DEHYDROXYLATION BEHAVIOR OF HEAT-TREATED AND STEAM-TREATED HOMOIONIC CIS-VACANT MONTMORILLONITES

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Abstract—Li⁺, Na⁺, Ca²⁺, Sr²⁺, Cu²⁺, or Zn²⁺-saturated samples of a *cis*-vacant montmorillonite from Linden, Bavaria, were heated to temperatures between 200–700°C. Half of each heated sample was subsequently autoclaved under steam at 200°C (\sim 1.5 MPa) to promote rehydroxylation. The smectites were characterized by cation-exchange capacity (CEC), determination of exchangeable cations, infrared (IR) spectroscopy, and thermoanalytical investigations of evolved water in a thermobalance linked with a mass spectrometer.

Changes in the montmorillonite structure and dehydroxylation behavior are related to three respective mechanisms: type of the interlayer cation, interlayer cation radius, and the movement of the interlayer cation. The migration of the smaller Li⁺, Cu²⁺, and Zn²⁺ ions after heating produces a strong reduction of CEC due to the Hofmann-Klemen effect before the initiation of dehydroxylation. Thereafter, the CEC of these smectites remains constant over a large temperature interval during dehydroxylation. After rehydroxylation, Cu²⁺ and Zn²⁺-rich samples release 16–23 meq/100 g of Mg²⁺ from the structure. No Mg²⁺ release is observed for the Li⁺-rich montmorillonite. Also the dehydroxylation behavior after rehydroxylation differs between the Cu²⁺, Zn²⁺, and Li⁺-rich samples. The mass curves of the evolved water during thermoanalysis of the rehydroxylated Cu²⁺ and Zn²⁺-rich smectites show a peak doublet between 480–700°C. For the Li⁺, Na⁺, Ca²⁺, and Sr²⁺-rich montmorillonites, the second peak disappeared and a third peak at ~760°C developed after rehydroxylation. The resulting structure after rehydroxylation of all samples is celadonite-like.

Key Words—Celadonite, Cis-Vacant Octahedra, Dehydroxylation, Homoionic Montmorillonite, Mass Spectrometry, Rehydroxylation, Thermogravimetry, Trans-Vacant Octahedra.

INTRODUCTION

Smectites and other dioctahedral 2:1 layer silicates lose water in three temperature ranges during heating: 1) dehydration at temperatures <220°C involves the loss of H₂O from the interlayer of expandable clays, 2) OH groups bound to the surface are released between 220–350°C, and 3) dehydroxylation between 350–1000°C results from the reaction between the octahedral hydroxyls according to 2(OH) \rightarrow O_r + H₂O↑ (Brindley and Lemaitre, 1987; Köster, 1993; Drits *et al.*, 1995), when r means residual and the oxygen atom remains in the structure.

In general, montmorillonites are characterized by dehydroxylation temperatures between 500–800°C. Three montmorillonites may be distinguished by dehydroxylation. Two types of montmorillonite show one endothermic peak in differential thermal (DT) curves at ~550 or ~700°C, whereas the third group shows a peak doublet between 500–700°C (*e.g.*, Grim and Kulbicki, 1961; Heller *et al.*, 1962; Schultz, 1969; Koster van Groos and Guggenheim, 1987, 1990; Guggenheim, 1990; Drits *et al.*, 1995). Furthermore, dehydroxylated samples regain many of the hydroxyl groups by treatment under steam between 200–300°C (*e.g.*, Grim and Bradley, 1948; Jonas, 1954; Heller *et al.*, 1962; Hamilton, 1971; Alvero *et al.*, 1994). After rehydroxylation the montmorillonites show one dehydroxylation peak near 600°C or two peaks near 500 and 650°C. Based on all investigations, the rehydroxylated montmorillonites apparently lose hydroxyls at lower temperatures than the original samples.

According to Tsipursky and Drits (1984), octahedral cations of montmorillonites are distributed over the trans- and cis-sites for dioctahedral smectites, *i.e.*, there exist trans-vacant (tv) and cis-vacant (cv) sites. Drits et al. (1995) developed a model, relating the dehydroxylation temperature to the position of the Al³⁺ cations and their movement to other sites in the octahedral sheet. Heller-Kallai and Rozenson (1980) found, that Fe³⁺ do not migrate. Smectites consisting of cv 2:1 layers are characterized by dehydroxylation temperatures of 150-200°C higher than those for a similar mineral consisting of tv 2:1 layers. During dehydroxylation, Al³⁺ ions in the cv smectite move from trans-octahedral sites to cis-octahedral sites. This requires additional thermal energy and explains the higher dehydroxylation temperature of cv smectites. Drits et al. (1995) assumed that during heating the transformation occurs in a portion of the sample, because two dehydroxylation peaks are observed after rehydroxylation of preheated montmorillonites. For this interpretation, the heating conditions are important, because the loss of hydroxyl groups starts at lower temperatures with a slower heating rate (Hamilton, 1971).

Table 1. Variation of the CEC and exchangeable interlayer cations.

	meq/100 g dry clay								
Homoionic sample	CEC _{theor}	CEC _{mcasur}	Main interlayer cation	Σ exchangeable cations					
LI-Li	101	101	90	99					
LI-Na	99	103	100	103					
LI-Ca	99	101	98	102					
LI-Sr	97	100	101	103					
LI-Cu	98	99	96	101					
LI-Zn	98	99	98	102					

Drits *et al.* (1995) further noted that structural changes and Al^{3+} movement in cv montmorillonites during dehydroxylation are not restricted by the interlayer cations. The movement of a small Li⁺, Cu²⁺, or Zn²⁺ cation from the interlayer to the hexagonal cavities of the tetrahedral sheet and then to the vacant octahedral positions of dioctahedral clays upon heating is a well known phenomenon. (*e.g.*, Hofmann and Klemen, 1950; Greene-Kelly, 1955; Calvet and Prost, 1971; McBride and Mortland, 1974; Heller-Kallai and Moser, 1995).

The goal of this study is to test the hypothesis that dehydroxylation behavior indicates changes in the clay structure developed after dehydroxylation and rehydroxylation depending on the type of the interlayer cations.

