

Correlation of calcareous nannofossil zones to the local first occurrence of *Pachydiscus neubergicus* (von Hauer, 1858) (Ammonoidea) in European Upper Cretaceous sections

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Abstract

The first occurrence (FO) of the ammonite *Pachydiscus neubergicus* (von Hauer, 1858) has been correlated to calcareous nannofossil zonations in several European sections along the northern margin of the Tethyan palaeobiogeographic realm. Both the proposed stratotype section of Tercis (SW France) and complete, ammonite-bearing sections in northern Spain document the FO of *P. neubergicus* within standard nannofossil zone CC23a (UC16), below the LO of *Broinsonia parca constricta*. Other sections such as the type locality Neuberg (Austria), Nagoriani (the Ukraine) and Bjala (Bulgaria) indicate considerable diachroneity of local FOs and show *P. neubergicus* to range up to nannofossil zone CC25b/c (UC20; Late Maastrichtian).

Keywords: Ammonoidea, calcareous nannoplankton, Maastrichtian, Europe, correlation

Introduction

The incomplete definition of Cretaceous stages by stratotype sections has been discussed in previous decades, leading to proposals for the designation of boundary sections during the Brussels meeting on Cretaceous stage boundaries in 1995 (Rawson et al., 1996). For the lower boundary of the Maastrichtian Stage, the Tercis section (Landes, SW France) was proposed for the definition of a Global Boundary Stratotype Section and Point, since this displayed a more continuous section than the type section of the Maastrichtian in the Maastricht area (the Netherlands). The first occurrence (FO) of the ammonite *Pachydiscus neubergicus* (von Hauer, 1858) was chosen as the biostratigraphic boundary criterion (Odin, 1996, 2001).

The present paper provides an overview concerning the correlations of calcareous nannofossil zones with the FO of *P. neubergicus* in several European sec-

tions along the transition from the Tethyan into the Temperate realm, e.g. in northern Spain, Austria and Bulgaria (Fig. 1). We demonstrate that *P. neubergicus* has a rather long stratigraphical range, from the base of the Maastrichtian up to the 'late' Maastrichtian in nannofossil terms (e.g., Burnett, 1998), and that the local first appearance of this ammonite is diachronous.

Definition of the Campanian-Maastrichtian boundary and nannofossil zonations

A number of macrofossil and microfossil events occur around the Campanian-Maastrichtian boundary interval (see Birkelund et al., 1984; Hancock et al., 1992; Odin, 1996). During the Brussels meeting on Cretaceous stage boundaries (1995), the FO of *P. neubergicus* was proposed as the boundary criterion (Odin, 1996), based on the proposed stratotype section at the Tercis quarry, which was chosen on ac-

