau *et al.*, 2000) in which the bond lengths of Fe(III)–O and Fe(II)–O of nontronite were found to be ~2.01 Å and ~2.10 Å, respectively, by EXAFS analyses. In Figure 2, Al_{Fe} denotes the Al atom connecting with Fe and Al_{AI} refers to the remaining Al atoms. The four Al–O curves are almost the same, ranging from ~1.7 to 2.1 Å in length and are centered at ~1.92 Å, which agrees well with previous reports on systems without Fe substitution (Sainz-Diaz *et al.*, 2002; Refson *et al.*, 2003; Larentzos *et al.*, 2007).

 Si_{Fe} in Figure 3 denotes the Si atom connecting with the Fe atom *via* a bridging oxygen atom, and Si_{A1}

indicates those connecting with Al through bridging oxygen. The four Si–O profiles comprise very similar sharp peaks ranging from 1.5 to 1.7 Å and centered at ~1.62 Å. This result is consistent with the previous experiments and quantum simulations (Sainz-Diaz *et al.*, 2002; Refson *et al.*, 2003; Larentzos *et al.*, 2007). The results above suggest that neither Fe(III) nor Fe(II) substitution influences the local structures of the neighboring octahedra and tetrahedra.

Orientations of hydroxyls

The hydroxyl orientation is characterized by the angle between the OH vector and the (100) direction, as depicted in Figure 4. H_{Fe} refers to the H from the hydroxyls coordinating with the Fe atom and H_{A1} denotes the other H atoms. Because one Fe atom is in the model, the total number of H atoms is given by 2 H_{Fe} and 6 H_{A1} . In the oxidized smectite (Figure 4a), both curves represent similar distributions of the OH vector ranging from ~0 to 60° and greatest at ~15–20°. This



Figure 3. Si-O pair distribution functions in the tetrahedral sheets of the (a) oxidized and (b) reduced models.



Figure 4. Distributions of hydroxyl orientations in the (a) oxidized and (b) reduced models.

distribution is consistent with the previous *ab initio* molecular dynamics simulations of pyrophyllites (Larentzos *et al.*, 2007) and also matches previous static density functional theory (DFT) calculations (Bridgeman *et al.*, 1996; Stackhouse *et al.*, 2001; Refson *et al.*, 2003; Sainz-Diaz *et al.*, 2004). In the reduced model, both curves show distributions ranging from ~0 to 40° and $H_{Fe(II)}$ shows another prominent maximum at ~90° (Figure 4b).

Density profiles of hydroxyl hydrogen (Figure 5), in which the middle plane of the clay layer was taken as the origin, revealed very similar distributions for the H_{Fe(III)} and the two H_{A1}, ranging from ~0.5 to 2 Å away from the middle plane (Figure 5a). Remarkably, the H_{Fe(II)} curve exhibited a greater distribution extending to 3 Å (Figure 5b), which is consistent with the distributions reported in Figure 4b.

As noted in the past (Larentzos *et al.*, 2007), the hydroxyls point toward the apical oxygen to maximize the electrostatic interaction, leading to a low-angle

distribution. The Fe reduction causes the clay layer to be negatively charged and such octahedral substitution is well recognized as being able to make all framework oxygens more negative; therefore, this may be responsible for the narrower low-angle distribution in the reduced model. On the other hand, the Fe-O interaction is weakened after reduction because the Fe-O distance increases, which can impose a greater mobility on the hydroxyl. All of these factors can contribute to the highangle orientation distribution of the $OH_{Fe(II)}$. The halfwidth of the clay layer is ~3.3 Å, so when the $OH_{Fe(II)}$ orients to the large-angle region, it is very close to the surface. Structural hydroxyls can significantly affect the interfacial properties, so the special orientation probably contributes to some unique properties of Fe-bearing smectites.

Vibration dynamics

To characterize the vibration dynamics, vibration density of states (VDOS) were calculated from a Fourier



Figure 5. Density distributions of hydroxyl hydrogen in the (a) oxidized and (b) reduced models. The calculations were done $\rho(z) = \langle N(z - \Delta z, z + \Delta z/2) \rangle / (\Delta z \times S)$, where $\langle N(z - \Delta z, z + \Delta z/2) \rangle$ is the averaged atom number occurring in the interval of $(z - \Delta z, z + \Delta z/2)$ and S is the basal surface area.



Figure 6. VDOS of hydrogen in the (a) oxidized and (b) reduced models.

transform of the atomic velocity auto-correlation functions (VACF) (Allen and Tildesley, 1987). Such a method was used in simulation analyses of minerals (*e.g.* Boek and Sprik, 2003; Wang *et al.*, 2003; Larentzos *et al.*, 2007). Generally, the peaks of VDOS indicate some vibration modes which could be either IR or Raman active. Hence, the peak positions can be compared with experimental spectra. The present study focuses on the spectra of H and Si, which are important for experiments (Stucki, 2006). For the calculation of spectra, the atomic velocities were collected every 1.45 fs during the simulations and a correlation time of 0.5 ps was used, giving a resolution of ~20 cm⁻¹.