MATERIAL

White to yellowish montmorillonite from Linden, Bavaria, $Ca_{0.10}Mg_{0.08}(Na,K)_{0.01}[(Si_{3.95}Al_{0.05})(Al_{1.46}Fe_{0.18})]$ $Mg_{0.38}O_{10}(OH)_2$] (Molecular weight: 371.3 g mol⁻¹, cation-exchange capacity (CEC) of 100 meq/100 g) was used in this study. Free iron oxides were removed using the procedure of Mehra and Jackson (1960); 8.7 meq iron per 100 g dried (105°C) material and 17.1 meg aluminum were removed. Excess citrate was removed by dialysis. The $<2-\mu m$ fraction was obtained by sedimentation from ~ 2 wt. % suspension. Homoionic samples of Li⁺, Na⁺, Ca²⁺, Sr²⁺, Cu²⁺, and Zn²⁺ were prepared by saturating ~ 15 g of clay for 12 h once with 700 mL of a 0.2 M chloride solution and three times with a 0.1 M chloride solution. The clay was then dialyzed using distilled water until chloride free. The suspended samples were dried at 55°C in polyethylene bottles, gently crushed in a mortar, and stored at room temperature. CEC measurement and determination of exchangeable cations showed an exchange of 90-99% relative to the prepared starting materials (Table 1). The different interlayer cations produce a change of molecular weight and small changes in CEC. The clay shows one dehydroxylation reaction at ~670°C and was classified as a cis-vacant dioctahedral smectite (Drits et al., 1995).

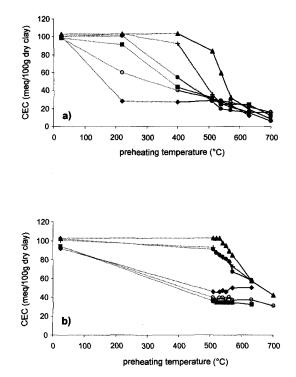


Figure 1. The evolution of the CEC (a) of the heated homoionic montmorillonites (- \bullet -LI-Li, -- LI-Na, -- LI-Ca, -- LI-Sr, -- -LI-Cu, -- LI-Zn) and (b) of the rehydroxylated homoionic montmorillonites (- \bullet -LI-Li rehydr., -- LI-Na rehydr., -- LI-Ca rehydr., -- LI-Sr rehydr., -- LI-Cu rehydr., -- LI-Zn rehydr.).

METHODS

Samples of 1 g of the homoionic clays were heated in a programmable furnace with a heating rate of 150° C h⁻¹ to 200, 400, and to temperatures between $510-570^{\circ}$ C in intervals of 10°C, and also to 630 and 700°C, respectively. The final temperature was maintained for 20 h. Half of each sample was immediately rehydroxylated under steam at 200°C (~1.5 MPa) for 24 h in a stainless steel autoclave. Selected samples were also steamed at several temperatures between 120-200°C without preheating.

CEC and exchangeable cations were determined from 60 mg of each sample dried at 105°C. Cations were removed by washing with 25 mL of a 1 M ammonium acetate solution (pH = 7) for 12 h. The suspension was centrifuged and the supernatant was replaced by 25 mL of fresh 1 M ammonium acetate solution, and the sample was redispersed and centrifuged. After repeating the procedure four times, the clay was washed with ethanol six times by using the centrifuge (Mackenzie, 1951). All supernatant liquids were collected in porcelain bowls, evaporated, and prepared for atomic absorption spectroscopy (AAS) to determine exchangeable cations. For the CEC determination, the NH₄ was exchanged by Na⁺ with 2 mL 50 wt. % NaOH solution. Liberated NH_3 was transferred by steam distillation in a Büchi 323 distillation Unit into an Erlenmeyer flask and titrimetrically determined with 0.0067 N H_2SO_4 . Potentiometric titrations were performed to determine exchangeable H⁺. 20 mL of a 0.5% clay suspension were titrated with 0.01 M KOH in steps of 0.2 mL. The pH was measured every 3 min.

Thermoanalytical investigations were made in a Mettler thermobalance linked to a Balzers quadrupole mass spectrometer (MS) as described in Kahr *et al.* (1996). The combination makes it possible to register simultaneously selected masses of the evolved gases during thermal reactions in the thermobalance. Samples of 100 mg were investigated with a heating rate of 10 K min⁻¹ in streaming air (3 L h⁻¹) in the range from 30 to 1000°C. MS curves of evolved water (m/e = mass divided by charge 18) were used to determine the dehydroxylation temperature.

Clay samples of 100 mg were dried at 105° C. They were decomposed in a KOH melt, dissolved in 20 mL 1:1 HCl and adjusted to 100 mL with Aqua bidest. This solution was analyzed by AAS, excluding K⁺ and Na⁺, which were determined in the CEC procedure. The structural formulae were computed after van Olphen (1963).

Two sample preparation methods were used for Xray diffraction (XRD) analysis. 1) For oriented samples, 15 mg clay were ultrasonically dispersed in 1.5 mL distilled water and the suspension was pipetted onto a special glass slide and air-dried. 2) The samples were mixed with 10 wt. % LiF as internal standard for random oriented powder samples. All samples were step scanned (0.02 ° θ /2 s) from 1.5 to 65 2 θ ° in a Philips APD 1900 X-ray diffractometer with CuK α radiation.

Infrared (IR) transmission spectra were obtained on a Fourier transform infrared (FTIR) spectrometer Perkin Elmer System 2000 using KBr pellets (2 mg sample diluted in 198 mg dry KBr). For every sample, 24 scans were recorded between 4000–370 cm⁻¹. The pellets were measured again after being heated for 16 h at 150°C (Russell and Fraser, 1994) and 200°C, respectively, to remove adsorbed water.

RESULTS

Changes of CEC and exchangeable cations

The effects of heating and rehydroxylation on CEC and exchangeable cations are shown in Table 2a and 2b. A good agreement exists between the CEC and the sum of exchangeable cations for the unheated homoionic montmorillonites (Table 2a). Difference of 1 or 2 meq/100 g dry material are within error.

Thermal treatment of the homoionic montmorillonites reduced the CEC with increasing temperature (Table 2a). The clays with Na^+ , Ca^{2+} , or Sr^{2+} as interlayer cations (>0.97 Å ionic radius) showed no, or just a slight, CEC reduction after heating to 220°C. The Ca²⁺ and Sr²⁺ remain fully exchangeable by NH₄⁺, but 13 meq Na⁺ did not. Further heating caused a weak and continuous CEC decrease until ~400°C for the Ca²⁺ and Sr²⁺-rich clay and until 510°C for the Na⁺-rich clay, followed by a sharp drop in CEC between 510–630°C. Heating at 700°C caused only a slight further reduction in CEC (Figure 1a).