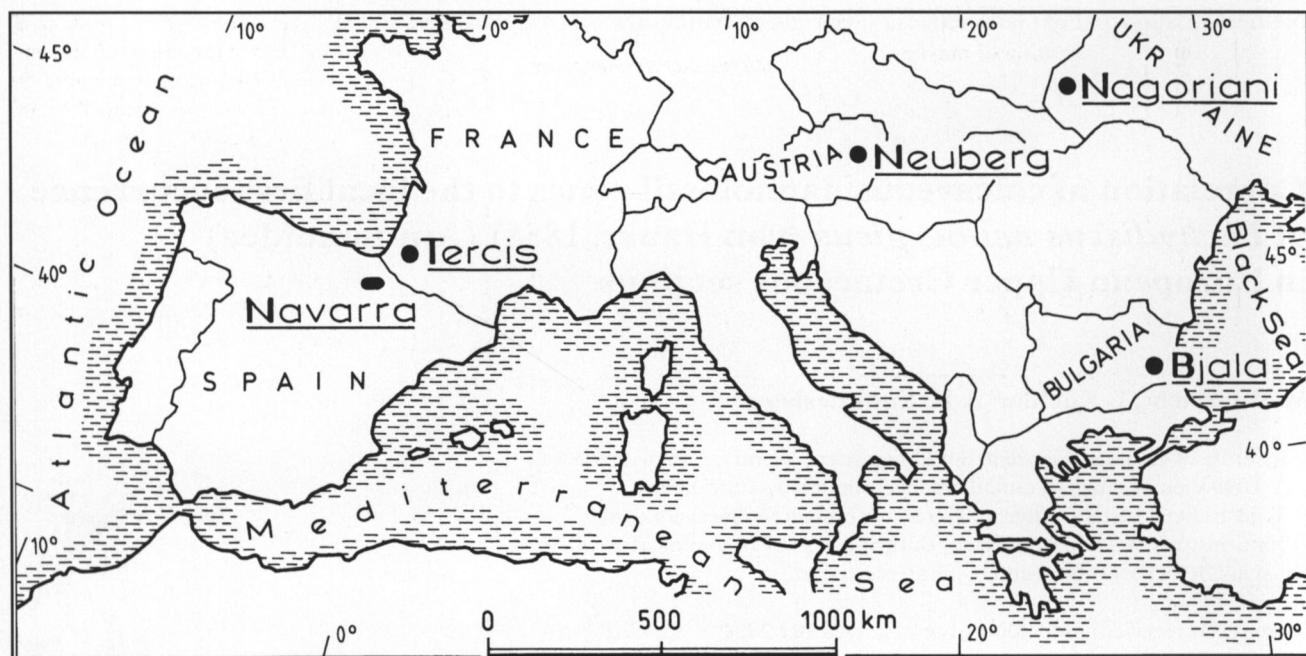


Fig. 1. Sketch map indicating the positions of European sections mentioned in the text.

count of its rich macrofossil content and its transitional position from the Tethys to the temperate/boreal biogeographic realms. The type locality for *P. neubergicus* is in Austria, at Krampen near Neuberg (Steiermark). The Neuberg section was described by Kennedy & Summesberger (1986); that paper also included data on nannofossil assemblages by Wagreich.

Published correlations between the FO of *P. neubergicus* and nannofossil zones either indicate a level within nannofossil zone CC23a/UC16 (e.g., Hancock et al., 1993; Barchi et al., 1997; Wagreich et al., 1998), below the last occurrence (LO) of *Broinsonia parca constricta*, or a level within CC23b/UC17 (e.g., Burnett et al., 1992; Cunha et al., 1997; Burnett, 1998), above the LO of *B. p. constricta*, and below the LO of *Uniplanarius trifidus* and *Tranolithus orionatus*. The sequence of FOs and LOs of primary nannofossil marker species during this time interval is well established both for the Tethyan and the temperate/boreal realms (Perch-Nielsen, 1985; Burnett, 1998), but secondary markers are still under debate. In the latest Campanian and Early Maastrichtian, nannofossil marker events for standard zonations of low to mid-latitudes (Sissingh, 1977; Perch-Nielsen, 1985; Burnett, 1998) are almost exclusively based on the last occurrences of a number of species, namely *Eiffelithus eximius*, *Reinhardtites anthophorus*, *Tranolithus orionatus*, *Broinsonia p. parca* (= *Aspidolithus p. parcus* of Perch-Nielsen, 1985), *Broinsonia p. constricta* (= *A. p. constrictus* of Perch-Nielsen, 1985), and *Reinhardtites levis* (Fig. 2). These LOs of marker species are prone to reworking, which therefore may intro-

duce a significant error to age interpretations of nannofossil assemblages around this time interval, especially in turbiditic deep-water sections (e.g., Wagreich & Krenmayr, 1993). Some of the secondary nannofossil events of Perch-Nielsen (1985) such as the LOs of *Uniplanarius trifidus* and *Uniplanarius sissinghii* may be truly diachronous from one palaeogeographic domain to another, as it is known for the FO of *Nephrolithus frequens* in the Maastrichtian (e.g., Burnett, 1998).

Considering these secondary markers the sequence of nannofossil events around the Campanian-Maastrichtian boundary as described for the Bottaccione section (Italy; see Monechi & Thierstein, 1985; Gardin et al., 1995) and the Tercis section (France) varies in parts from the sequence of events given in the standard zonations by Perch-Nielsen (1985) and Burnett (1998), i.e. the LO of *U. trifidus* is below the LO of *B. parca constricta* both in the Bottaccione section (Gardin et al., 1995) and at Tercis (Barchi et al., 1997), but above the LO of *B. p. constricta* in the standard schemes of Perch-Nielsen (1985) and Burnett (1998).