Because the deuterium mass was used for H in simulations, the calculated VDOS (Figure 6) values were scaled according to $v_{\rm H} = 1.4 \times v_{\rm D}$. Three main bands are present in each profile, consistent with other simulation reports (Boek and Sprik, 2003; Sainz-Diaz *et*

al., 2004; Larentzos et al., 2007). Based on comparisons with previous studies, the band at $\sim 3500 \text{ cm}^{-1}$ can be assigned unambiguously to the OH-stretching mode. It also coincides well with the experimental range from 3400 to 3600 cm⁻¹ (Farmer, 1974; Bishop *et al.*, 2002; Petit, 2006). Botella et al. (2004) reported that Fe(III) substitution can shift the stretching to the lowerfrequency region, which is not detected on the calculated spectrum. However, the presence of Fe(II) yields a lesser stretching frequency than Fe(III), which is consistent with experiments (Manceau et al., 2000; Fialips et al., 2002). In both the oxidized and reduced smectites, the spectra of H_{A1} showed the in-plane bending mode of Al₂-OH at ~930 cm^{-1} , which agrees well with experimental measurements (Farmer, 1974; Bishop et al., 2002). Redshifts of the in-plane bending modes can be observed in both H_{Fe} spectra: ~890 cm⁻¹ if Fe(III) and 820 cm⁻¹ if Fe(II). The reducing trend is consistent with the previous



Figure 7. VDOS of silicon in the (a) oxidized and (b) reduced models.



Figure 8. Power spectra of Si-O bonds in the (a) oxidized and (b) reduced models.

calculation (Botella *et al.*, 2004) and experiments (Goodman *et al.*, 1976). Because one Fe atom is present in each model, *i.e.* only two (AlFe)-OH hydroxyls, the red-shifts have not been reflected in the spectra of all hydrogen (black lines in Figure 6). As for the bands below 700 cm⁻¹, the low-frequency modes such as the out-of-plane bending of OHs and some libration modes are included (Farmer, 1974; Bishop *et al.*, 2002).

In the vibrational spectra of modes involving Si (Figure 7), the following domains can be distinguished: ~900-1100 cm⁻¹, 700-900 cm⁻¹, 650 cm⁻¹, and below 400 cm⁻¹. In the spectra of $Si_{Fe(II)}$ (Figure 7b), the 700-900 cm⁻¹ domain splits into two peaks.

In order to assign the Si–O stretching modes, the power spectra of Si–O bonds (Figure 8) were calculated by Fourier transforming the auto-correlation functions of Si–O bond lengths (Gaigeot and Sprik, 2003). In one Si–O tetrahedron, the three Si–O_{basal} interactions were very similar, but they were different from the Si–O_{apical} interaction. By comparing Figure 7 with Figure 8, the bands at ~900–1100 cm⁻¹, ~700–900 cm⁻¹, and

650 cm⁻¹ are all attributed to the Si–O stretching modes, consistent with previous reports (Bougeard *et al.*, 2000). The Si–O_{basal} stretching contributes to all three bands, whereas the Si–O_{apical} stretching contributes mainly to the latter two bands. In Figure 7b, an additional peak at ~900 cm⁻¹ can be observed on the Si–O_{apical} spectrum, which corresponds to the peak in the spectrum of Si_{Fe(II)} (Figure 7b).

The analyses above can also be confirmed by the component VDOS. The anisotropic VDOS contributions (XX, YY, and ZZ; Figure 9) are calculated for each Si type using only the relevant \mathbf{x} , \mathbf{y} , and \mathbf{z} projections of the velocity vectors. In the present systems, the \mathbf{x} and \mathbf{y} components of velocities are within the plane, parallel to the basal surface, and the \mathbf{z} component is perpendicular to the basal surface. Therefore, XX and YY reflect the correlations of motions within the plane parallel to the basal surface and ZZ reflects that perpendicular to the basal surface. The Si $-O_{\text{basal}}$ vector is almost parallel to the basal plane, so its stretching contribution is mainly on XX and YY. As can be seen in Figure 9, all XX and



Figure 9. Component VDOS of silicon in the (a) oxidized and (b) reduced models.

YY curves contribute to the bands at 800 cm⁻¹ and 1000 cm⁻¹, which thus correspond to the contribution of Si-O_{basal} stretching. Because the Si-O_{apical} vector is approximately perpendicular to the basal surface, its stretching mode contributes mainly to the out-of-plane vibration, *i.e.* on ZZ. The spectra show that the ZZ contribution to the band at 1000 cm⁻¹ is trivial. In Figure 9b, the ZZ of Si_{Fe(II)} is shifted to higher frequency, which corresponds to the peak located at ~900 cm⁻¹ (Figures 7b, 8b). A previous IR spectroscopy study of Fe-containing smectites (Yan and Stucki, 1999) indicated that Fe reduction can increase the out-of-plane Si-O stretching frequency, which is qualitatively consistent with the finding in the Si_{Fe(II)} spectrum.

SUMMARY

In the present study, the structural and vibrational properties of Fe-bearing smectites were investigated in detail using *ab initio* molecular dynamics. The average Fe-O bond length was found to be ~2.08 Å in the reduced smectite and ~2.02 Å in the oxidized model. Neither substitution imposed significant influence on the coordination structures of the Al-O and Si-O polyhedra. For hydroxyl orientations, Fe(III) substitution had no obvious influence whereas Fe(II) forced the coordinated hydroxyls to present a broader distribution. Both substitutions were found to shift the in-plane bending of hydroxyls downward and the Fe(II) substitution can redshift the stretching frequencies. The Si-O stretching was not influenced by Fe(III) substitution, but the outof-plane stretching mode can be shifted upwards by Fe reduction.

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