For montmorillonites with smaller interlayer cations, different CEC patterns were observed. The CEC of the Li⁺-rich clay reached a minimum at 28 meq/ 100 g after heating to 220°C and no significant changes of the CEC occurred for Li⁺-rich samples heated at temperatures to 570°C (Figure 1a). Note that the exchangeable Li⁺ increased from 11 to 26 meq/100 g after preheating at temperatures >510°C. Heating at 630°C entailed a further CEC reduction to 12 meq/100 g with still 25 meq/100 g exchangeable Li⁺.

The CEC of the Cu²⁺-rich clay decreased to 60 meq/ 100 g, after the sample was heated at 220°C and still 20 meq/100 g Cu²⁺ remained exchangeable. Further heating at higher temperatures caused a CEC decrease to 16 meq/100 g at 700°C, whereas the amount of exchangeable Cu²⁺ ions decreased to 2–4 meq/100 g above 400°C. Acid conditions equal to 16 meq/100 g exchangeable protons were found for the Cu²⁺-rich montmorillonite preheated at 540°C from the first derivative of the potentiometric titration curves.

The behavior of Zn^{2+} -rich clay is quite similar to that of the Cu²⁺-rich clay. The CEC underwent only a small reduction until 220°C and 74 meq/100 g Zn²⁺ remained exchangeable. The exchangeable Zn²⁺ reached a minimum of 6 meq/100 g at 400°C, whereas the CEC decreased slowly to a minimum of 12 meq/ 100 g during further heating to 700°C (Figure 1a).

The final CEC after rehydroxylation also depends on the nature of the interlayer cation (Table 2b). For the Na⁺-rich clay the CEC remained at 103 meq/100 g until preheating at 530°C (Figure 1b), whereas the amount of exchangeable Na⁺ decreased to 83 meq/100 g (data not shown). After preheating at 540°C and higher temperatures, the CEC and the exchangeable cations decreased slowly. The reduction in the Ca²⁺and Sr²⁺-rich clays was similar with a slightly greater CEC reduction from 101 to 58 meq/100 g and from 100 to 57 meq/100 g, respectively (Figure 1b).

Steaming at 200°C did not effect the CEC of the unheated Na⁺⁻, Ca²⁺⁻, and Sr²⁺-rich clays, as reported also by Oscarson and Dixon (1989), but only 93–96% of these ions remained exchangeable.

The rehydroxylated Cu²⁺-, Zn²⁺-, and Li⁺-rich clays reached a steady CEC at much lower preheating temperatures. After preheating to 510°C and higher, the rehydroxylated Cu²⁺-rich clays showed an almost constant CEC of 39 meq/100 g and the Zn²⁺-rich clays of 36 meq/100 g, which in both cases is 10 meq/100 g

Table 2a. Measured CEC and exchangeable cations [meq/100 g (105°C dried material)] of the preheated samples.

								Homoic	onic samp	le						
Hanting		L	l-Li		LI-Na				I	J-Ca		LI-Sr				
Heating T(°C)	Li+	Mg ²⁺	Σı	CEC	Na ⁺	Mg ²⁺	Σı	CEC	Ca ²⁺	Mg ²⁺	Σı	CEC	Sr ²⁺	Mg ²⁺	Σ^{1}	CEC
RT	90	7	99	101	100	2	103	103	98	3	102	101	101	1	103	100
220	11	4	16	28	87	3	92	103	99	3	103	102	99	3	104	103
400	11	3	15	27	85	3	90	104	53	2	56	55	90	3	95	93
510	26	2	29	29	73	2	76	84	20	2	23	28	24	2	28	36
540	27	1	29	28	56	1	59	60	15	1	17	20	17	2	21	26
570	25	1	27	24	31	2	34	33	11	2	14	18	13	2	17	22
630	25	1	27	12	19	2	22	19	7	2	10	15	10	1	13	16
700	3	1	5	5	8	1	10	8	3	1	5	6	4	2	8	16

¹ Including small amounts of exchangeable Al³⁺, Fe³⁺, K⁺, Ca²⁺, Na⁺.

higher than the CEC of samples not rehydroxylated. The Li⁺-rich clay regained a CEC averaging 48 meq/ 100 g, 20 meq higher than those clays that were only preheated. For the Cu²⁺- and Zn²⁺-rich clays a remarkably high amount of exchangeable Mg²⁺, between 16–23 meq/100 g, was measured. Steaming at temperatures between 120–200°C of nonheated Cu²⁺- and Zn²⁺-rich montmorillonites also caused nearly the same amount of exchangeable Mg²⁺ in interlayer positions, increasing with an increased steam temperature (Emmerich and Kahr, 1998). In contrast no exchangeable Mg²⁺ were detected for steam treated Li⁺-rich samples.

Thermal investigations and dehydroxylation behavior

Preheated samples. The natural Ca^{2+} and Mg^{2+} -rich montmorillonite from Linden and its homoionic forms have one dehydroxylation reaction between 625– 685°C (Figure 2). All homoionic samples lost 4.85 ± 0.1% of the initial weight of the dry clay as the result of dehydroxylation. This is in perfect agreement with the theoretical OH content of the dioctahedral smectite. After preheating the homoionic clays, the main dehydroxylation peak increased by ~30°C and lost intensity. A second peak near 450–480°C appeared (Figures 3 and 4).

Heating to 220°C did not apparently effect the number of OH groups, but heating to 400°C caused a small reduction of the mass of evolved water corresponding to octahedral hydroxyl groups. Further heating produced a continuous loss of octahedral OH groups with no appreciable differences due to the saturating cation. A completely dehydroxylated state was not attained even after heating to 700°C, because the dehydroxylated montmorillonites show a spontaneous rehydroxylation in the laboratory atmosphere (K. Emmerich, unpubl. data, 1998). The clays regained nearly 10% of their initial hydroxyls, which are possibly bound to the edges of the dehydroxylated layers (Figure 5).

Rehydroxylated samples. After heating and rehydroxylation of Cu²⁺-and Zn²⁺-rich clays, the main MS peak increased by $\sim 25^{\circ}$ C and lost intensity with increasing temperature. A second peak appeared near 475–495°C and increased in intensity at higher temperature. A weak shoulder at 860°C maintained position and intensity (Figures 4 and 6).