Tercis-les-Bains: proposed Global Boundary Stratotype Section and Point for the lower boundary of the Maastrichtian

Biostratigraphic data from Hancock & Kennedy (1993), Hancock et al. (1993), Odin et al. (2001) and magnetostratigraphic data by Barchi et al. (1997) from the Tercis section demonstrate a complete Cam-

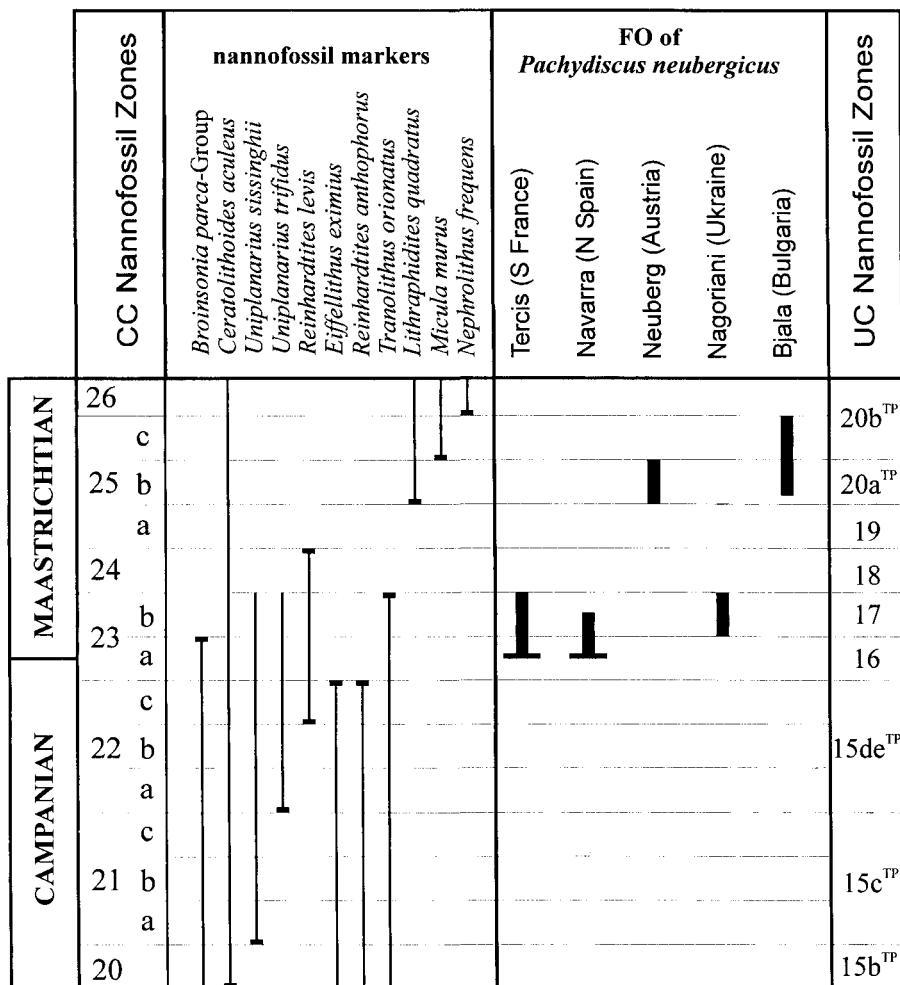


Fig. 2. Stratigraphical position of local first occurrences of the ammonite *Pachydiscus neubergicus* in European sections against the nannofossil standard zonations (CC zones) of Sissingh (1977) and Perch-Nielsen (1985), modified according to Wagreich & Krenmayr (1993), and the nannofossil zonation of Burnett (1998; UC zones). Standard CC (sub)zones are shown in equal durations.

panian-Maastrichtian boundary interval. Although preservation of nannofossils is poor, a detailed nannofossil stratigraphy was established by Barchi et al. (1997), differing slightly from results of former investigations by Burnett in Hancock et al. (1993).

The sequence of events within the Tercis section based on Barchi et al. (1997) is as follows:

FO *Ceratolithoides aculeus*, FO *Uniplanarius gothicum* (= *U. sissinghii* in the present paper), FO *U. trifidus*, LO *Eiffellithus eximus*, LO nannoconids, LO *U. gothicum*, LO *U. trifidus*, LO *Tranolithus phacelosus* (= *T. orionatus* in the present paper) and *B. parca*.

Nannofossil assemblages were interpreted by Barchi et al. (1997) to indicate the predominance of Tethyan taxa; typically high temperate/boreal taxa are not present.

The FO of *Pachydiscus neubergicus* according to Odin (in Christensen et al., 2000; see also Odin et al., 2001) is at level 115.7 m; this is practically contemporaneous with the LO of forms referred either to *P. perfidus* or to an intermediate *P. perfidus/neubergicus*. This event lies above the LO of *Uniplanarius gothicum*

(96 m) and below the LO of *Uniplanarius trifidus* (129 m) and *Broinsonia parca* (144 m) and may be correlated with the standard nannofossil zone CC 23a of Perch-Nielsen (1985) or UC16 of Burnett (1998).

Correlations of these biostratigraphic data for the Tercis section to the magnetostratigraphic scale indicate a long interval of normal polarity (C33n) in the lower part followed by a short reversed polarity zone (C32r2; Barchi et al., 1997). The C32r2 transition appears considerably below the FO of *P. neubergicus* (80–81 m). According to Barchi et al. (1997), this level can be correlated to the *Globotruncanella havanensis* Zone of the Tethyan foraminiferal zonation of the Bottaccione section (Monechi & Thierstein, 1985; Premoli Silva & Sliter, 1994) and the lower part of the *Belemnitella langei* Zone of the boreal Lägerdorf-Kronsmoor section (northern Germany; see Schönfeld et al., 1996), thus indicating a late Late Campanian age for this transition.

The type locality of *Pachydiscus neubergicus*

The type material of *Pachydiscus neubergicus* was restudied by Kennedy & Summesberger (1986). At

the type locality, a disused quarry near Krampen (Neuberg an der Mürz, Austria), basal conglomerates and sandstones without ammonites grade into marly sandstones containing *P. neubergicus*, *Pachydiscus epiplectus* (Redtenbacher, 1873) and *Hoploscaphites constrictus* (J. Sowerby, 1817). Nannofossil samples from the Neuberg section have yielded a nannoflora including *Lithraphidites quadratus*, which indicated standard nannofossil zone CC25b/UC20a, i.e. a 'late' Maastrichtian age according to nannofossil scales (e.g. Sissingh, 1977; Perch-Nielsen, 1985; Burnett, 1998), although the occurrence of *P. epiplectus* suggests an Early Maastrichtian age. Correlations with German sections by means of nannofossils (Kennedy & Summesberger, 1986) provided evidence for a middle Early Maastrichtian age (*Belemnella sumensis* Zone) of the type locality of *P. neubergicus*.

Campanian-Maastrichtian boundary sections in Navarra (northern Spain)

Sections in northern Spain provide important data for the correlation of Late Cretaceous macrofossil and microfossil scales, in view of the co-occurrence of ammonites, echinoids, inoceramids, foraminifera and calcareous nannoplankton (Küchler & Kutz, 1989; Kutz 1995; Wagreich et al., 1998; Küchler & Wagreich, 1999; Küchler et al., 2001). Several sections comprise a nearly continuous ammonite record across the Campanian-Maastrichtian boundary, including well-defined FOs of *P. neubergicus*. Two ammonite zones could be distinguished, a 'Late' Campanian *Nostoceras hyatti* Zone and an 'Early' Maastrichtian *Pachydiscus neubergicus/P. epiplectus* Zone. In the Erro section, the FO of *P. neubergicus* was found to overlie a *Nostoceras* bed, which could be assigned to nannofossil zone CC23a, characterised by the presence of *Broinsonia p. constricta* and *Uniplanarius trifidus* and the absence of both *Reinhardtites anthophorus* and *Eiffellithus eximius*. In Burnett's (1998) nannofossil zonation, this interval corresponds to zone UC16. In all the investigated localities in northern Spain, e.g. the Imiscoz, Juandechaco and Barranca sections (Erice II), the FO of *P. neubergicus* can be correlated consistently to a level within nannofossil zone CC23a/UC16.

The *Nostoceras hyatti* Bed, as defined in the Erro II section, lies near the base of CC23a, about 2 m above the LO of *Eiffellithus eximius*. Also in the Juandechaco and Erice II sections, at least portions of the latest Campanian *Nostoceras hyatti* Zone may be assigned to nannofossil zone CC23a.