For the Na⁺-, Ca²⁺-, and Sr²⁺-rich clays a different phenomenon was observed. After preheating until 520° C and rehydroxylation, the Na⁺-, Ca²⁺-, and Sr²⁺rich clays and the Li⁺-rich clay, respectively, behaved similarly to the clays described above (curves not shown). After preheating to higher temperatures, however, the peak between 660–680°C decreased in intensity and a peak between 750–790°C became more intense than the peak at 480–495°C (Figures 3 and 6).

The water content corresponding to hydroxyl groups did not reach the initial level upon rehydroxylation

Table 2b. Measured CEC and exchangeable cations [meq/100 g ($105^{\circ}C$ dried material)] of the rehydroxylated samples (steam temperature of $200^{\circ}C$).

								Homoio	mic samp	ole						
n . 1 . 4	LI-Li			LI-Na			LI-Ca			LI-Sr						
Pre-heating T (°C)	Li+	Mg ²⁺	Σ1	CEC	Na+	Mg ²⁺	Σ^1	CEC	Ca ²⁺	Mg ²⁺	Σ1	CEC	Sr ²⁺	Mg ²⁺	Σ^1	CEC
RT	81	6	89	92	96	2	100	103	93	3	97	102	96	4	102	101
510	42	3	47	46	90	1	93	103	88	2	91	91	87	3	92	93
540	41	2	45	47	84	2	89	98	65	2	68	83	74	4	82	82
570	41	2	45	50	79	1	81	85	57	1	59	67	68	3	72	73
630	41	2	45	50	52	2	54	57	50	2	53	58	50	4	56	57
700				_	39	2	41	42					_	_		

¹ Including small amounts of exchangeable Al³⁺, Fe³⁺, K⁺, Ca²⁺, Na⁺.

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	Homoionic sample											
	LI	-Cu			L	-Zn						
Cu ²⁺	Mg ²⁺	Σ١	CEC	Zn ²⁺	Mg ²⁺	Σ1	CEC					
96	4	101	99	98	3	102	99					
20	4	26	60	74	5	81	91					
4	3	9	40	6	4	12	44					
2	4	8	31	7	4	13	32					
2	4	8	29	6	4	12	25					
2	4	8	27	6	4	12	25					
3	5	10	22	7	4	13	24					
3	5	10	16	4	4	10	12					

Table 2a. Extended.

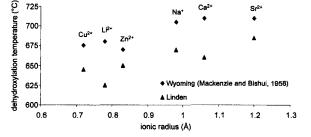


Figure 2. Dehydroxylation temperatures vs. the ionic ratio of the interlayer cations.

under steam at 200°C. Doubling the reaction time to 48 h did not increase the water uptake. Whereas the heated montmorillonites saturated with Na⁺, Ca²⁺, and Sr²⁺ regained 75–95% of their OH groups, the Li⁺-, Cu^{2+} , and Zn^{2+} -rich montmorillonites regained only 60-80% of their OH groups, although the number of liberated OH during preheating did not show large differences (Figure 5).

IR spectroscopy

OH-stretching region. The absorption band at 3630 cm⁻¹, which is assigned to stretching vibrations of octahedral OH groups, was found to be independent of interlayer-cation species for the unheated homoionic montmorillonites. Also heating between 510-700°C caused similar shifts to higher wavenumbers (~3675 cm⁻¹) for all samples. Moreover, a shoulder between 3630-3650 cm⁻¹ appears in spectra of the heated samples (spectra not shown). The absorption strongly diminished with continued dehydroxylation. Rehydroxylation resulted in increased absorption but no shifts of these bands.

Vibrations of molecular water. Vibrations corresponding to adsorbed water at 3440 and 1630 cm⁻¹ diminished with increasing temperature of drying of the KBr pellets. These vibrations nearly disappeared after drying at 200°C.

OH-bending and M-O vibration regions. Between 1200-450 cm⁻¹, there are seven major changes, which

Table 2b. Extended.

Homoionic sample											
	LI	·Cu			Ll	Zn					
Cu ²⁺	Mg ²⁺	Σı	CEC	Zn ²⁺	Mg ²⁺	Σ^1	CEC				
66	18	86	94	70	17	- 89	94				
6	16	23	39	10	16	28	36				
4	23	29	39	7	19	28	34				
3	19	23	37	5	19	26	34				
3	22	27	37	4	16	22	32				
3	20	23	31		_		_				

are described below. However, there were no significant differences in the spectra relating to saturating cations after heating the homoionic clays at temperatures between 510-700°C. Unless discussed below, rehydroxylation did not cause any other changes in the bands (Figures 7, 8, and 9). The seven major changes are: 1) The band at 798 cm⁻¹, corresponding to OH associated with MgMg pairs (Calvet and Prost, 1971), apparently moved after heating at 510 or 540°C to wave numbers between 800-818 cm⁻¹ and returned to 799 cm⁻¹ after heating the samples at higher temperatures, as well as after rehydroxylation. Only the rehydroxylated Cu²⁺-rich samples show a band higher than 808 cm⁻¹. 2) The band at 845 cm⁻¹, assigned to OH groups surrounded by AlMg pairs (Calvet and Prost, 1971), nearly disappeared after heating for all samples and can be detected only as a small shoulder. 3) The AlFeOH band at 882 cm⁻¹ shifted to lower wave numbers (to 865 cm⁻¹), and the AlAlOH band at 916 cm⁻¹ shifted to higher wavenumbers (to 942 cm⁻¹). The band near 880 cm⁻¹ returned for rehydroxylated samples, which were preheated at <540°C. 4) The Si-O stretching bands at 1038 and 1115 cm⁻¹ moved to higher wave numbers at 1051 and 1135 cm^{-1} . 5) The absorption band at 733 cm^{-1} increased with heating of the samples at $>510^{\circ}$ C. 6) The peak at 622 cm⁻¹, assigned to Al-O in the octahedral sheet (Madejová et al., 1996), shifted to 640-647 cm⁻¹. 7) The Si-O bending vibrations near 523 cm⁻¹, due to Si-O-Al (Madejová et al., 1996), disappeared or may have moved to 570 cm⁻¹. After rehydroxylation, two broad absorption bands at 523 and 565 cm⁻¹ were detected. The position of the Si-O-Si vibration moved from 465 to 483 cm⁻¹.