Pachydiscus neubergicus at Bjala (Bulgaria) and Nagorianiy (the Ukraine)

Pachydiscus neubergicus has previously been recorded from sections at Bjala (Bulgaria; Ivanov & Stoykova, 1994; Rögl et al., 1995) and at Nagoriani (the Ukraine; Kennedy & Summesberger, 1987). In the Bjala section, *P. neubergicus* ranges up to nannofossil zone CC25b/c (*Micula murus* Zone of Ivanov & Stoykova, 1994), i.e. into the Upper Maastrichtian. The Bjala section also allows a correlation with planktonic foraminifera, indicating the LO of *P. neubergicus* to be within the *Abathomphalus mayaroensis* Zone (Rögl et al., 1995).

At Nagoriani, samples from ammonite specimens including *P. neubergicus*, indicated nannofossil zone CC23b (Wagreich, 1987). Similar to the type section of *P. neubergicus* in Austria, these sections provided evidence for a rather long stratigraphical range of this ammonite, with respect to nannofossil and foraminiferal zonations.

Nannofossil zones and the base of the Maastrichtian in European sections

Both at Tercis, the proposed Global Boundary Stratotype Section and Point for the base of the Maastrichtian, and in sections in northern Spain (between c. 50 and 80 km south of Tercis), *P. neubergicus* first appears within nannofossil zone CC23a/UC16 (Fig. 2). Nannofossil standard zone CC23a is characterised by the absence of *E. eximius* and *R. anthophorus* and the presence of *B. p. constricta* and *U. trifidus*. Nannofossil assemblages from both areas show Tethyan affinities, providing evidence for a palaeogeographic position along the northern Tethyan margin.

Nannofossil data from these sections and their correlation to other European key sections such as the Lägerdorf-Kronsmoor section (northern Germany; Schönfeld et al., 1996) yield convincing evidence for the assumption that the event defining the base of the Maastrichtian in the Boreal realm, i.e. the FO of the belemnite *Belemnella lanceolata* (von Schlottheim, 1813), defines a level within the same nannofossil zone, CC23a/UC16 (Schönfeld & Burnett, 1991). The nannofossil assemblages from the Lägerdorf-Kronsmoor sections are generally quite different from those found in Spain, because some low-latitude species (e.g., *U. trifidus*) are missing (see Wagreich et al., 1999). Nevertheless, the marker events are similar across this boundary interval, so that nannofossil zone CC23a could still be determined in these Boreal sections (Burnett, 1990, 1998; Schönfeld et al., 1996). Comparison to nannofossil data from the

Maastrichtian type area is hampered by a considerable hiatus across the Campanian/Maastrichtian boundary and by strong reworking in this area (Manivit in Robaszynski et al., 1985; Burnett, 1996; see also Jagt & Felder, 2003).

Our nannofossil data also support the conclusions reached by Schönfeld & Burnett (1991), Hancock et al. (1992) and Simmons et al. (1996), in that the extinction level of *Globotruncanita calcarata*, a widely used marker in Tethyan planktonic foraminiferal zonations (e.g., Caron, 1985), is considerably older than the base of the Maastrichtian as defined either by *P. neubergicus* or by *B. lanceolata*. The extinction of *G. calcarata* is commonly reported from nannofossil zone CC22b/c (Schönfeld & Burnett, 1991; Wagreich & Krenmayr, 1993), at least one zone below the FO of *P. neubergicus* as indicated by our data.

Correlations of nannofossil zones to other key sections containing *P. neubergicus* indicate considerable diachroneity in the local FO of this marker species. Evidence from the local first appearance at the type locality near Neuberg (Austria), as well as from Bjala (Bulgaria) and Nagorani (the Ukraine), indicate a stratigraphical range of this ammonite into the 'Late' Maastrichtian up to nannofossil zones CC25b/c (Wagreich in Kennedy & Summesberger 1986; Ivanov & Stoykova, 1994). The total known range of *P. neubergicus* thus spans at least five nannofossil subzones, starting from CC23a/UC16 and extending into CC25c/UC20b^{TP}. As a consequence of the long stratigraphical range of *P. neubergicus*, local first appearances in non-continuous ammonite sections may be diachronous. It is recommended that checks be made against other fossil groups, or a magnetostratigraphical scale, before age interpretations are attempted.

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