XRD

All homoionic montmorillonites show a well-defined (001) reflection with the position of this reflection dependent on the interlayer cation and its hydration state under laboratory conditions at ~24°C and relative humidity at \sim 45–55% (Table 3a). All samples expanded after ethylene glycol treatment to d(001) =16.9–17.1 Å. The d(001) of the Na⁺-, Ca²⁺-, and Sr²⁺rich montmorillonites collapsed to 9.7-9.8 Å after

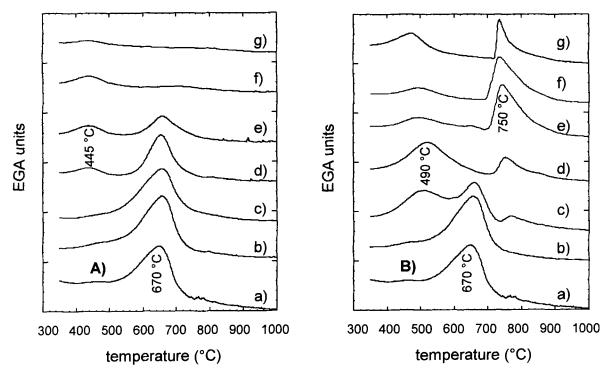


Figure 3. MS curves of the evolved water mass fragment 18 m/e (A) of Na⁺-montmorillonite a) untreated; and preheated b) 220°C, c) 400°C, d) 510°C, e) 540°C, f) 570°C, and g) 630°C and (B) of Na⁺-montmorillonite a) untreated; b) steamed at 200°C; and rehydroxylated after preheating at c) 520°C, d) 530°C, e) 550°C, f) 570°C, and g) 630°C.

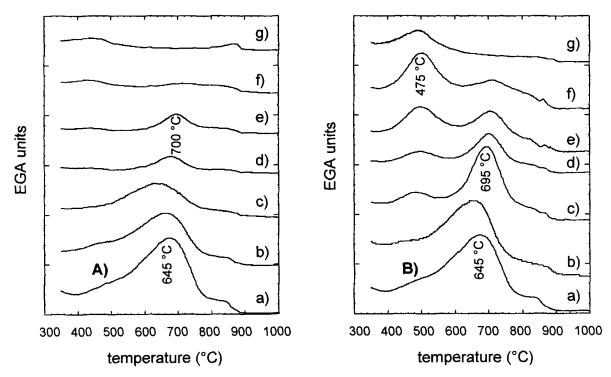


Figure 4. MS curves of the evolved water mass fragment 18 m/e (A) of Cu^{2+} -montmorillonite a) untreated; and preheated b) 220°C, c) 400°C, d) 510°C, e) 540°C, f) 570°C, and g) 630°C and (B) of Cu^{2+} -montmorillonite a) untreated; b) steamed at 200°C; and rehydroxylated after preheating at c) 510°C, d) 530°C, e) 550°C, f) 570°C, and g) 630°C.

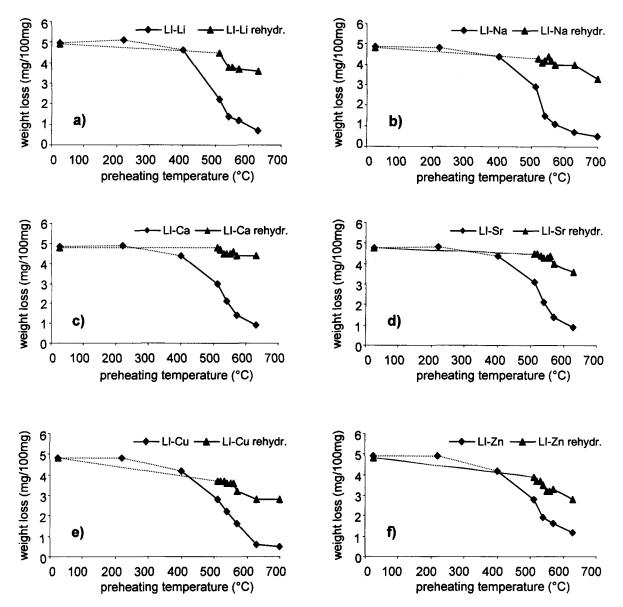


Figure 5. Weight loss caused by evolved water, which corresponds to the hydroxyl groups, during thermal analysis between $350-1000^{\circ}$ C of the heated and rehydroxylated homoionic a) Li⁺-, b) Na⁺-, c) Ca²⁺-, d) Sr²⁺-, e) Cu²⁺-, and f) Zn²⁺-mont-morillonites.

heating at temperatures of $\geq 400^{\circ}$ C (Table 3a), but they expanded partially in water and ethylene glycol after rehydroxylation as shown for selected Na⁺-rich samples in Figure 10a. After heating at 220–700°C the Cu²⁺-, Zn²⁺-, and Li⁺-rich montmorillonites collapsed to 9.6–9.8 Å and did not swell with ethylene glycol. Rehydroxylation did not produce re-expansion (Table 3b). In contrast, the steamed unheated samples were still able to swell, as shown for the Cu²⁺-rich samples in Figure 10b.

The heated and the steam-treated samples show a peak at 3.16–3.22 Å. This peak moved toward 3.22 Å

and increased in intensity relative to the (001) reflection with increasing temperature of heating.

DISCUSSION

The natural Ca²⁺ and Mg²⁺-rich montmorillonite from Linden dehydroxylated near 670°C. Therefore it consists of *cis*-vacant 2:1 layers (Drits *et al.*, 1995). The homoionic forms have one dehydroxylation peak between 625–685°C depending on the interlayer cations in the following order: Li⁺ < Cu²⁺ < Zn²⁺ < Ca²⁺ < Na⁺ < Sr²⁺. These results are in fair agreement with the relation between the dehydroxylation temperature

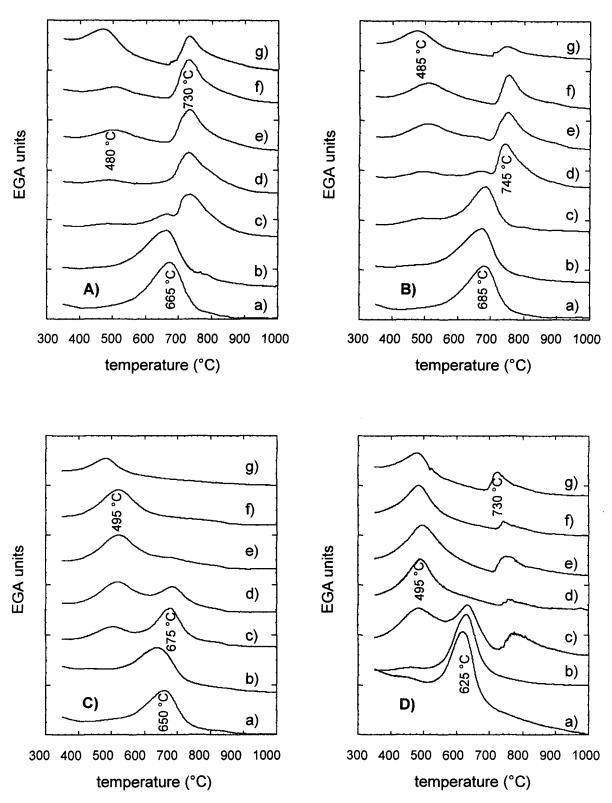


Figure 6. Mass spectrometer curves of the evolved water mass fragment 18 m/e of (A) Ca^{2+} -montmorillonite, (B) Sr^{2+} -montmorillonite, (C) Zn^{2+} -montmorillonite, and (D) Li^{+} -montmorillonite a) untreated; b) steamed at 200°C; and rehydroxylated after preheating at c) 510°C, d) 530°C, e) 550°C, f) 570°C, and g) 630°C.

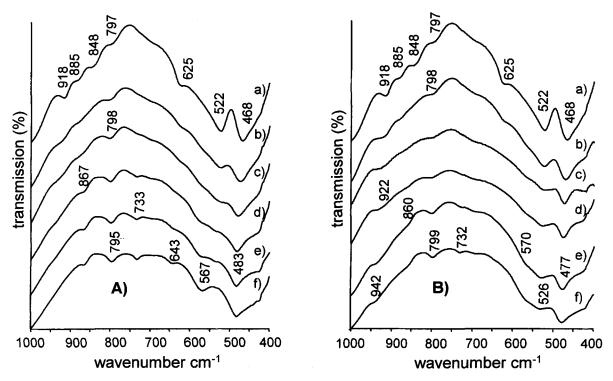


Figure 7. IR spectra (range $1000-400 \text{ cm}^{-1}$) of Na⁺-montmorillonite; (A) preheated and (B) rehydroxylated after preheating at a) RT, b) 510° C, c) 540° C, d) 570° C, e) 630° C, and f) 700° C.

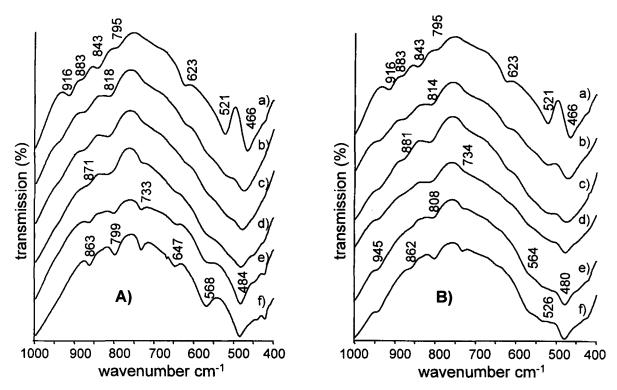


Figure 8. IR spectra (range 1000–400 cm⁻¹) of Cu²⁺-montmorillonite; (A) preheated and (B) rehydroxylated after preheating at a) RT, b) 510° C, c) 540° C, d) 570° C, e) 630° C, and f) 700° C.

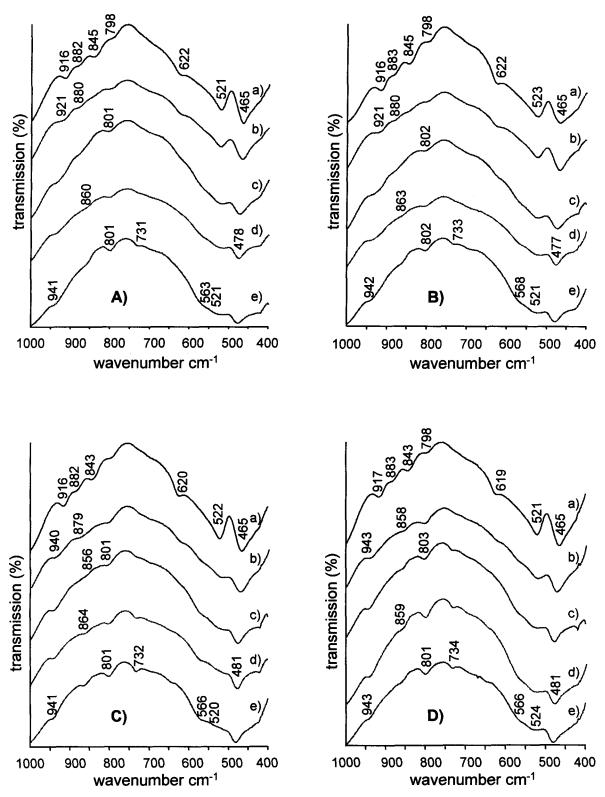


Figure 9. IR spectra (range 1000–400 cm⁻¹) of (A) Ca²⁺-montmorillonite, (B) Sr²⁺-montmorillonite, (C) Zn²⁺-montmorillonite, and (D) Li⁺-montmorillonite a) untreated; and rehydroxylated after preheating at b) 510°C, c) 540°C, d) 570°C, and e) 630°C.

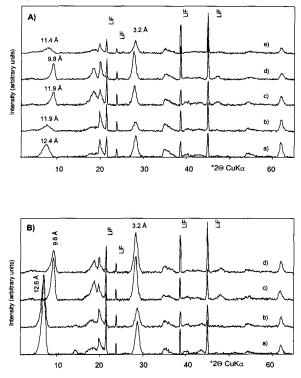


Figure 10. XRD patterns of (A) selected heat and steam treated Na⁺-montmorillonites a) untreated, b) steamed at 200°C, c) heated at 400°C, d) heated at 540°C; and e) rehydroxylated after preheating at 540°C; (B) selected heat and steam treated Cu²⁺-montmorillonites a) untreated, b) steamed at 180°C, c) heated at 400°C, and d) rehydroxylated after preheating at 540°C.

and the Goldschmidt ionic radius (Mackenzie and Bishui, 1958) for ionic radii < 1.2 Å (Figure 2).

Migration kinetics of small ions into the octahedral sheet are important, because this affects the dehydroxylation temperature during thermal analysis. After heating Li⁺-saturated montmorillonite at 180°C and maintaining this temperature for 1 h, Calvet and Prost (1971) found that 73% of the total Li⁺ becomes nonexchangeable and many of the Li⁺ migrated into octahedral sites. Thus, it appears likely that Li⁺, Cu²⁺, and Zn²⁺ migrate into octahedral sites during thermal analysis, and the resulting structure has a lower dehydroxylation temperature than otherwise.

Structural changes below 510°C

After preheating to remove ligand water, small interlayer cations (Li⁺, Cu²⁺, Zn²⁺) migrate into the hexagonal holes of the tetrahedral sheet (McBride and Mortland, 1974) and into octahedral sites (*e.g.*, Brindley and Ertem, 1971; Madejová *et al.*, 1996). This migration caused a CEC reduction to an apparent lower limit of 14–31 meq/100 g greater than the tetrahedral charge. This is consistent with McBride and Mortland (1974).

Table 3a. The d(001)-values of heated samples.

Heating temperature (°C)	<i>d</i> (001)-values (Å)								
	LI-Li	LI-Na	LI-Ca	LI-Sr	LI-Cu	LI-Zn			
25	12.7	12.4	15.2	15.5	12.6	14.9			
220	9.8	10.3	13.8	13.6	9.7	9.8			
400	9.8	9.8	9.7	9.9	9.6	9.6			
510	9.8	9.8	9.7	9.7	9.6	9.6			
540	9.8	9.8	9.7	9.7	9.6	9.7			
570	9.8	9.8	9.7	9.8	9.6	9.7			
630	9.8	9.8	9.7	9.8	9.6	9.7			
700	9.8	9.8	9.8	9.8	9.7	9.7			

Clays with interlayer cations of ionic radius >0.97 Å show either no change in CEC after heating to 220°C or a small reduction upon heating to 400°C. Larger cations (Na⁺, Ca²⁺, and Sr²⁺) are not as mobile owing to size, and they do not move into octahedral sites after complete dehydration (*e.g.*, Calvet and Prost, 1971).

All heated homoionic samples showed a slightly higher dehydroxylation temperature compared to the initial samples. Heating to 400°C caused a small reduction in evolved water, which corresponded to hydroxyl groups of the octahedra. This indicates that some of the hydroxyl groups were lost during preheating prior to the dehydroxylation event, presumably as a result of protonation of the OH groups, as reported for Li⁺ charge-reduced montmorillonites (Ertem, 1972; Heller-Kallai and Rozenson, 1980; Madejová *et al.*, 1996).

Structural changes during dehydroxylation

CEC of Na⁺-, Ca²⁺-, Sr²⁺-rich clays decreased with increased heating temperature. These large ions do not enter octahedral sites. Apparently, the layer charge is reduced as a result of relatively small positional changes and stronger bonding to the hexagonal cavities of the tetrahedral sheets. This results in the observed decrease in basal spacing.

The dehydroxylation reaction $OH_{cis} + OH_{cis} \rightarrow H_2O\uparrow + O_r$ produces five octahedral configurations: 1) five-fold coordinated dehydroxylated polyhedra (former *cis* positions), 2) six-fold distorted *trans* octahedra with one O_r and one OH group, 3) six-fold distorted *trans* octahedra, with two O_r, 4) distorted *cis* octahedra,

Table 3b. The d(001)-values of rehydroxylated samples.

Preheating	d(001)-values (Å)									
temperature (°C)	LI-Li	LI-Na	LI-Ca	LI-Sr	LI-Cu	LI-Zn				
25	12.8	11.9	13.3	12.5	12.7	14.0				
510	9.8	11.7	14.4	12.3	9.6	9.7				
540	9.8	11.4	14.3	13.2	9.6	9.7				
570	9.8	10.7	13.6	11.4	9.6	9.6				
630	9.7	10.2	9.9	10.2	9.7	9.7				

and 5) nearly unaffected *cis* and *trans* positions in the structure.

Guggenheim et al. (1987), Koster van Groos and Guggenheim (1987, 1990) and Guggenheim (1990) showed, using Pauling's rules, that Al-rich pentagonal prisms form during partial dehydroxylation. The Al-OH bond strength increases in the Al-rich octahedra having one or two shared edges with the pentagonal prisms. The OH groups that are bound more strongly with Al³⁺ require additional thermal energy for the dehydroxylation process to proceed. Drits et al. (1995) described a movement of Al3+ from trans positions into vacant cis positions, which requires also additional thermal energy. With respect to observed dehydroxylation, lattice distortion during partial dehydroxylation causes the movement of some Al3+ from trans positions into cis positions because the thermal energy is too low (Drits et al., 1995). In contrast, some Al³⁺ are more strongly bound and cannot move until sufficient thermal energy is provided.

After heating at 540°C, the partially dehydroxylated samples show a dehydroxylation curve typical for cv 2:1 layer or a mixture of cv and tv 2:1 layer with a main peak near 700°C. The observed water release in MS curves between 420-450°C after heating the samples to <630°C is assigned to hydroxyl groups regained by spontaneous rehydroxylation under laboratory atmosphere. Rehydroxylation occurs because the clay structure is under stress by lattice distortions and by cations in the hexagonal holes. No evidence was found in IR spectra for the development of a porous network in the clay as proposed by Villieras et al. (1994) for dehydroxylated chlorites. This supports the single-crystal study of Zhan and Guggenheim (1995). We believe that the observed decrease of the intensity of the IR band at 1635 cm⁻¹ with preheating and the change of its intensity with relative humidity are consistent with the work of Mosser et al. (1997) and Trillo et al. (1992). Thus, this band is assigned to the hydration shell of partly dehydrated interlayer cations in hexagonal holes of the silicate ring.

It appears that the movement of Cu^{2+} , Zn^{2+} , and Li^+ into vacant octahedral positions does not hinder the movement of Al^{3+} parallel to the (001) plane. Thus, the mechanism of dehydroxylation is similar to that described above. Heating produces a continuous loss of octahedral OH groups without appreciable differences between the different clays.

A difference between the CEC and the total amount of exchangeable cations for the Li⁺-, Cu²⁺-, and Zn²⁺rich samples is related to proton liberation as suggested by Russell and Farmer (1964), Ertem (1972), and Madejová *et al.* (1996). This is confirmed by potentiometric titrations in the present study.

Note that the CEC of all heated homoionic montmorillonites converge to the value of the the tetrahedral charge (\sim 13 meq/100 g) before lattice destruction occurs at temperatures >700°C.

Structural changes during rehydroxylation

For rehydroxylated homoionic montmorillonites, the loss of OH groups occurs at two temperatures, ~480 and ~700°C (Cu²⁺-, Zn²⁺-rich clay) and ~480 and ~760°C (Na⁺-, Ca²⁺-, Sr²⁺-, Li⁺-rich clay), respectively. The shift of the second peak to 700 or 760°C does not support the general interpretation that rehydroxylated montmorillonites lose hydroxyls at lower temperatures than the initial material. For example, montmorillonites with one endothermic peak at ~700°C showed a peak doublet at ~500 and 650°C after rehydroxylation (*e.g.*, Grim and Bradley, 1948; Drits *et al.*, 1995).

The temperature shift of the second hydroxyl loss for the rehydroxylated Na⁺-, Ca²⁺-, Sr²⁺-, or Li⁺-rich montmorillonite indicates an interaction between these cations and the regained hydroxyl groups, which causes a stabilization of the OH bond in the octahedral sheet. The observed dehydroxylation peak at 760°C may be regarded as pyrophyllite-like, because pyrophyllite is a 2:1 phyllosilicate with a vacant site in the trans position much like the thermally treated montmorillonite samples (Drits et al., 1995) and some pyrophyllites also have one dehydroxylation peak near 760°C. However, the layer charge differs between the dehydroxylated and rehydroxylated montmorillonite and pyrophyllite. Whereas pyrophyllite layers are essentially neutral having a maximum CEC of 10 meq/ 100 g (Heller et al., 1962), the montmorillonite regained at least four times as much CEC. Considering that Na⁺, Ca²⁺, and Sr²⁺ did not penetrate into vacant octahedra, a tv structure similar to that of celadonite (Tsipursky and Drits, 1984) is more likely than a pyrophyllite-like structure. On the other hand, Li⁺ migrated into vacant octahedral sites and generated a partial trioctahedral layer silicate with reduced permanent layer charge.

Also for the Cu²⁺, Zn²⁺-rich clay, a small shift of the second dehydroxylation peak in MS curves was observed. Cu^{2+} and Zn^{2+} occupied previously empty octahedral sites to produce sauconite-like trioctahedral areas. After steam treatment, the CEC increased and a noteworthy amount of Mg²⁺ occupied exchangeable sites in interlayer positions, whereas the amount of exchangeable Cu²⁺ and Zn²⁺, respectively, did not increase after steam treatment. The Mg2+ therefore, must be displaced from octahedral positions. It is likely that H₃O⁺ ions are generated from hydrolysis caused by the acidity of Cu^{2+} and Zn^{2+} , which remained in the hexagonal holes after heat treatment. Then H₃O⁺ replaced Mg^{2+} with the H_3O^+ ions gaining access to the octahedral positions from the grain edges or through the hexagonal cavities of the tetrahedral sheet (Kaviratna and Pinnavaia, 1994). More detailed information

about this process is expected after further study. In contrast, the CEC of the rehydroxylated Li^+ -rich clay also increased, but no Mg^{2+} was displaced from octahedral positions, and the amount of exchangeable Li^+ increased. This is explained by the different charge and acidity of these ions.

After dispersing the rehydroxylated Cu²⁺-, Zn²⁺-, and Li⁺-rich clays (preheating temperature of 540°C) in a 5 wt. % soda solution and after washing until salt free, the CEC increased to 62 meq/100 g (Cu²⁺rich clay) and 50 meq/100 g (Zn²⁺- and Li⁺-rich clay). From the evolved water during thermoanalytical measurements, the following chemical formulae were calculated: (1) $Na_{0.17}Mg_{0.017}Cu_{0.013}[(Si_{3.94}Al_{0.06})]$ $(Al_{1.42}Fe_{0.18}Mg_{0.33}Cu_{0.18})O_{10}O_{0.25}(OH)_{1.5}];$ (2) $Na_{0.10}$ $Mg_{0.025}Zn_{0.02}[(Si_{3.96}Al_{0.04})(Al_{1.43}Fe_{0.18}Mg_{0.33}Zn_{0.18})O_{10}]$ $O_{0.28}(OH)_{1.44}$]; (3) $Na_{0.17}Mg_{0.003}Li_{0.014}[(Si_{3.92}Al_{0.08})]$ $(Al_{1,46}Fe_{0.18}Mg_{0.38}Li_{0.21})O_{10}O_{0.32}(OH)_{1.46}]$. These formulae provide the evidence of Mg²⁺ displacement by Cu²⁺ and Zn²⁺ and a nearly complete storage of these cations in previously vacant octahedral sites. Also, only some of the Li⁺ moved into octahedral positions and no Mg²⁺ was replaced. The interlayer Li⁺, which migrated to the hexagonal holes of the tetrahedral sheet, became fully exchangeable after steam treatment.

In the IR spectrum, we could not clearly distinguish the influence of the interlayer cations on the OHstretching vibrations at 3675 cm⁻¹ of rehydroxylated 2:1 phyllosilicates. These vibrations are more or less sharper than the vibrations of the starting material, but they are influenced in a similar way by small cations (Li⁺, Cu²⁺, Zn²⁺) which moved into octahedral vacancies (producing trioctahedral areas) and by larger interlayer cations (Na⁺, Ca²⁺, Sr²⁺) which position themselves deeper into the hexagonal holes of the tetrahedral sheet (producing nonswellable mica structures) (Mosser *et al.*, 1997; Trillo *et al.*, 1992; Russel and Fraser, 1994; Wise and Eugster, 1964).

We conclude that: 1) only one-fifth of the vacant octahedral sites are actually occupied by the small Cu^{2+} (Zn²⁺ and Li⁺) cations after rehydroxylation, 2) that a large number of nonexchangeable Cu^{2+} (Zn²⁺ and Li⁺) cations is located in the hexagonal holes of the 2:1 layers, 3) that the dehydroxylation peak at 760°C disappeared after dispersing the clays by soda or ultrasonic treatment, whereas the amount of hydroxyls stayed constant for the rehydroxylated Li⁺-, Na^{+} , Ca^{2+} , and Sr^{2+} -rich montmorillonite, and 4) that the structural changes detected in IR spectra and XRD patterns are similar for all samples preheated at temperatures >510°C and rehydroxylated. Thus, the occurrence of the proposed celadonite-like structure may be assumed for all rehydroxylated homoionic montmorillonites used in this study.

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